Dynamics 2
Lecturer: Prof. Dr. G. Lohmann
Due date: 19.4.2020
Exercise 1, Summer semester 2021 Tutors: Justus Contzen, Lars Ackermann 12.4.2020

## 1. Scaling of the dynamical equations (2 points)

We work in the rotating frame of reference of the Earth. The equation can be scaled by a length-scale L, determined by the geometry of the flow, and by a characteristic velocity U . We can estimate the relative contributions in units of $\mathrm{m} / \mathrm{s}^{2}$ in the horizontal momentum equations:

$$
\begin{equation*}
\underbrace{\frac{\partial \mathbf{v}}{\partial t}}_{U / T \sim 10^{-8}}+\underbrace{\mathbf{v} \cdot \nabla \mathbf{v}}_{U^{2} / L \sim 10^{-8}}=\underbrace{-\frac{1}{\rho} \nabla p}_{\delta \mathbf{P} /(\rho \mathbf{L}) \sim 1 \mathbf{0}^{-5}}+\underbrace{2 \boldsymbol{\Omega} \times \mathbf{v}}_{\mathbf{f}_{0} \mathbf{U} \sim \mathbf{1 0 ^ { - 5 }}}+\underbrace{\text { fric }}_{\nu U / H^{2} \sim 10^{-13}} \tag{1}
\end{equation*}
$$

where fric denotes the contributions of friction due to eddy stress divergence (usually $\sim \nu \nabla^{2} \mathbf{v}$ ). Typical values are given in Table 1. The values have been taken for the ocean.
a) Please repeat the estimate for the atmosphere using Table 1 .
b) The Rossby number Ro is the ratio of inertial (the left hand side in (1)) to Coriolis (second term on the right hand side in (11) terms

$$
\begin{equation*}
R o=\frac{\left(U^{2} / L\right)}{(f U)}=\frac{U}{f L} \tag{2}
\end{equation*}
$$

Ro is small when the flow is in a so-called geostrophic balance. Please calculate Ro for the atmosphere and ocean using Table 1 .

|  | Quantity | Atmosphere | Ocean |
| :---: | :---: | :---: | :---: |
| horizontal velocity | U | $10 \mathrm{~ms}^{-1}$ | $10^{-1} \mathrm{~ms}^{-1}$ |
| horizontal length | L | $10^{6} \mathrm{~m}$ | $10^{6} \mathrm{~m}$ |
| horizonal Pressure changes | $\delta \mathrm{P}$ (horizontal) | $10^{3} \mathrm{~Pa}$ | $10^{4} \mathrm{~Pa}$ |
| time scale | T | $10^{5} \mathrm{~s}$ | $10^{7} \mathrm{~s}$ |
| Coriolis parameter at $45^{\circ} \mathrm{N}$ | $f_{0}=2 \Omega \sin \varphi_{0}$ | $10^{-4} \mathrm{~s}^{-1}$ | $10^{-4} \mathrm{~s}^{-1}$ |
| density | $\rho$ | $1 \mathrm{kgm}^{-3}$ | $10^{3} \mathrm{kgm}^{-3}$ |
| viscosity (turbulent) | $\nu$ | $10^{-5} \mathrm{kgm}^{-3}$ | $10^{-6} \mathrm{kgm}^{-3}$ |

Table 1: Table shows the typical scales in the atmosphere and ocean system.

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## 2. Concept of dynamic similarity (3 points)

For the case of an incompressible flow, assuming the temperature effects are negligible and external forces are neglected, the Navier-Stokes equations consist of conservation of mass

$$
\begin{equation*}
\nabla \cdot \mathbf{u}=0 \tag{3}
\end{equation*}
$$

and conservation of momentum

$$
\begin{equation*}
\partial_{t} \mathbf{u}+(\mathbf{u} \cdot \nabla) \mathbf{u}=-\frac{1}{\rho_{0}} \nabla p+\nu \nabla^{2} \mathbf{u} \tag{4}
\end{equation*}
$$

where $\mathbf{u}$ is the velocity vector and p is the pressure, $\nu$ denotes the kinematic viscosity.
a) Show: The equations (3|4) can be made dimensionless by a length-scale L , determined by the geometry of the flow, and by a characteristic velocity U. For example: $u=U \cdot u_{d}$.
Note: the units of $\left[\rho_{0}\right]=k g / m^{3},[p]=k g /\left(m s^{2}\right)$, and $[p] /\left[\rho_{0}\right]=m^{2} / \mathrm{s}^{2}$. Therefore the pressure gradient term in (4) has the scaling $U^{2} / L$.
b) Show: The scalings vanish completely in front of the terms except for the $\nabla^{2} \mathbf{u}_{\mathrm{d}}{ }^{-}$ term! The dimensionless parameter is the Reynolds number and the only parameter left!

Remark: For large Reynolds numbers, the flow is turbulent. In most practical flows Re is rather large $\left(10^{4}-10^{8}\right)$, large enough for the flow to be turbulent.

## 3. Advection (3 points)

A ship is steaming northward at a rate of $10 \mathrm{~km} / \mathrm{h}$. The surface pressure increases toward the northwest at a rate of $5 \mathrm{~Pa} / \mathrm{km}$. What is the pressure tendency recorded at a nearby island station if the pressure aboard the ship decreases at a rate of $100 \mathrm{~Pa} / 3 \mathrm{~h}$ ?
4. Download and install the $\mathbf{R}$ version for your operating system (for many linux distributions R is also available in the package management system). Furthermore, look at the web page for R studio http://www.rstudio.com/, R studio is a free and open source user interface for R. One particular package is Shiny. This makes it super simple for R users like you to turn analyses into interactive web applications that anyone can use. The latest version of R for Linux, OS X and Windows is freely available on the CRAN webpage: http://cran.r-project.org (Fig. 1).


Figure 1: R is available for download from the CRAN webpage: http://cran.r-project.org.

## 5. Short programming questions. (2 points)

Write down the output for the following R-commands:
a) $\mathrm{a}<-\mathrm{c}(0,-5,4,20)$; mean (a)
b) $\max (a)-\min (a)$
c) $a * 2+c(3,1,-1,0)$
f) Plot the potential

```
y=-100:100
x=y/50
r=1
z=-r * x^2/2 + r * x^3/3
plot(x,z,type='l')
```

and the derivative of $\mathrm{z}(\mathrm{x})$

Notes on submission form of the exercises: Working in study groups is encouraged, but each student is responsible for his/her own solution. The answers to the questions can be send until the due date (12:00) to Justus Contzen (Justus.Contzen@awi.de), Lars Ackermann (Lars.Ackermann@awi.de).

