(4)

## 1. Wind-driven ocean circulation (6 points)

The **Sverdrup transport** V for the depth-integrated flow is calculated by

$$\rho_0 \beta V = \frac{\partial}{\partial x} \tau_y - \frac{\partial}{\partial y} \tau_x \tag{1}$$

where  $\tau_x$  and  $\tau_y$  are the components of the wind stress.

The **Ekman transports**  $V_E, U_E$  describe the dynamics in the upper mixed layer:  $fV_E = -\tau_x/\rho_0$ ,  $fU_E = \tau_y/\rho_0$  (2)

where  $U_E = \int_{-E}^{0} u dz$  and  $V_E = \int_{-E}^{0} v dz$  are the depth-integrated velocities in the thin friction-dominated Ekman layer at the sea surface.

**Ekman vertical velocity**  $w_E$ : Using the continuity equation, the divergence of the Ekman transports leads to a vertical velocity  $w_E$  at the bottom of the Ekman layer:

$$w_E = -\int_{-E}^{0} \frac{\partial w}{\partial z} dz = \frac{\partial}{\partial x} U_E + \frac{\partial}{\partial y} V_E = \frac{\partial}{\partial x} \left( \frac{\tau_y}{\rho_0 f} \right) - \frac{\partial}{\partial y} \left( \frac{\tau_x}{\rho_0 f} \right) \quad . \tag{3}$$

a) Assume that the windstress is only zonal with  $\tau_x = -\tau_0 \cos(\pi y/B)$  for an ocean basin with 0 < x < L, 0 < y < B.

Calculate the Sverdrup transport, Ekman transports, and Ekman pumping velocity for this special case.

b) Make a schematic diagram of the windstress, Sverdrup transport, Ekman transports, and Ekman pumping velocity.

c) Using a), at what latitudes y are |V| and  $|V_E|$  maximum? Calculate their magnitudes. Take constant  $f = 10^{-4} \text{ s}^{-1}$  and  $\beta = 1.8 \cdot 10^{-11} \text{ m}^{-1} \text{s}^{-1}$  and  $B = 5000 \text{ km}, \tau_0/\rho_0 = 10^{-4} \text{ m}^2 \text{s}^{-2}$ .

d) Using the values in b), calculate the maximum of  $w_E$  for constant f.

## 2. Angular momentum and Hadley cell (6 points)

Consider a zonally symmetric circulation (i.e., one with no longitudinal variations) in the atmosphere. In the inviscid upper troposphere one expects such a flow to conserve absolute angular momentum, i.e.,

$$\frac{DA}{Dt} = 0, (5)$$

| Dynamics II, Summer semester 2025             | Exercise 6          |
|---|---------------------|
| Lecturer: Prof. Dr. G. Lohmann, Dr. M. Ionita | 19.5.2025           |
| Tutors: Alessandro Gagliardi, Georg Hüttner   | Due date: 26.5.2025 |

where A is the absolute angular momentum per unit mass (parallel to the Earth's rotation axis)

$$A = r\left(u + \Omega r\right) = \Omega R^2 \cos^2 \varphi + uR \cos \varphi \quad . \tag{6}$$

 $\Omega$  is the Earth rotation rate, u the eastward wind component,  $r = R \cos \varphi$  is the distance from the rotation axis, R the Earth's radius, and  $\varphi$  latitude.

a) Show, for inviscid zonally symmetric flow, that the relation  $\frac{DA}{Dt} = 0$  is consistent with the zonal component of the equation of motion

$$\frac{Du}{Dt} - fv = -\frac{1}{\rho} \frac{\partial p}{\partial x} \tag{7}$$

in (x, y, z) coordinates, where  $y = R\varphi$ .

b) Use angular momentum conservation to describe in words how the existence of the Hadley circulation explains the existence of both the subtropical jet in the upper troposphere and the near-surface trade winds.

c) If the Hadley circulation is symmetric about the equator, and its edge is at  $20^{\circ}$  latitude, determine the strength of the subtropical jet. Use (5, 6).

d) Is the Hadley Cell geostrophically driven or not?

<u>Notes on submission form of the exercises:</u> Working in study groups is encouraged, but each student is responsible for his/her own solution. The answers to the questions can be send until the due date (12:00 pm) to Alessandro Gagliardi (Alessandro.Gagliardi@awi.de), Georg Huettner (Georg.Huettner@awi.de).