## 1. Ocean thermohaline circulation (3 points)

Consider a geostrophic flow (u, v)

$$-fv = -\frac{1}{\rho_0} \frac{\partial p}{\partial x} \tag{1}$$

$$fu = -\frac{1}{\rho_0} \frac{\partial p}{\partial y} \quad . \tag{2}$$

The meridional overturning stream function  $\Phi(y, z)$  is defined via

$$\Phi(y,z) = \int_0^z \frac{\partial \Phi}{\partial \tilde{z}} d\tilde{z}$$
(3)

$$\frac{\partial \Phi}{\partial \tilde{z}} = \int_{x_e}^{x_w} v(x, y, \tilde{z}) \, dx \quad \text{(zonally integrated transport)}, \tag{4}$$

where  $x_e$  and  $x_w$  are the eastward and westward boundaries in the ocean basin (think e.g. of the Atlantic Ocean). Units of  $\Phi$  are  $m^3 s^{-1}$ . At the surface  $\Phi(y, 0) = 0$ .

Calculate  $\Phi(y, z)$  as a function of density  $\rho$  at the basin boundaries! Use

$$v(x, y, z) = \frac{1}{\rho_0 f} \frac{\partial p}{\partial x}$$

and the hydrostatic approximation

$$\frac{\partial p}{\partial z} = -g\rho$$

## 2. Laplace transform (2 points)

is given by 
$$\mathcal{L}\left\{x(t)\right\} = L(s) = \int_0^\infty e^{-st} x(t) dt$$
 (5)

a) Show that

$$\mathcal{L}\left\{\frac{d}{dt}x(t)\right\} = sL(s) - x(0) \tag{6}$$

through integration by parts.

b) Show furthermore that

$$\mathcal{L}\left\{\exp(-at)\right\} = \frac{1}{s+a} \tag{7}$$

Dynamics II, Summer semester 2023	Exercise 7
Lecturer: Prof. Dr. G. Lohmann, Dr. M. Ionita	12.6.2023
Tutors: Fernanda Matos, Ahmadreza Masoum	Due date: 19.6.2023

3. Laplace transformation of mixed layer model (5 points) Imagine that the temperature of the ocean mixed layer is governed by

$$\frac{dT}{dt} = -\lambda T + Q(t), \qquad (8)$$

where  $\lambda$  is the typical damping rate of a temperature anomaly and Q(t) a forcing.

(a) Use the Laplace transformation to show

$$L(s) = \frac{Q(s) + T(0)}{s + \lambda} \quad . \tag{9}$$

where  $Q(s) = \mathcal{L} \{Q(t)\}$ 

(b) Consider the special case  $Q(t) = \exp(i\omega_0 t)$ , then  $Q(s) = \frac{1}{s-i\omega_0}$ . The forcing and the temperature is of course a real number, but by representing it as a complex number we can simultaneously keep track of both phase components. Show that

$$L(s) = \frac{T(0) + Q(s)}{s + \lambda} = \frac{T(0)}{s + \lambda} + \frac{1}{(s + \lambda)} \frac{1}{(s - i\omega_0)}$$
(10)

and via the Laplace back-transformation and (7) of the exercise above as well as (15) that

$$T(t) = \exp(-\lambda t)T(0) + \frac{\left[\exp(i\omega_0 t) - \exp(-\lambda t)\right]}{\lambda + i\omega_0} \quad . \tag{11}$$

- (c) Calculate the real and imaginary part of (11).
- (d) Take the real part. Show: At low frequencies, the output T(t) is similar to the forcing Q(t). At high frequencies it rolls off as  $1/\omega$  (it is a low-pass filter) and is out of phase by 90°.
- (e) Instead of b), consider now the special case  $Q(t) = c \cdot u(t)$  with u(t) as unit step or the so-called Heaviside step function (step forcing). Show via Laplace transform that

$$\langle T(t) \rangle = \mathcal{L}^{-1}\{L(s)\}(t) = \mathcal{L}^{-1}\left\{\frac{\langle T(0) \rangle}{s+\lambda} + \frac{c}{s} \cdot \frac{1}{s+\lambda}\right\}$$
(12)

$$= T(0) \cdot \exp(-\lambda t) + \frac{c}{\lambda} \left(1 - \exp(-\lambda t)\right)$$
(13)

the equilibrium response

$$\Delta T = \lim_{t \to \infty} \langle T(t) \rangle = \frac{c}{\lambda} \,. \tag{14}$$

Hint:

$$\mathcal{L}\left\{-\exp(-at) + \exp(-bt)\right\} = \frac{-1}{s+a} + \frac{1}{s+b} = \frac{a-b}{(s+a)}\frac{1}{(s+b)}$$
(15)

<u>Notes on submission form of the exercises:</u> Working in study groups is encouraged, but each student is responsible for his/her own solution. The answers to the questions can be send until the due date (12:00) to Fernanda Matos (Fernanda.Matos@awi.de), Ahmadreza Masoum (Ahmadreza.Masoum@awi.de).