

**10:00-11:15 Overview with essential questions**

- 1) Astrid Bracher, Ocean Optics and Ocean Color Remote Sensing
- 2) Gerrit Lohmann, Dynamics II
- 3) Christoph Völker, Global Carbon Cycle
- 4) Gunnar Spreen: Remote Sensing of Ocean and Cryosphere /remote sensing of sea ice, polar regions
- 5) Mihalis Vrekoussis (Mathematics/Python)
- 6) Mihalis Vrekoussis Atm. Chemistry together with John P. Burrows,

**11:15-11:30 Break with Posters**

**11:30-12:45 continue with essential questions**

- 7) Atmospheric Modeling
- 8) Christian Melsheimer: Microwave Remote Sensing
- 9) Alexandra Klemme: Isotopes in Environmental Physics (elective)
- 10) John Burrows Atmospheric Physics
- 12) Gerrit Lohmann, Climate II

**12:45-13:30 Break with Posters and food**

**13:30-14:00 Tour through the building**

**14:00-15:00 Developing ideas in breakout groups**

# Basics of Environmental Physics and Climate

# *Aim of meeting, overview of key activities, essence*

**How does the Earth System evolve in a warming climate ?**

**Environmental Changes**

**Drivers, Thresholds, Variability & Extremes, ...**

# **Aim of meeting, overview of key activities, essence**

**How does the Earth System evolve in a warming climate ?**

**Environmental Changes**

**Drivers, Thresholds, Variability & Extremes, ...**

**Common understanding & Framework**

**Suitable diversity for interdisciplinary research & teaching**

**What can be done together?**

**Developments and gaps**





# Dynamics II

Gerrit Lohmann, Monica Ionita

- underlying dynamics of the atmosphere-ocean system
- The fundamental concepts of atmosphere-ocean flow, energetics, vorticity, and waves
- dynamical equations, climate data, basic physical concepts
- Exercises

$$\rho \left( \underbrace{\frac{\partial \mathbf{u}}{\partial t}}_{\text{Unsteady acceleration}} + \underbrace{\mathbf{u} \cdot \nabla \mathbf{u}}_{\text{Advective acceleration}} \right) = \underbrace{-\nabla p}_{\text{Pressure gradient}} + \underbrace{\mu \nabla^2 \mathbf{u}}_{\text{Viscosity}} + \underbrace{\mathbf{F}}_{\text{Other body forces}} .$$

The equation is annotated with physical interpretations: a large bracket above the left side labels it as 'Inertia (per volume)', a large bracket above the right side labels it as 'Divergence of stress', and individual brackets below each term provide further labels: 'Unsteady acceleration' for  $\frac{\partial \mathbf{u}}{\partial t}$ , 'Advective acceleration' for  $\mathbf{u} \cdot \nabla \mathbf{u}$ , 'Pressure gradient' for  $-\nabla p$ , 'Viscosity' for  $\mu \nabla^2 \mathbf{u}$ , and 'Other body forces' for  $\mathbf{F}$ .

# Fundamental aspects: Dynamics

**Goal: replace physical parameters with dimensionless numbers, which completely determine the dynamical behavior**

representative values for velocity ( $U$ ), time ( $T$ ), distances ( $L$ )

Using these values, the values in the dimensionless-system (written with subscript d) can be defined:

$$u = U \cdot u_d$$

$$t = T \cdot t_d$$

$$x = L \cdot x_d$$

with  $U = L/T$ .

From these scalings, we can also derive

$$\partial_t = \frac{\partial}{\partial t} = \frac{1}{T} \cdot \frac{\partial}{\partial t_d}$$

$$\partial_x = \frac{\partial}{\partial x} = \frac{1}{L} \cdot \frac{\partial}{\partial x_d}$$

$$\frac{\partial}{\partial t_d} \mathbf{u}_d + (\mathbf{u}_d \cdot \nabla_d) \mathbf{u}_d = -\nabla_d p_d + \frac{1}{Re} \nabla_d^2 \mathbf{u}_d \quad (1.27)$$

The dimensionless parameter  $Re = UL/\nu$  is the Reynolds number and the only parameter left!

# CIRCULATION

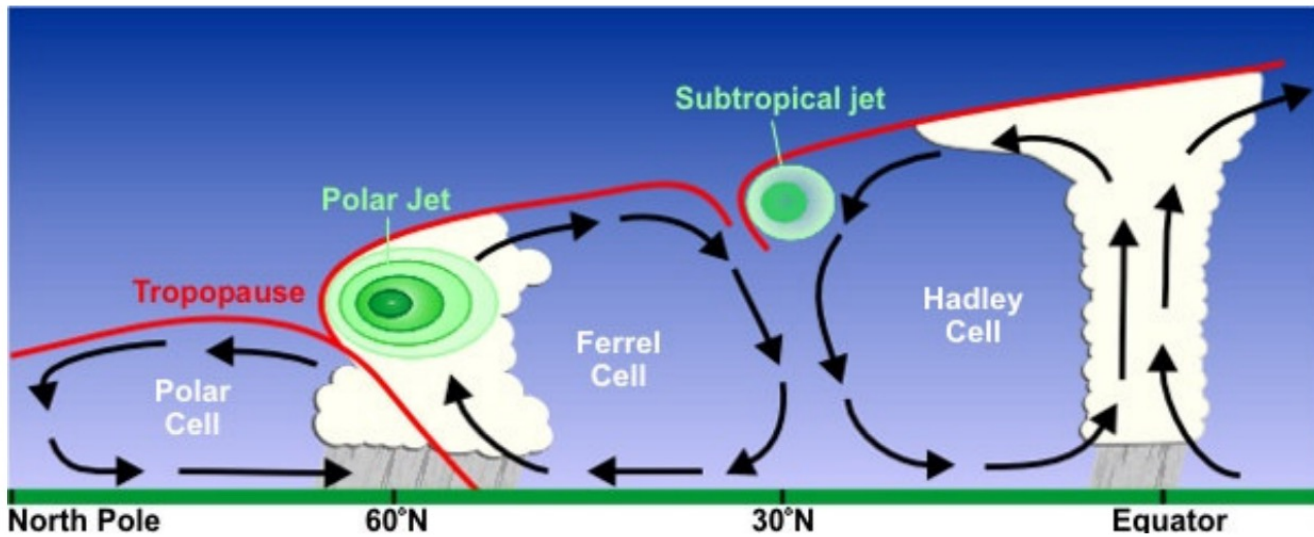


Figure 4.1: Cross section of the subtropical and polar jet streams by latitude.

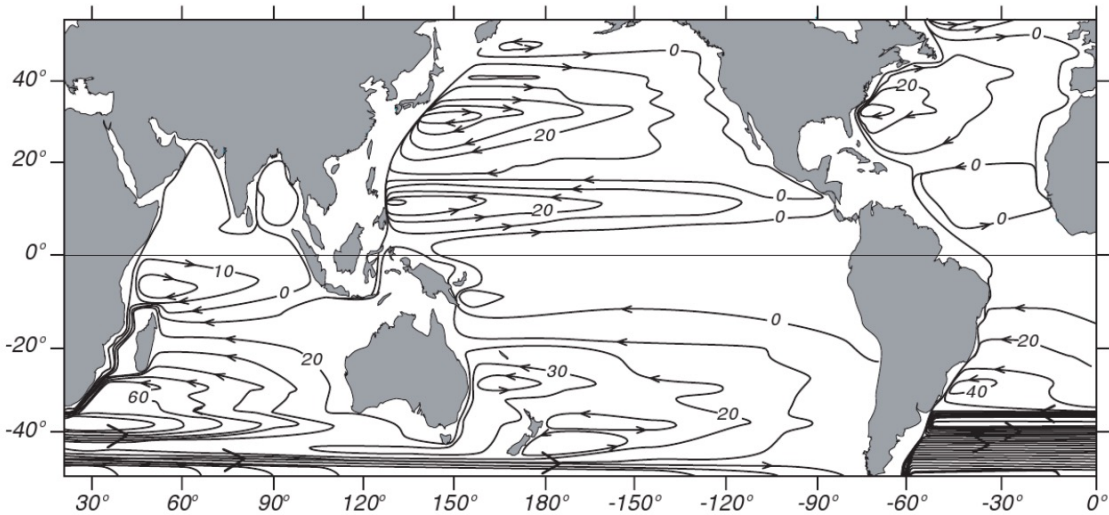


Figure 5.1: Depth-integrated Sverdrup transport applied globally using the wind stress from Hellerman and Rosenstein (1983). Contour interval is 10 Sverdrups (Tomczak and Godfrey, 1994).

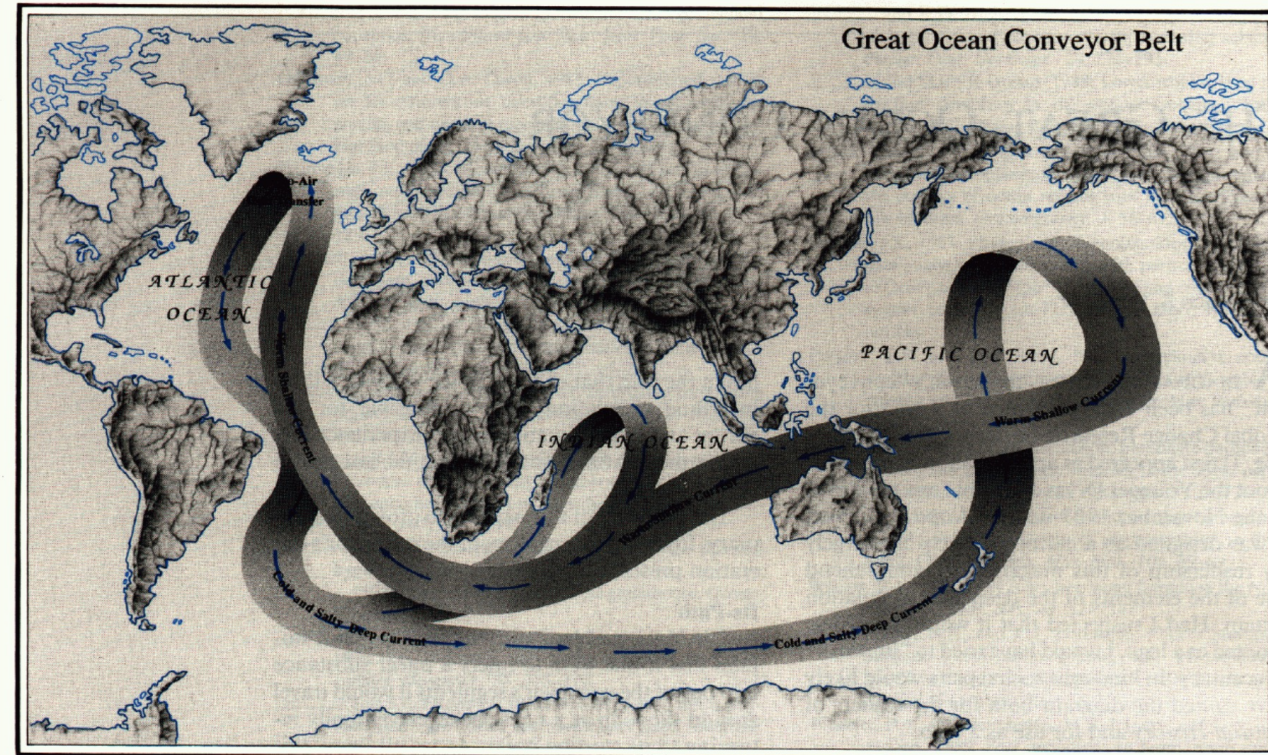
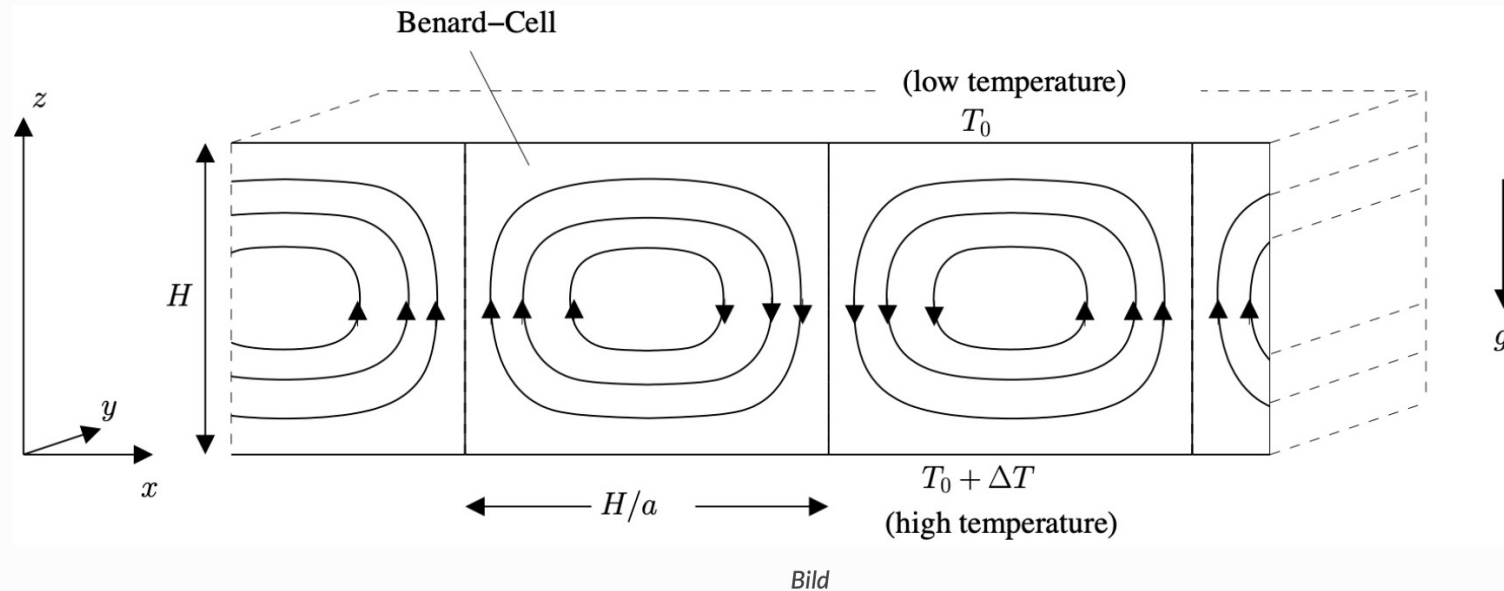


Fig. 1: The great ocean conveyor logo (Broecker, 1987). (Illustration by Joe Le Monnier, Natural History Magazine.)



# Fundamental aspects: Dynamics

## Convection in the Rayleigh-Benard system



Rayleigh (1916) temperature difference between the upper- and lower-surfaces

$$T(x, y, z = H) = T_0$$

$$T(x, y, z = 0) = T_0 + \Delta T$$

# Rossby waves atmosphere

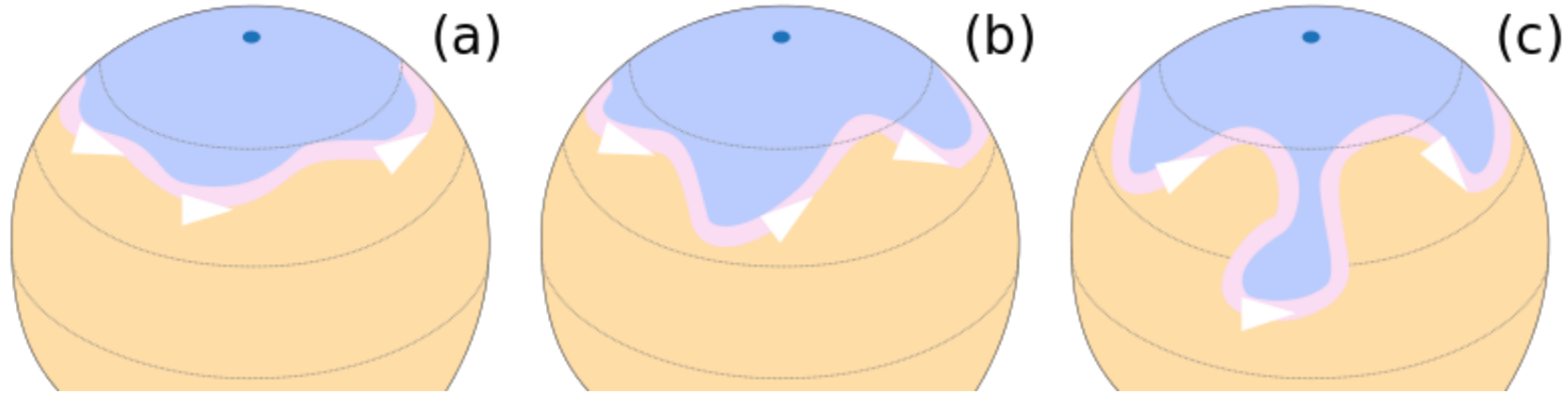


Figure 7.1: Meanders (Rossby Waves) of the Northern Hemisphere's polar jet stream developing (a), (b); then finally detaching a "drop" of cold air (c). Orange: warmer masses of air; pink: jet stream.

$$\partial_t u = f v - g \partial_x \eta$$

$$\partial_t v = -f u - g \partial_y \eta$$

$$\partial_t \eta = -\partial_x (H u) - \partial_y (H v)$$

# open shallow2D\_rossby.R

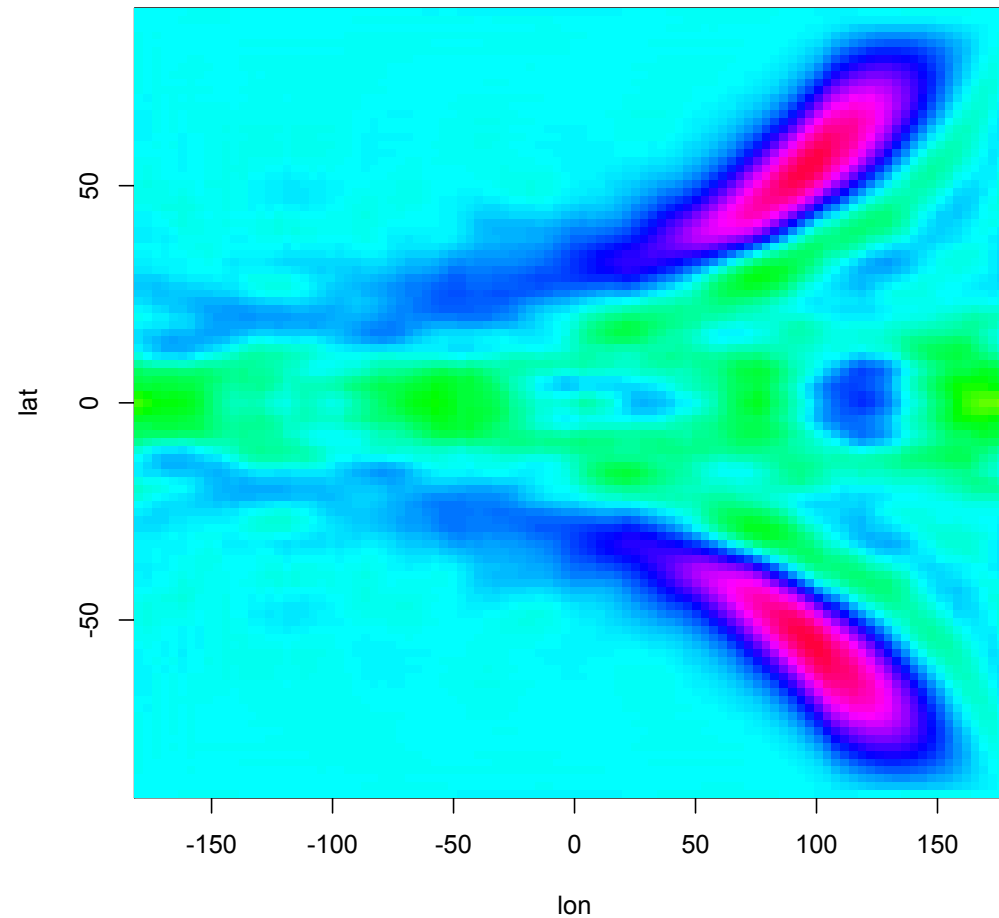
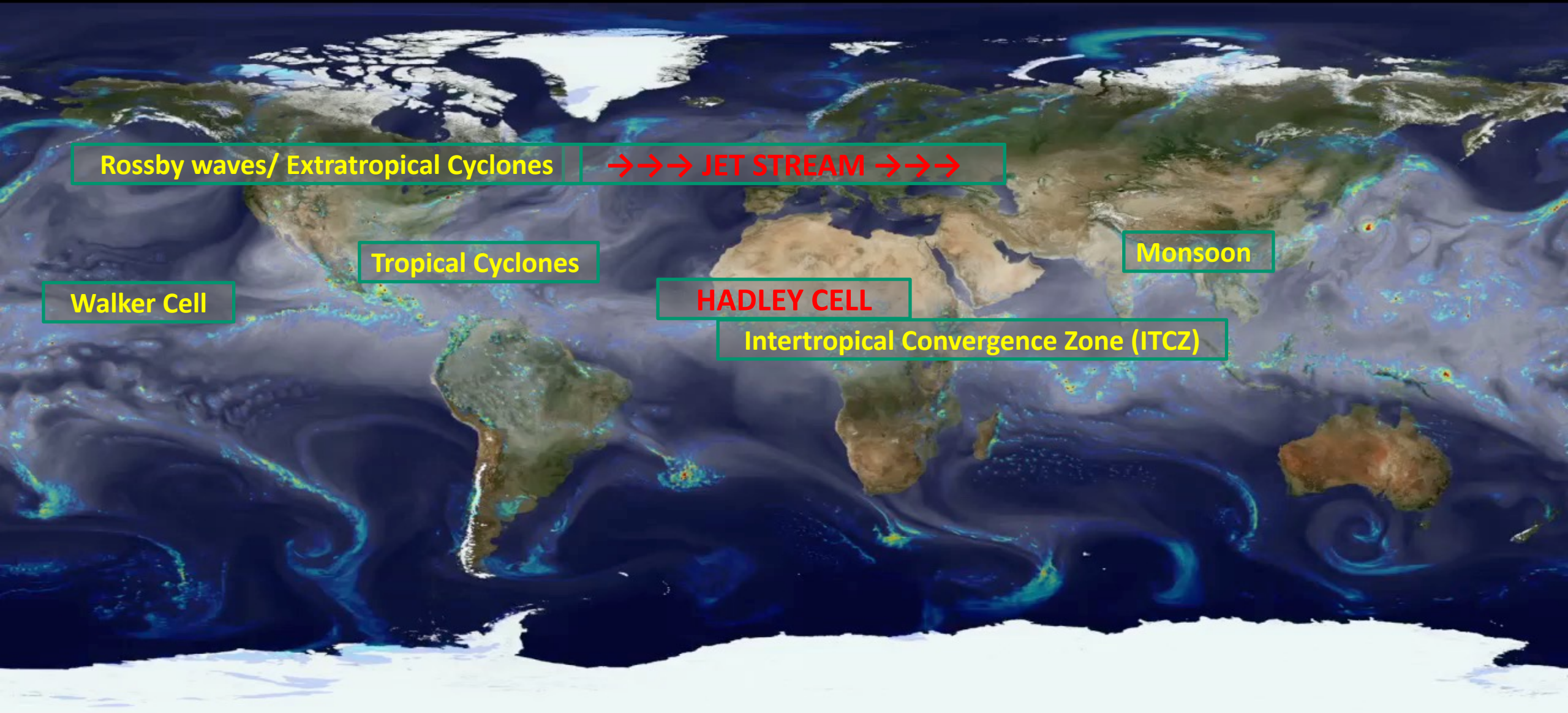
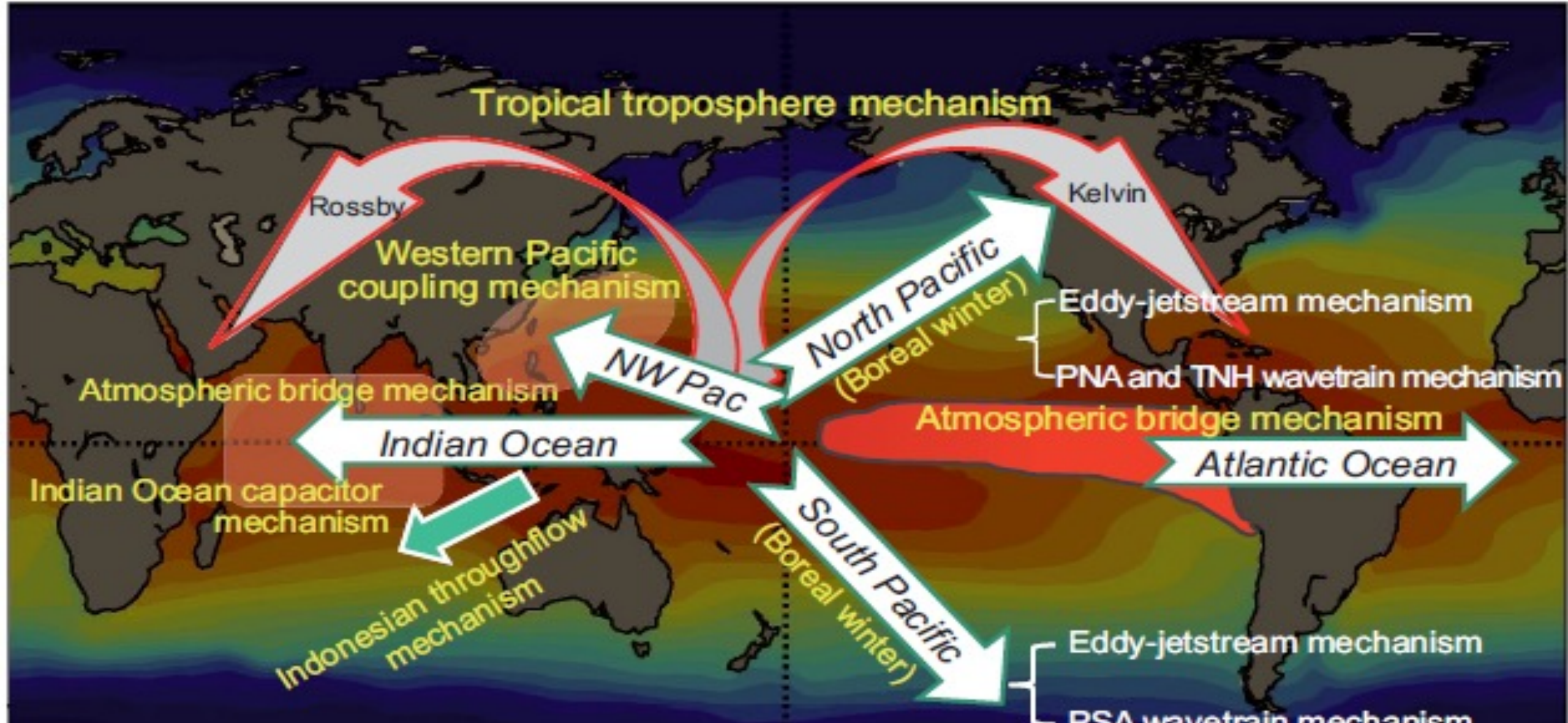


Figure 7.3: Global Rossby and Kelvin wave signatures in the exercise [49](#).

# Large scale atmospheric circulation







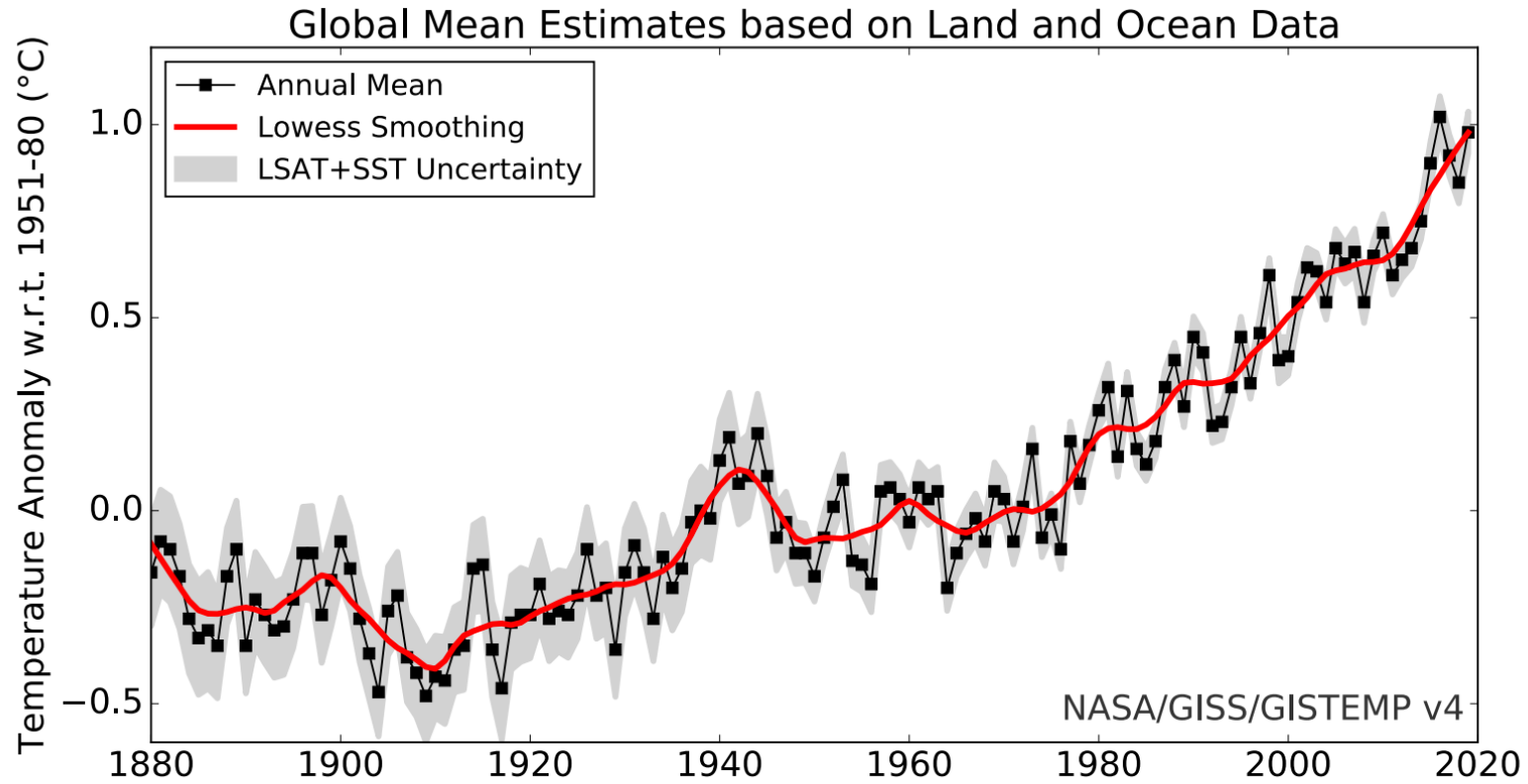
**How to analyze mode of variability and their relationship with our climate?!**





# Climate II

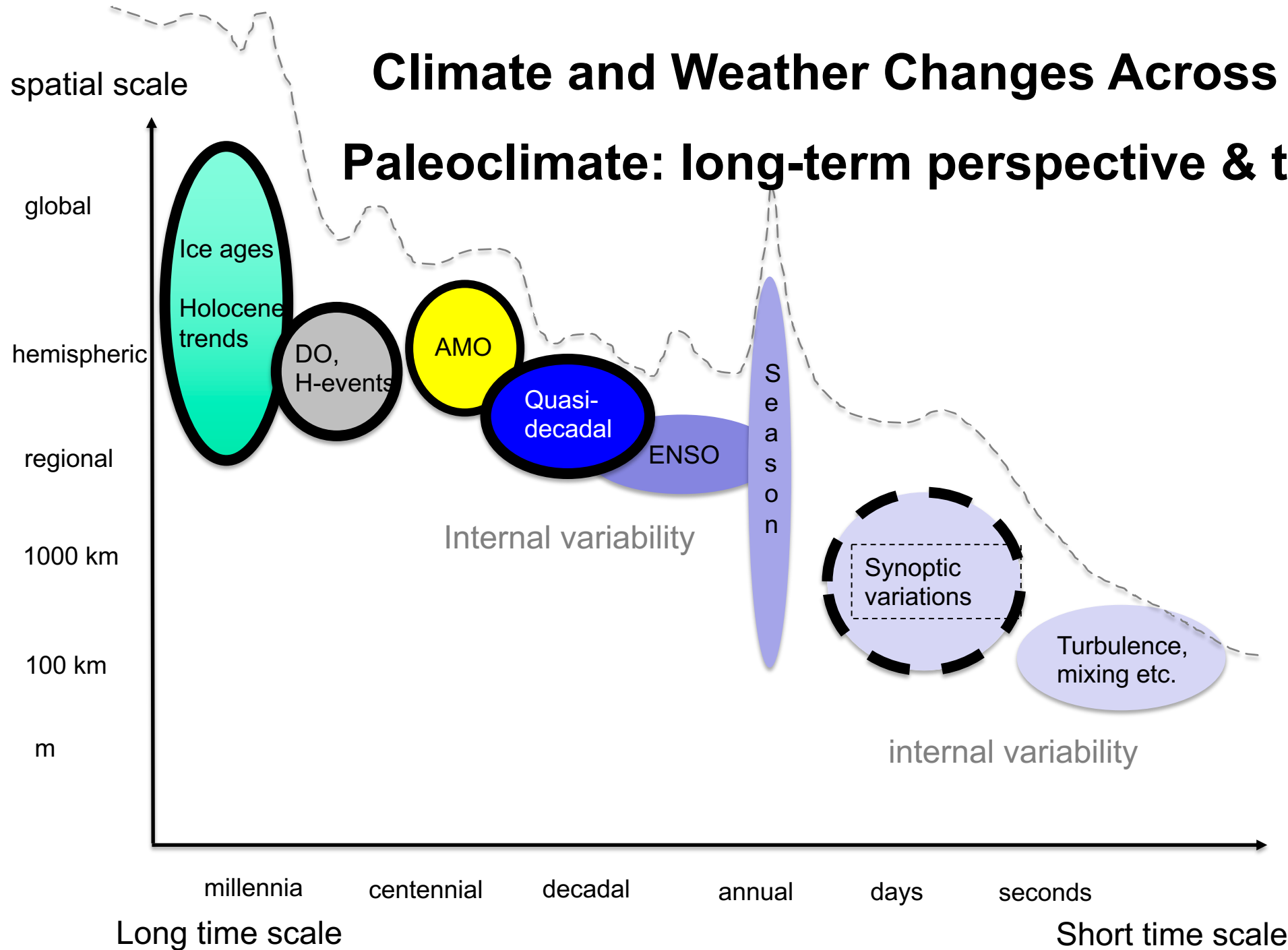
Gerrit Lohmann & Martin Werner



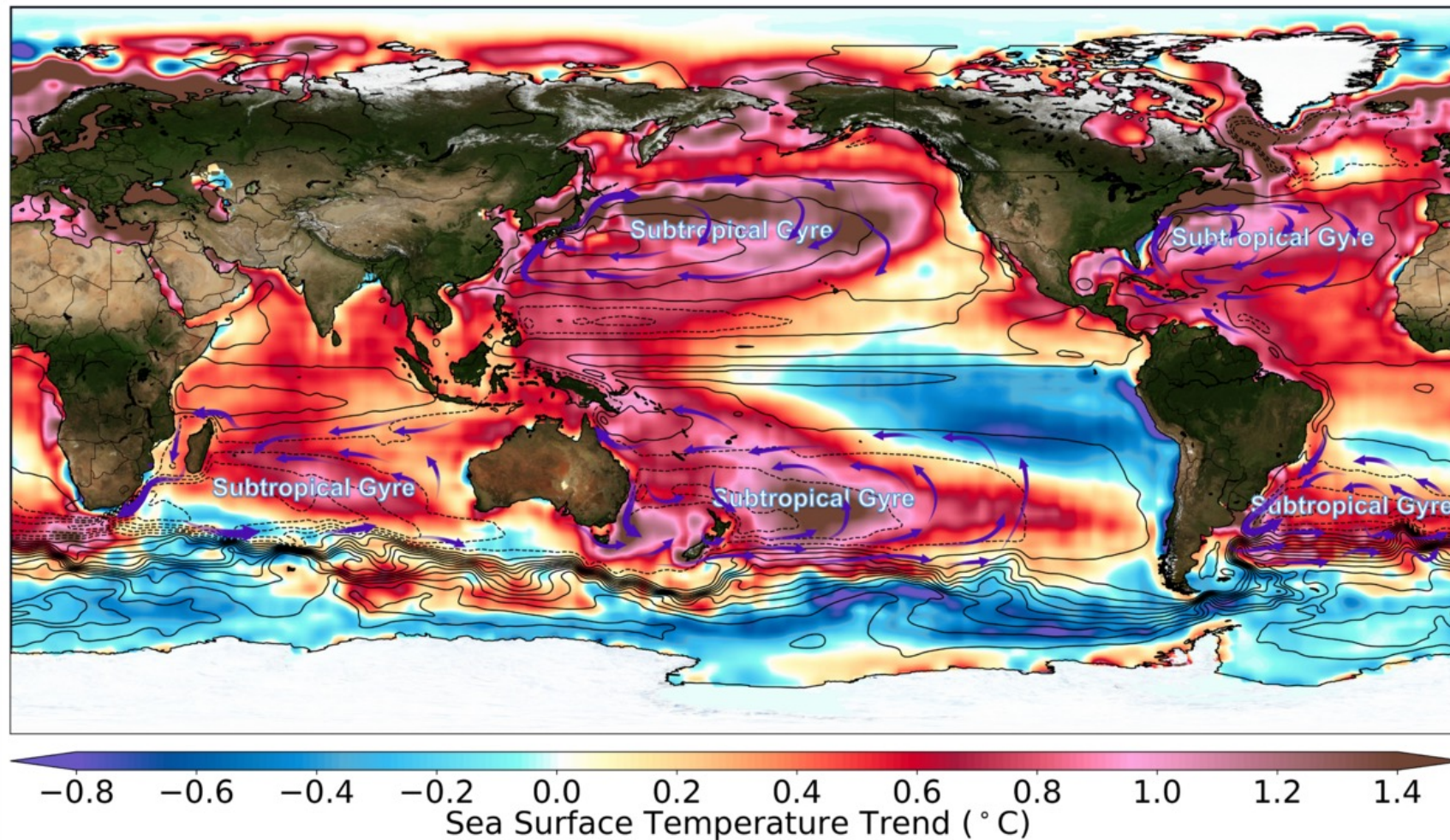
[https://data.giss.nasa.gov/gistemp/graphs\\_v4/](https://data.giss.nasa.gov/gistemp/graphs_v4/)

# Climate and Weather Changes Across Time Scales

## Paleoclimate: long-term perspective & test of models



# Temperature trend of the last 40 years

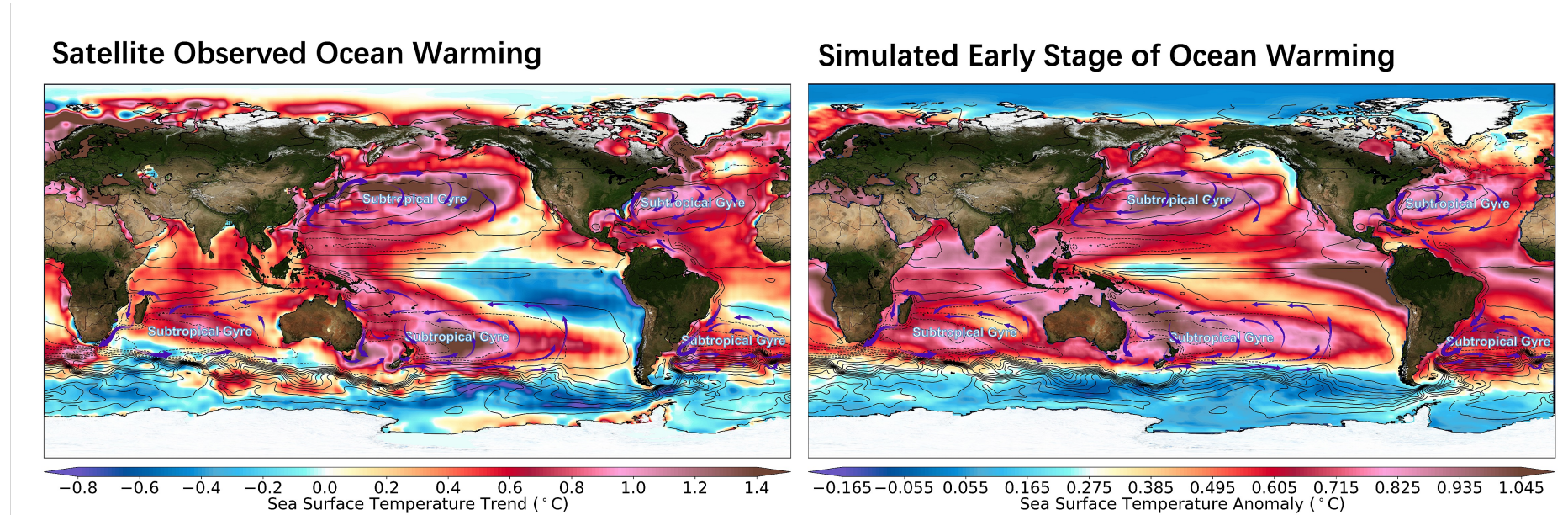


Information  
based on  
satellites

Yang, G. Lohmann, ... & J. Müller, 2023: The emergent pattern of infant stage ocean warming in satellite measurements. Communications Earth & Environment



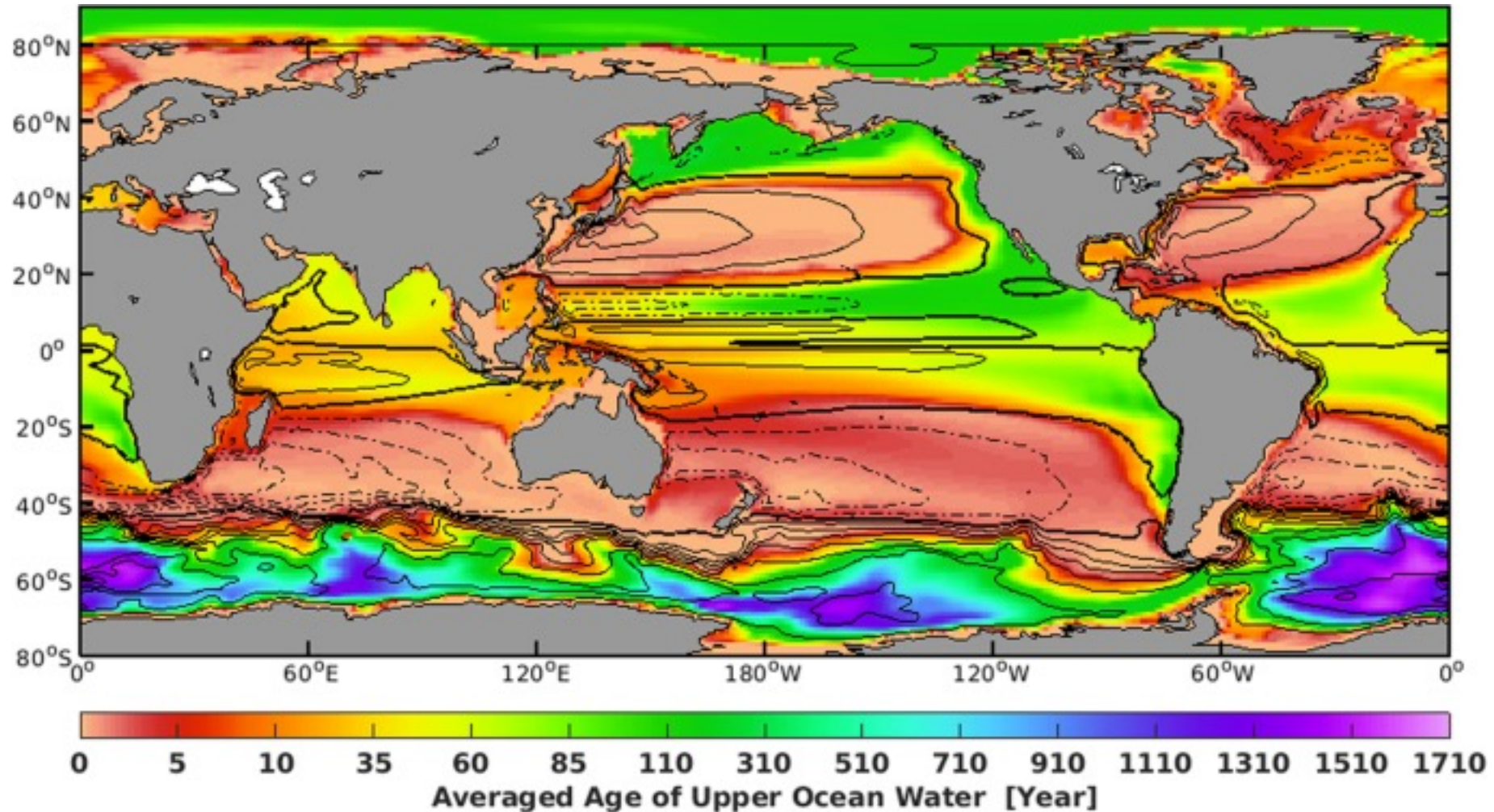
# Satellite-observed strong subtropical ocean warming as an early signature of global warming



AWI-ESM:  
initial shock

observed warming pattern is likely a short-term transient response to the increased CO<sub>2</sub> forcing, which only emerges during the early stage of anthropogenic warming.

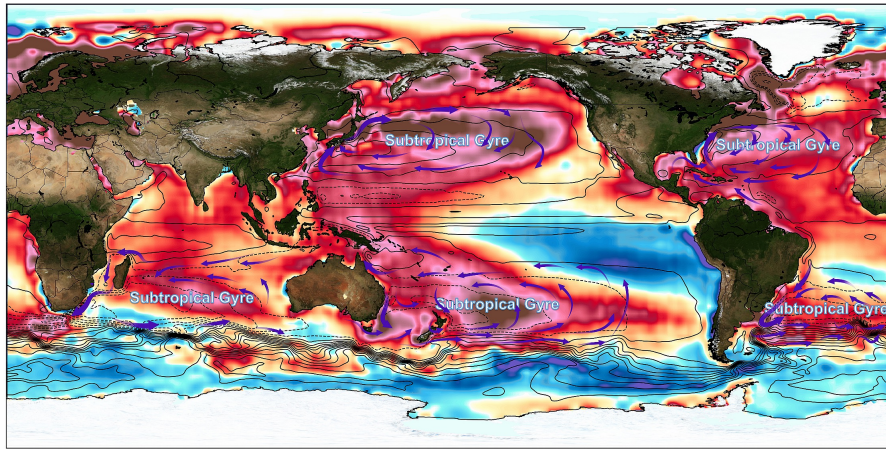
# Simulated averaged age of the upper 300-m ocean water column





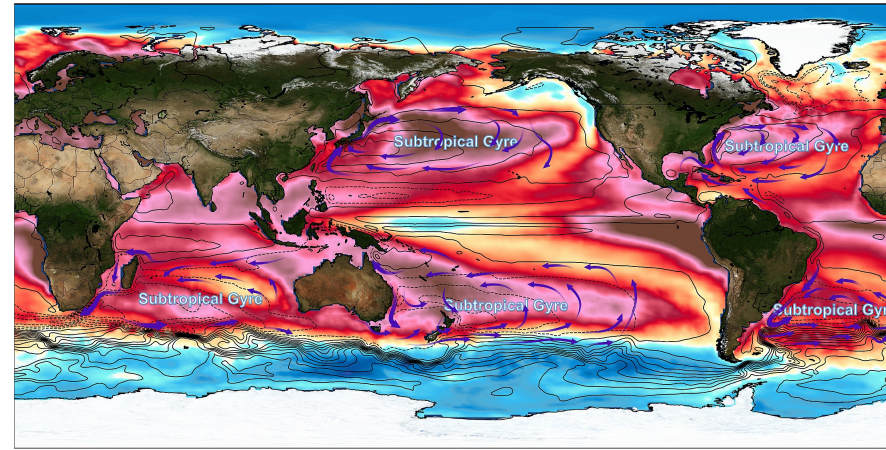
# Satellite-observed strong subtropical ocean warming as an early signature of global warming

Satellite Observed Ocean Warming



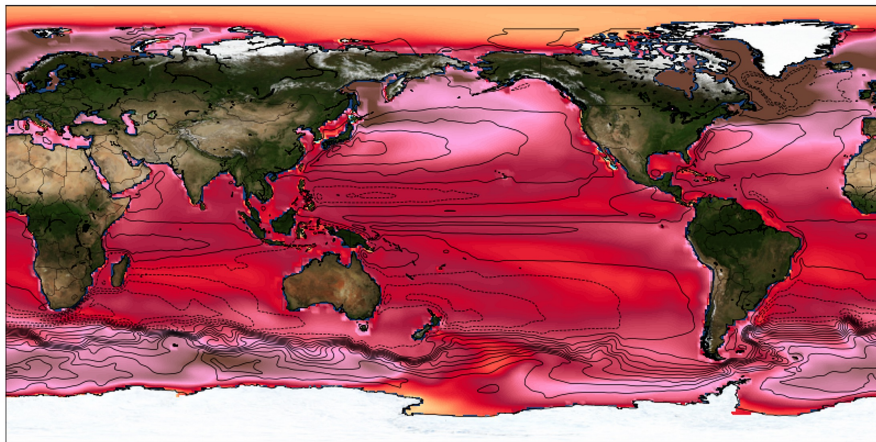
-0.8 -0.6 -0.4 -0.2 0.0 0.2 0.4 0.6 0.8 1.0 1.2 1.4  
Sea Surface Temperature Trend (°C)

Simulated Early Stage of Ocean Warming



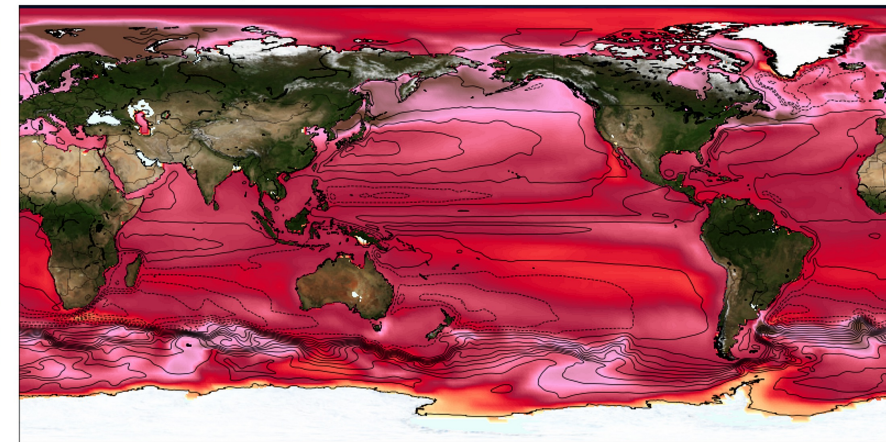
-0.165 -0.055 0.055 0.165 0.275 0.385 0.495 0.605 0.715 0.825 0.935 1.045  
Sea Surface Temperature Anomaly (°C)

Mid-Pliocene Ocean Warming



-2.8 -2.1 -1.4 -0.7 0.0 0.7 1.4 2.1 2.8 3.5 4.2 4.9  
Sea Surface Temperature Anomaly (°C)

Simulated Equilibrium Ocean Warming (4xCO<sub>2</sub>)



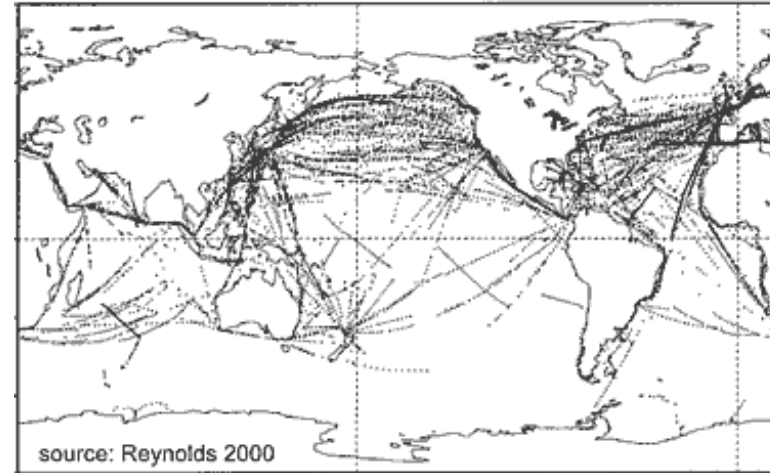
-5.6 -4.2 -2.8 -1.4 0.0 1.4 2.8 4.2 5.6 7.0 8.4 9.8  
Sea Surface Temperature Anomaly (°C)

Early stage

equilibrated stage

# The “Climate dilemma“

- The records of direct temperature measurements are short and already fall in the phase of strong human influence.
- Instrumental data are sparse

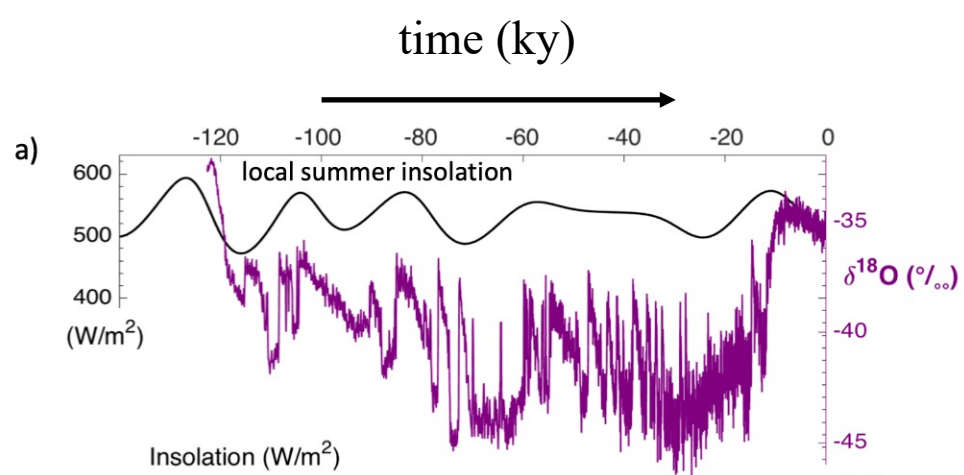


- For the time before instrumental records, one has to rely on information from **long-term data and modeling.**

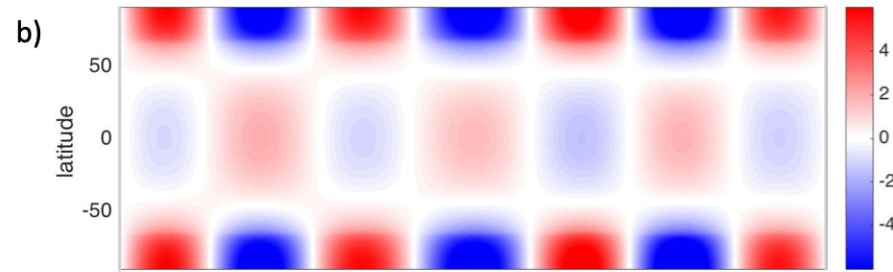




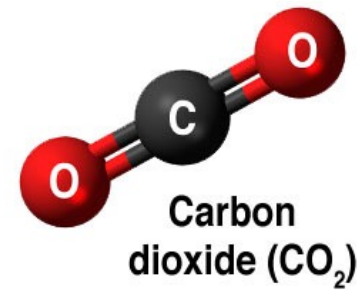
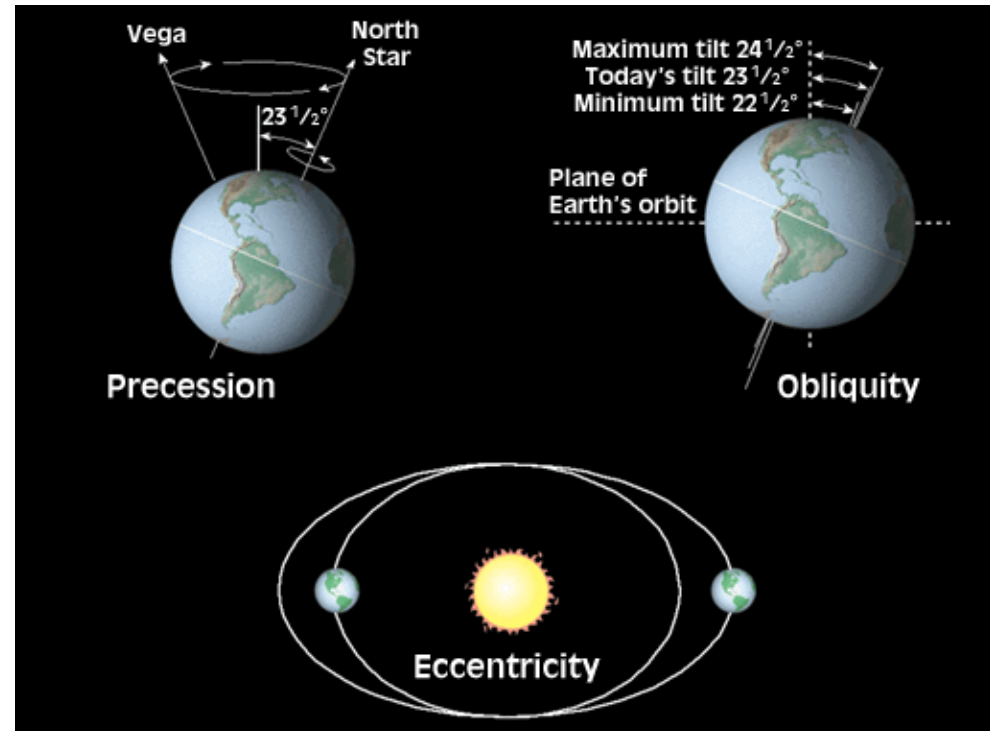
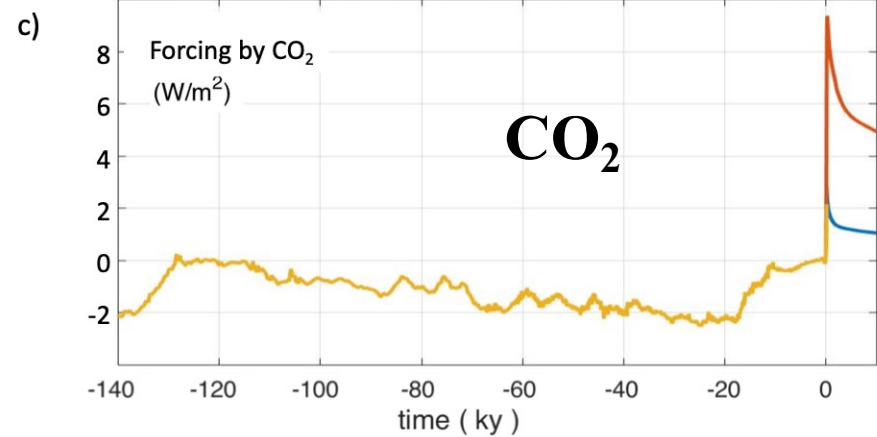
Abrupt climate change



Short wave radiation



Long wave radiation



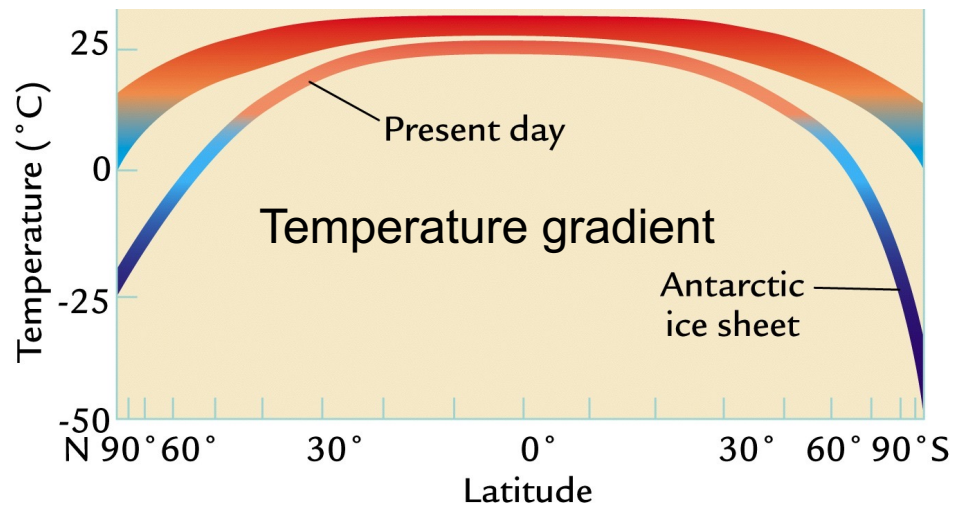
# Effect of **obliquity** on the **position Tropic of Cancer**

Highway in Mexico

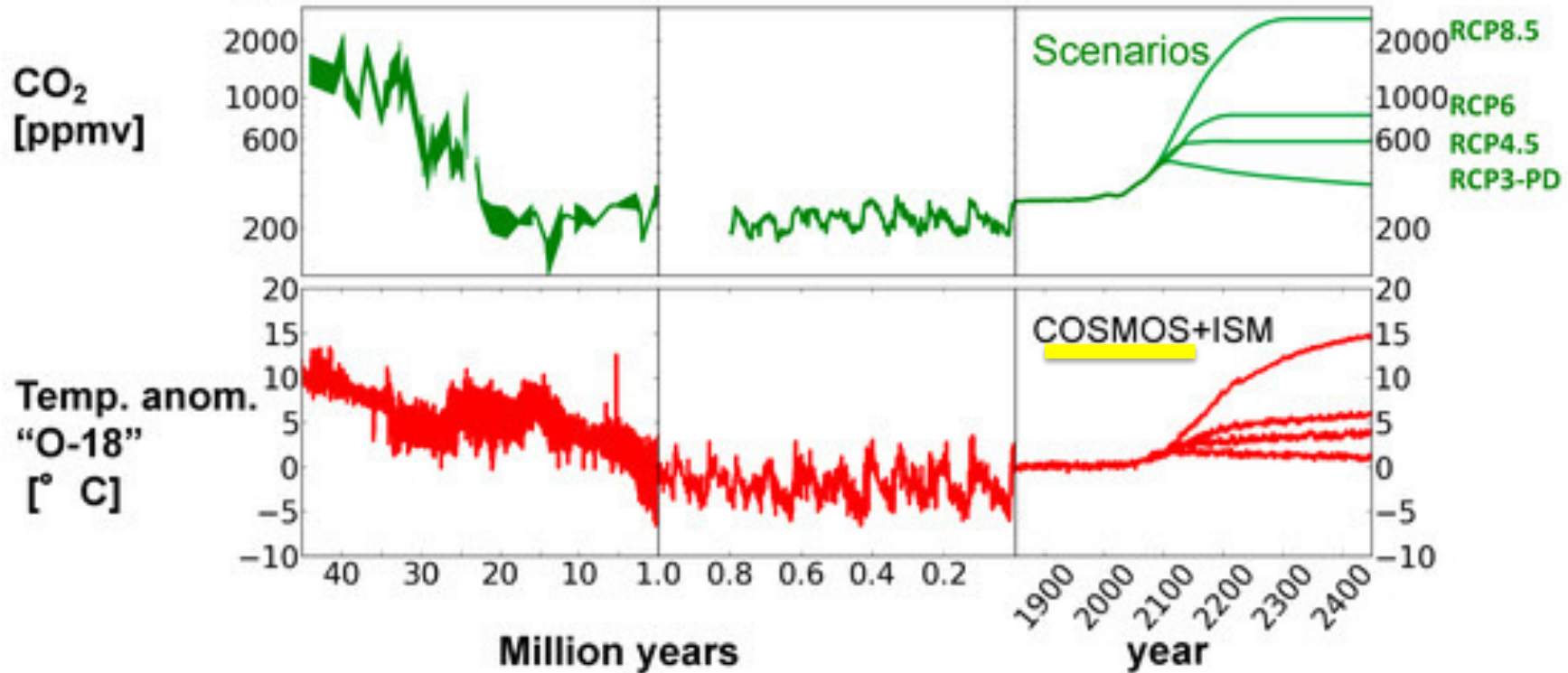
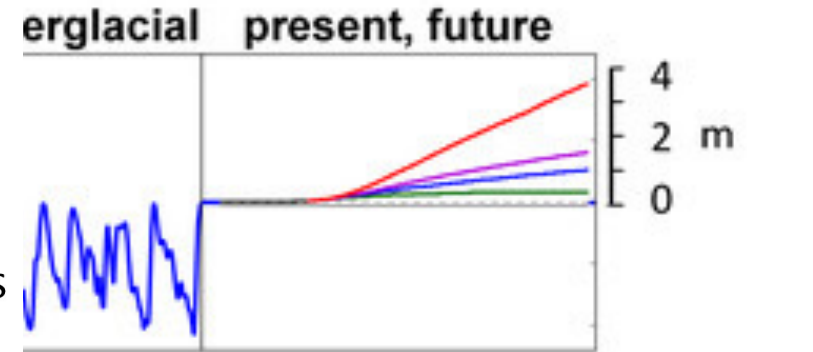


How many meters per year?

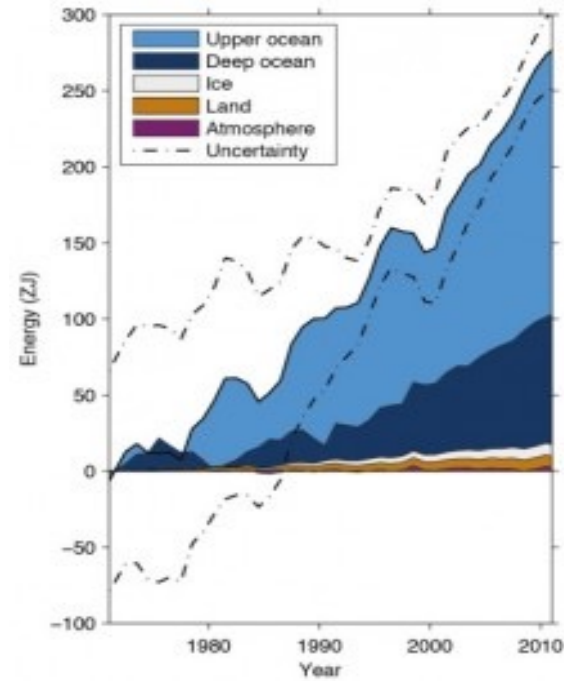
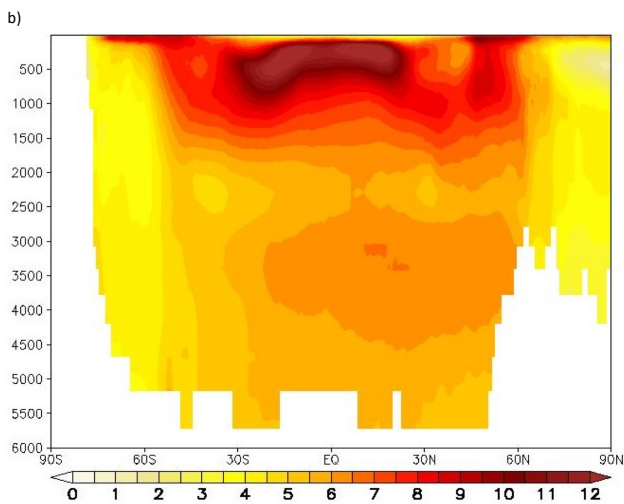
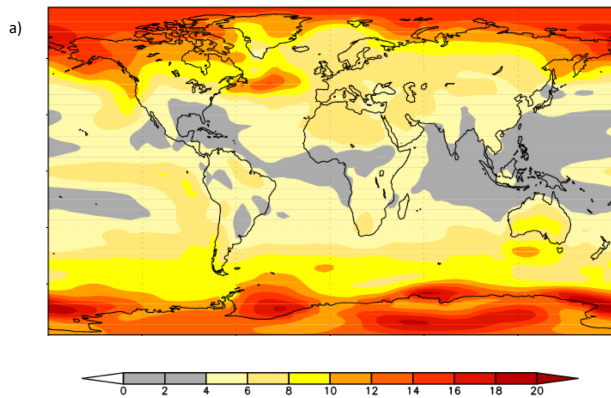
Earth's obliquity oscillates between  $22.1^\circ$  and  $24.5^\circ$  on a 41,000-year cycle.  
The Earth radius  $a=6371$  km



Disturbed climate



PACES program = our last 10 years



“old”  
IPCC  
report

## Effective heat capacity/heat uptake

$$C_p^o \partial_t T = \partial_z (k^o \partial_z T)$$

Increased  $k$  leads to high latitude warming  
& pronounced warming at the thermocline.



# Effects of CO<sub>2</sub> and Ocean Mixing on Miocene and Pliocene Temperature Gradients

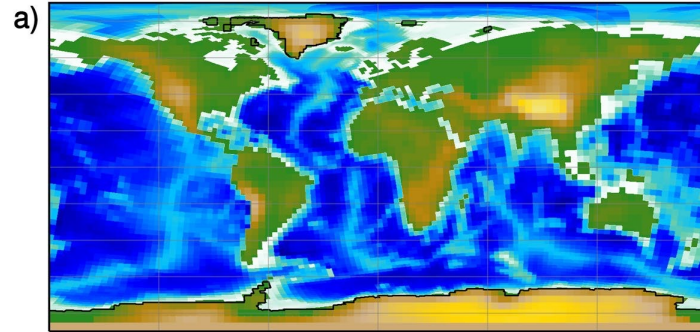
**Special Section:**

The Miocene: The Future of the Past

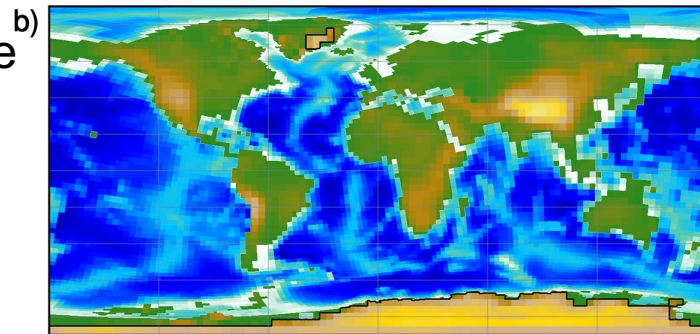
Gerrit Lohmann<sup>1,2</sup> , Gregor Knorr<sup>1</sup>, Akil Hossain<sup>1</sup> , and Christian Stepanek<sup>1</sup> 

<sup>1</sup>Alfred Wegener Institute (AWI) Helmholtz Center for Polar and Marine Research, Bremerhaven, Germany, <sup>2</sup>Department of Environmental Physics and MARUM, University of Bremen, Bremen, Germany

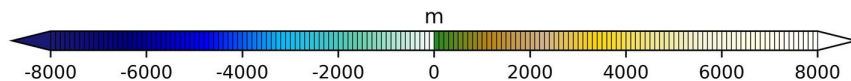
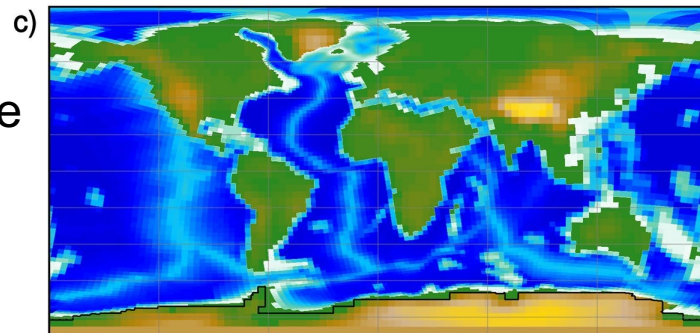
PI



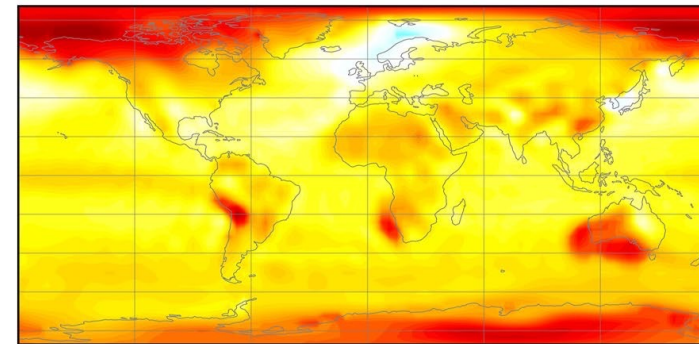
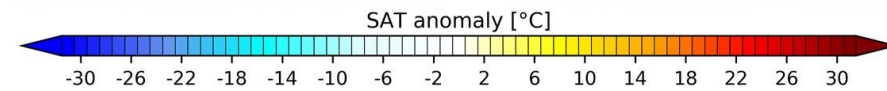
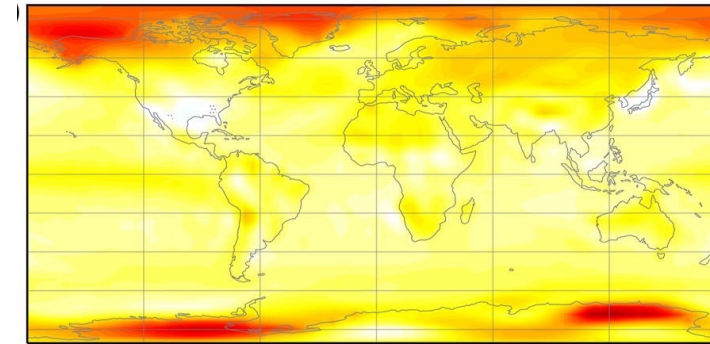
Pliocene



Miocene

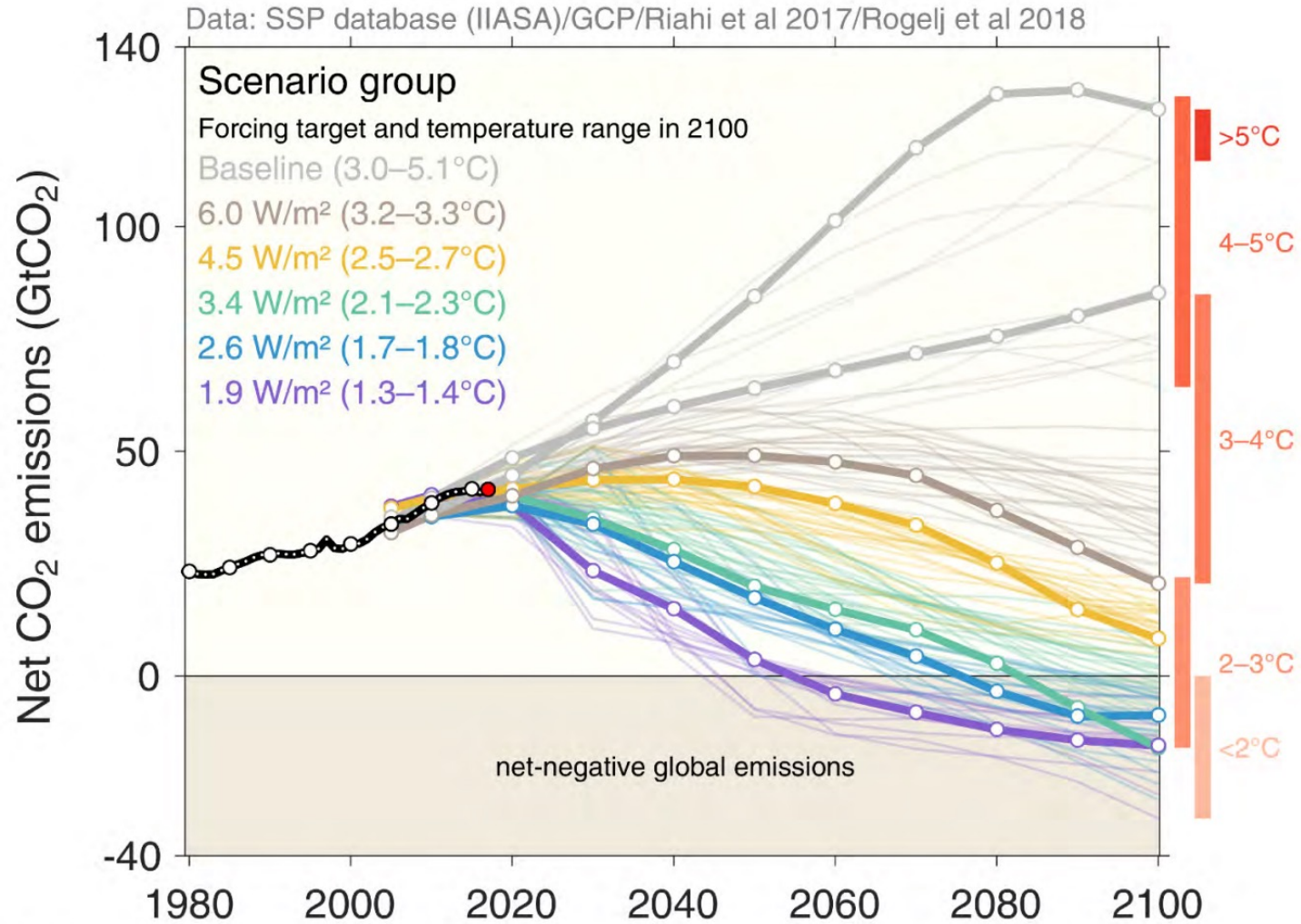


## Temperature anomalies



# Scenarios

## Global temperatures and CO<sub>2</sub> of the future?





# Modelling of the Earth System

## 1) 13.4. INTRODUCTION:

Earth system model components, definitions, processes. (MV)

## 2) 27.4. NUMERICAL APPROXIMATIONS I:

Finite differences: Ordinary differential equations (Runge-Kutta etc) (ST)

Gerrit Lohmann / Silke Thoms / Thomas Jung / Mihalis Vrekoussis

## 3) 4.5. NUMERICAL APPROXIMATIONS II:

Finite differences: Partial differential equations (Arakawa Grids etc) (ST)

## 4) 11.5. EXAMPLES: waves, diffusion, boundaries (ST)

## 5) 25.5. NUMERICAL APPROXIMATIONS III:

Finite Volume and Finite Elements and spectral methods (atmosphere and ocean) (TJ)

## 6) 14.7. HIGH-PERFORMANCE COMPUTING (scalability, Moore's law) (TJ)

## 7) 8.6. ATMOSPHERIC CHEMISTRY I:

Chemistry Transport Models (chemical processes including types of models, box models, grids, coordinates) (MV)

## 8) 15.6. ATMOSPHERIC CHEMISTRY II:

Inverse methods (MV)

## 9) 22.6. Earth system models including TRACERS and DYNAMICAL VEGETATION (GL)

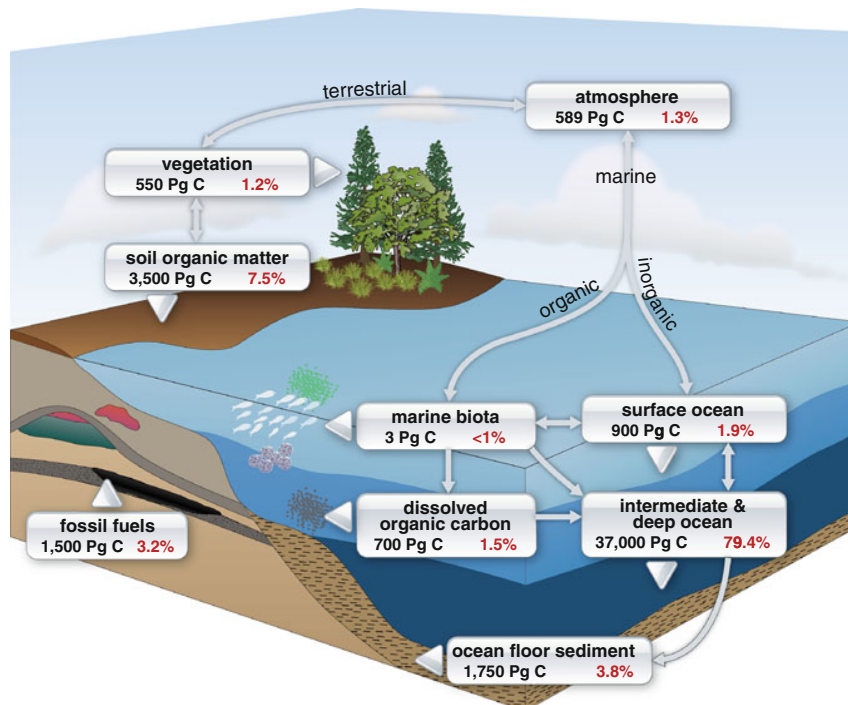
## 10) 29.6. DATA ASSIMILATION (Kalman filters etc) (TJ)

## 11) 6.7. RANDOM SYSTEMS (Stochastic differential equations, Lattice Gases) (GL)

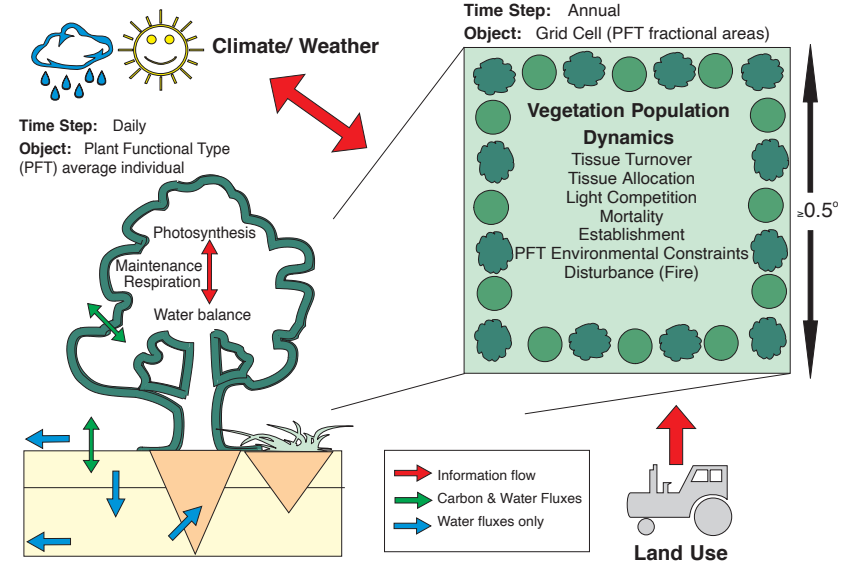
## 12) 13.7. CRYOSPHERE (Sea ice, ice sheets, and permafrost) (GL)



# Earth system models including tracers and dynamical vegetation



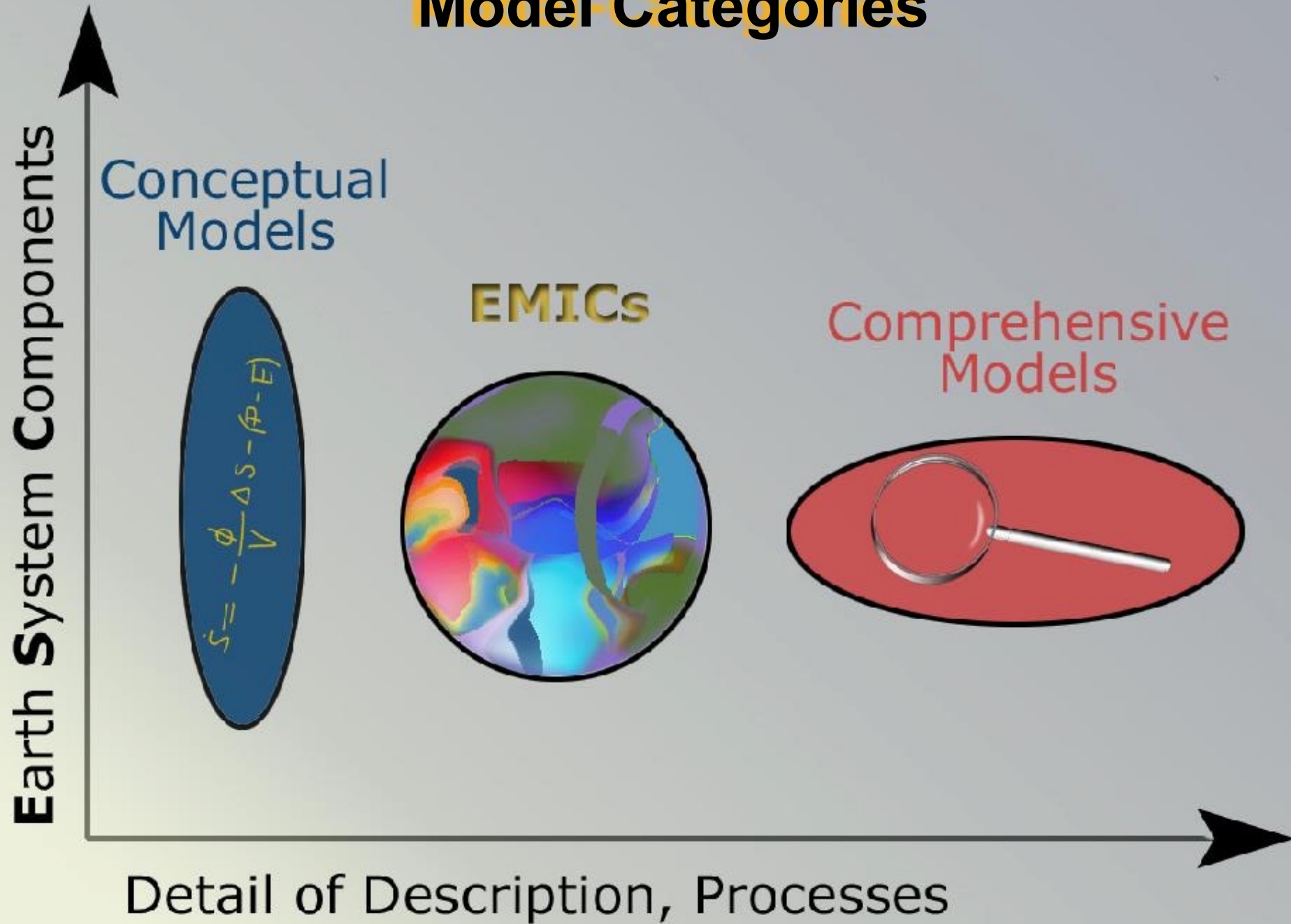
The Lund-Potsdam-Jena Dynamic Global Vegetation Model (DGVM)



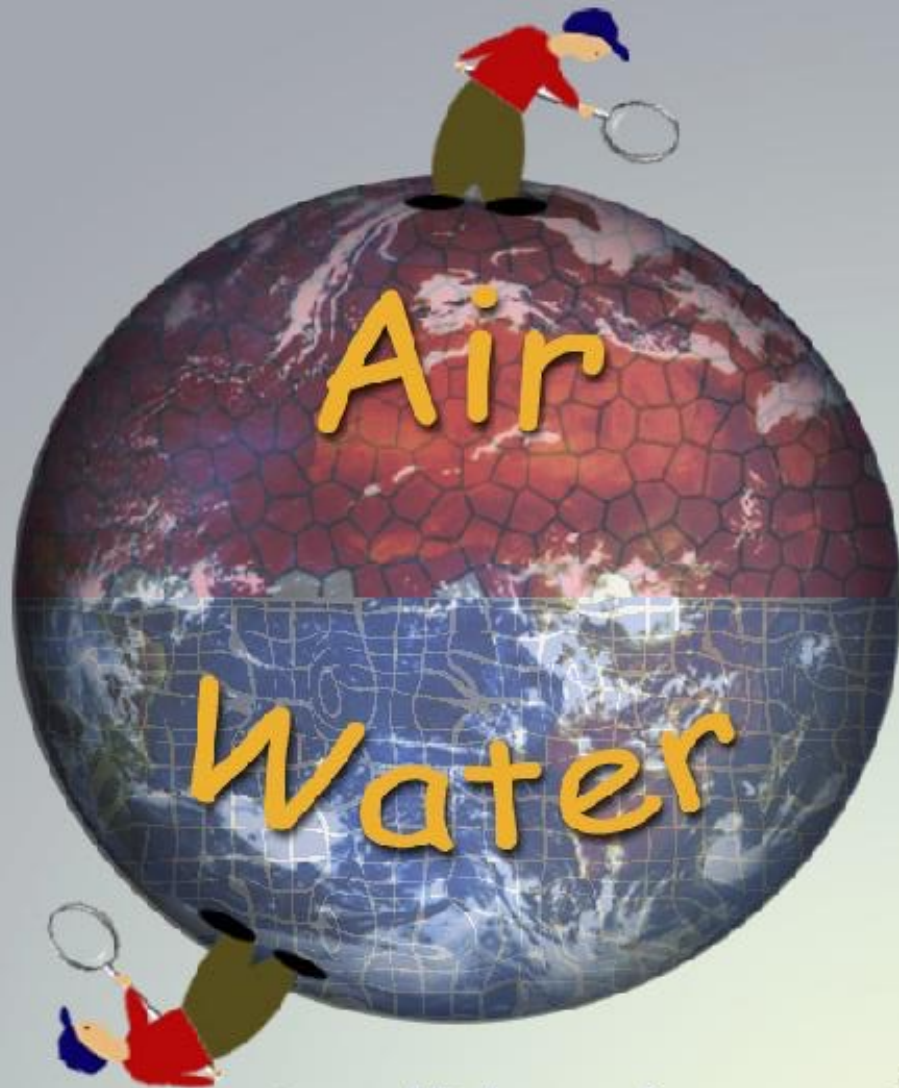
# Today's lecture

- Earth System Models
- Vegetation & Ecosystem models
- Practicals: Daisy World, vegetation dynamics  
<https://paleodyn.uni-bremen.de/study/MES/daisy/>
- Tracers in the Sea (Carbon, Radiocarbon)

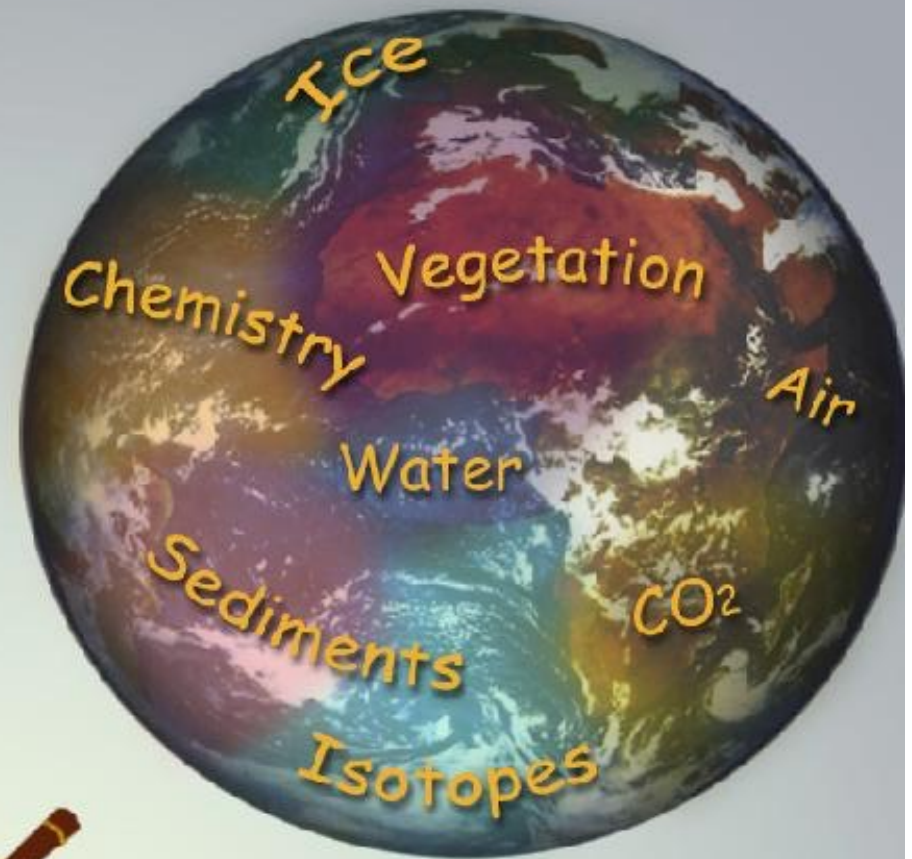
# Model Categories



# Different Point of View



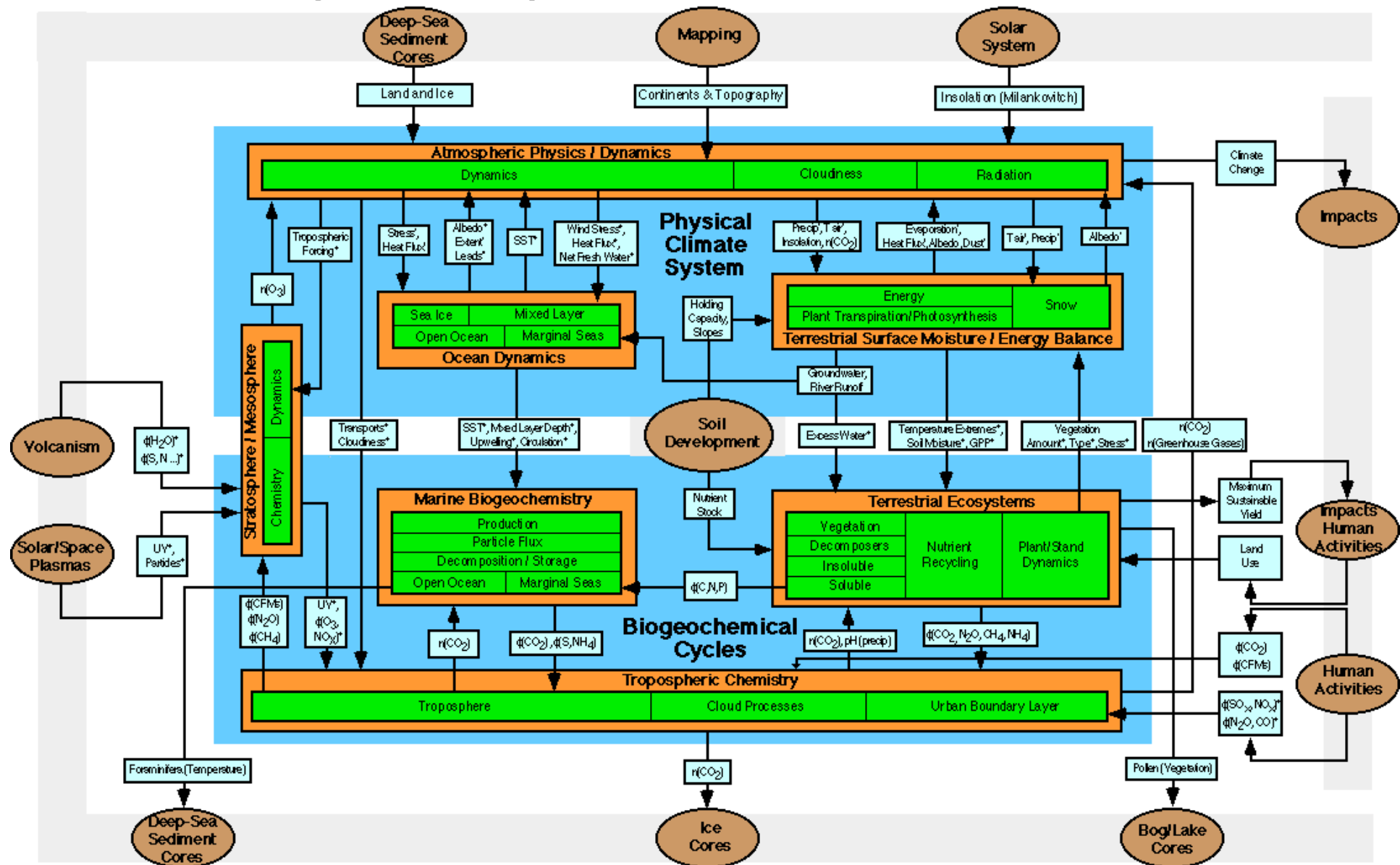
traditional  
GCMs



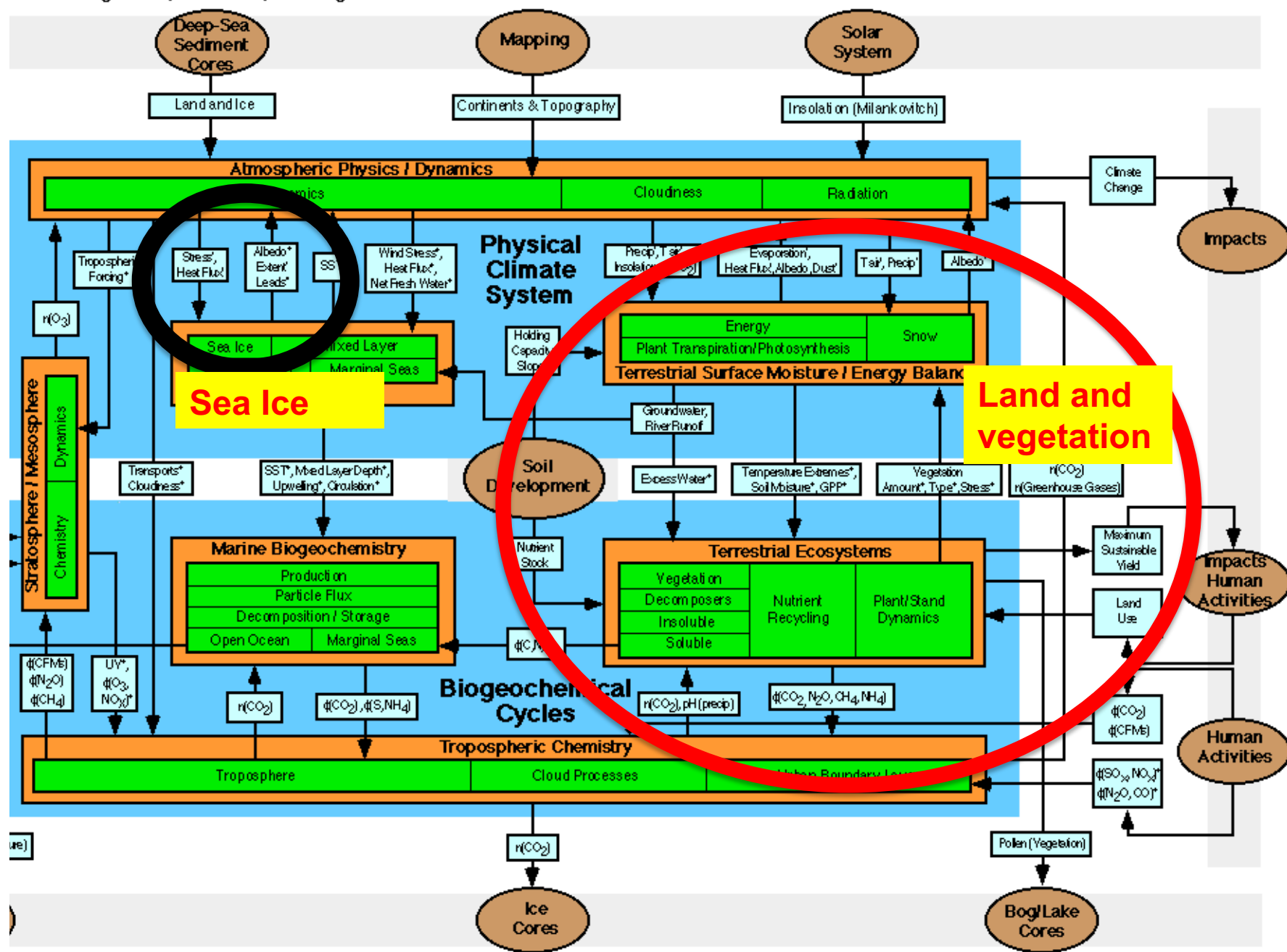
Earth System  
Models



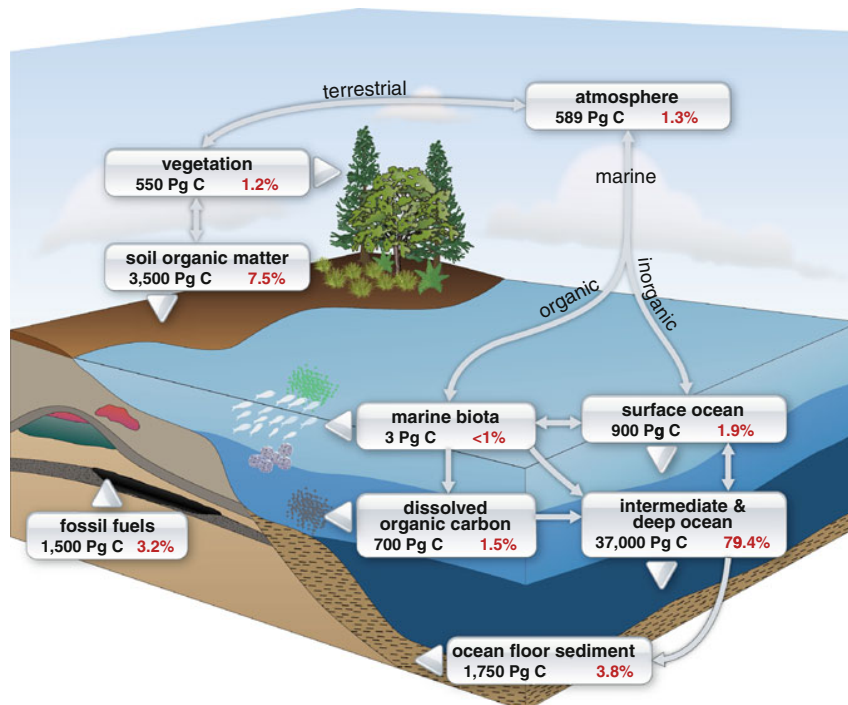
CONCEPTUAL MODEL of Earth System process operating on timescales of decades to centuries



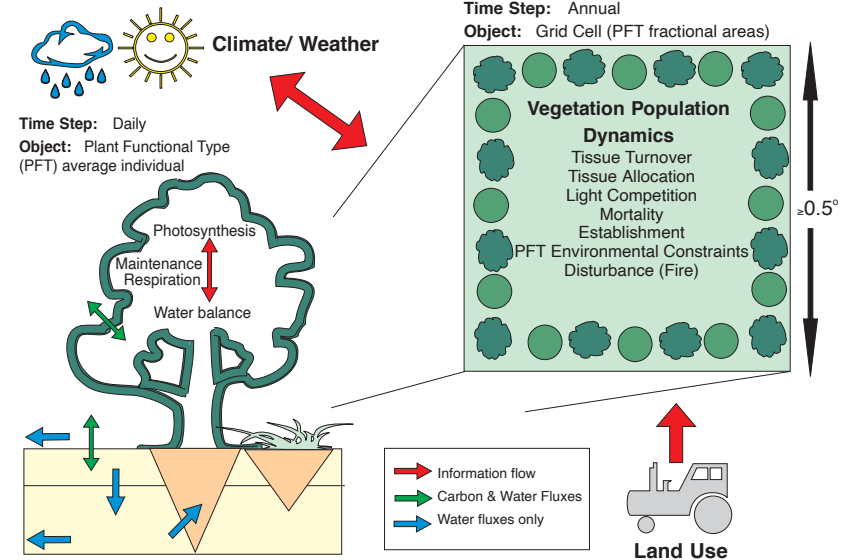
\* = on timescale of hours to days    \* = on timescale of months to seasons     $\phi$  = flux    n = concentration



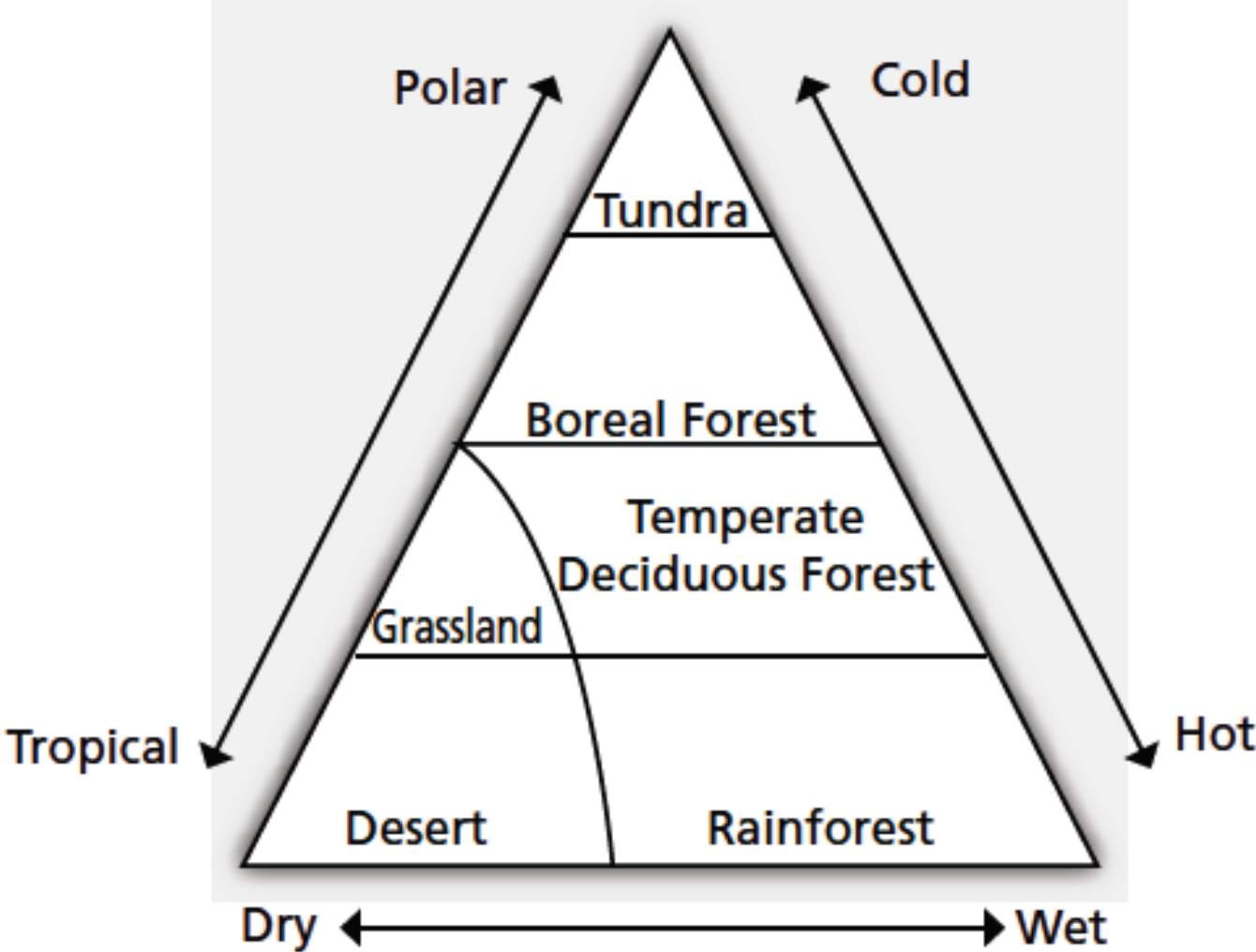
# Earth system models including tracers and dynamical vegetation



The Lund-Potsdam-Jena Dynamic Global Vegetation Model (DGVM)



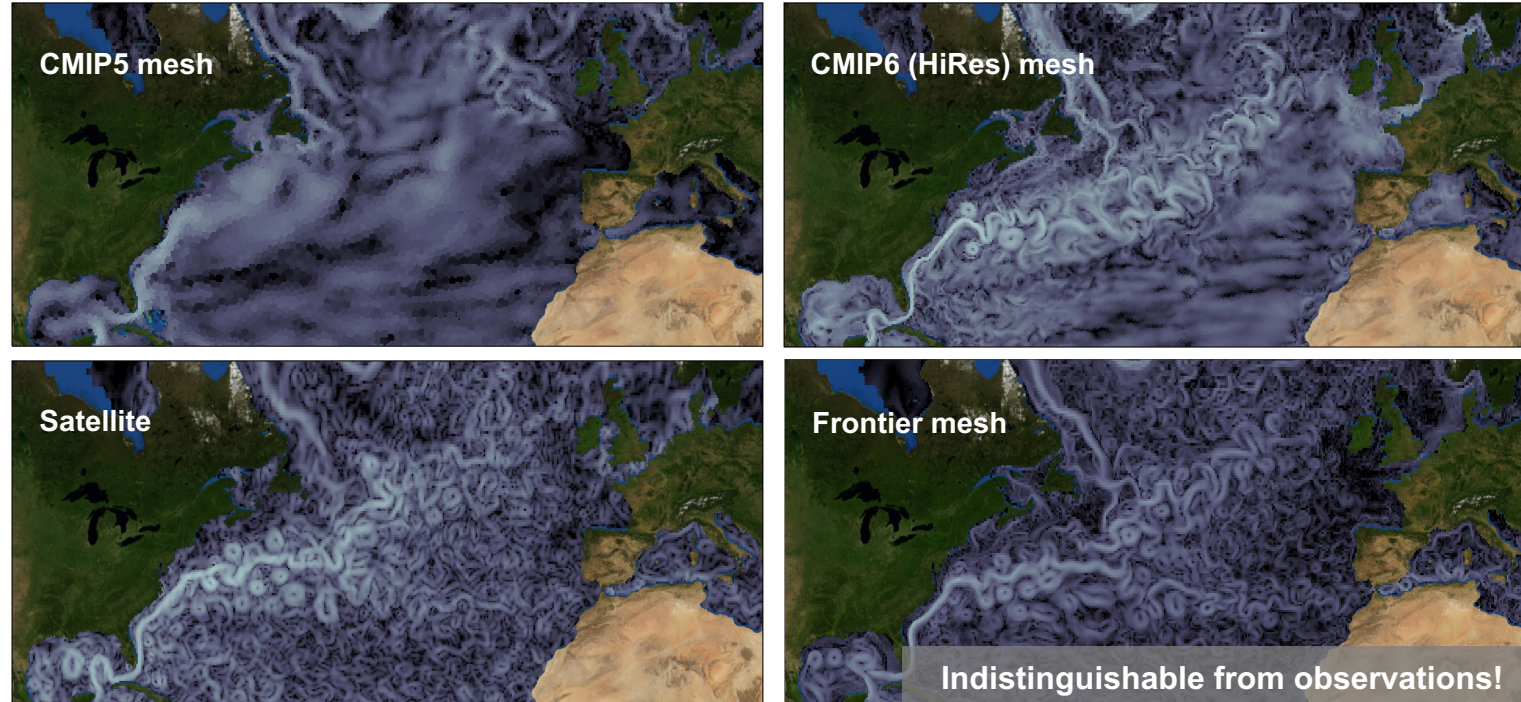
# Vegetation & Ecosystem



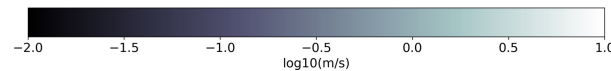


# How realistic are climate models ?

## Ocean velocity

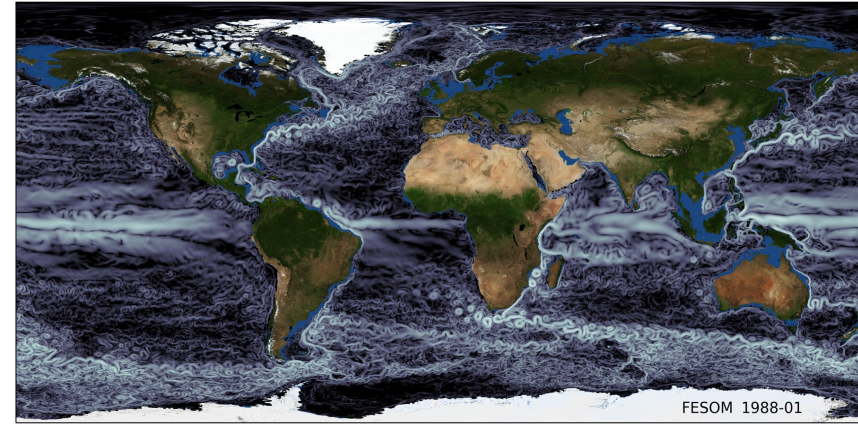
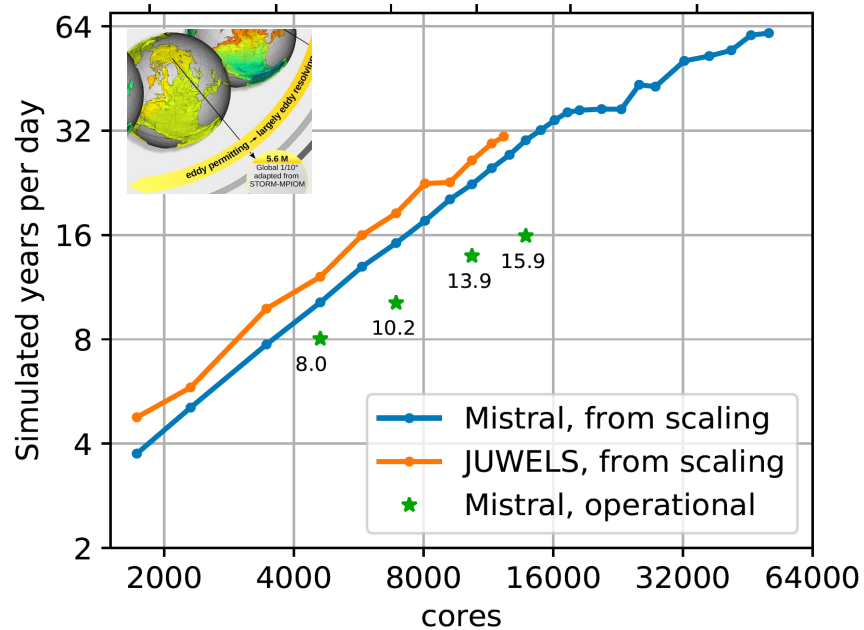


Displayed on a common 1/4° mesh



- Large uncertainties in regional changes
- Limitations for extreme events

# Scalability



Koldunov et al (2019)

Limited by available HPC capabilities (today)

Limited by our ability to use future HPC systems (tomorrow)

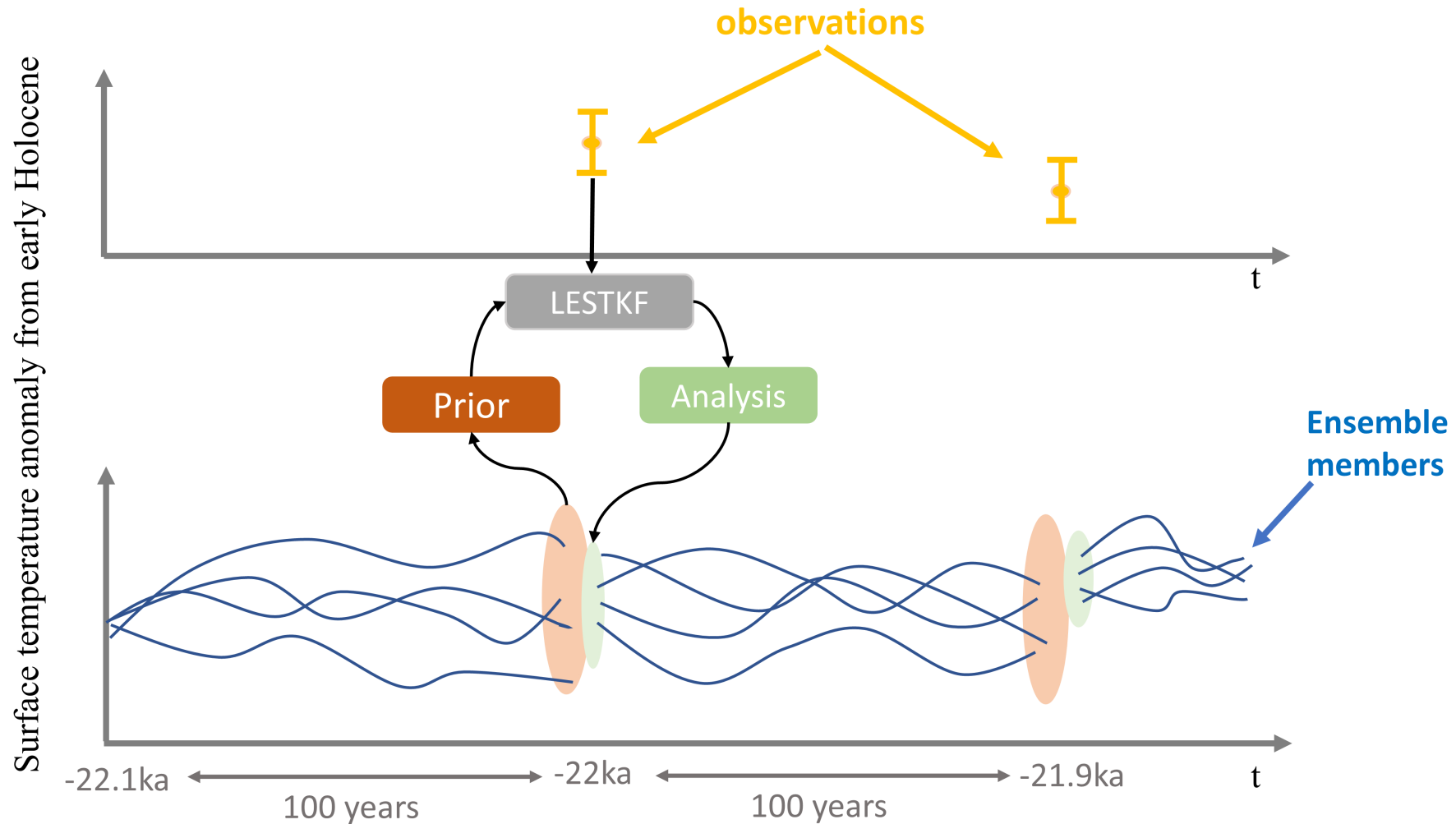
## Parameterizations

Some critical small-scale processes are *not* represented by the laws of physics, but by physically motivated rules of thumb (parameterizations)

- Large uncertainties in regional (global) climate change projections
- Limitations in predicting extreme events, major feedbacks (ice)

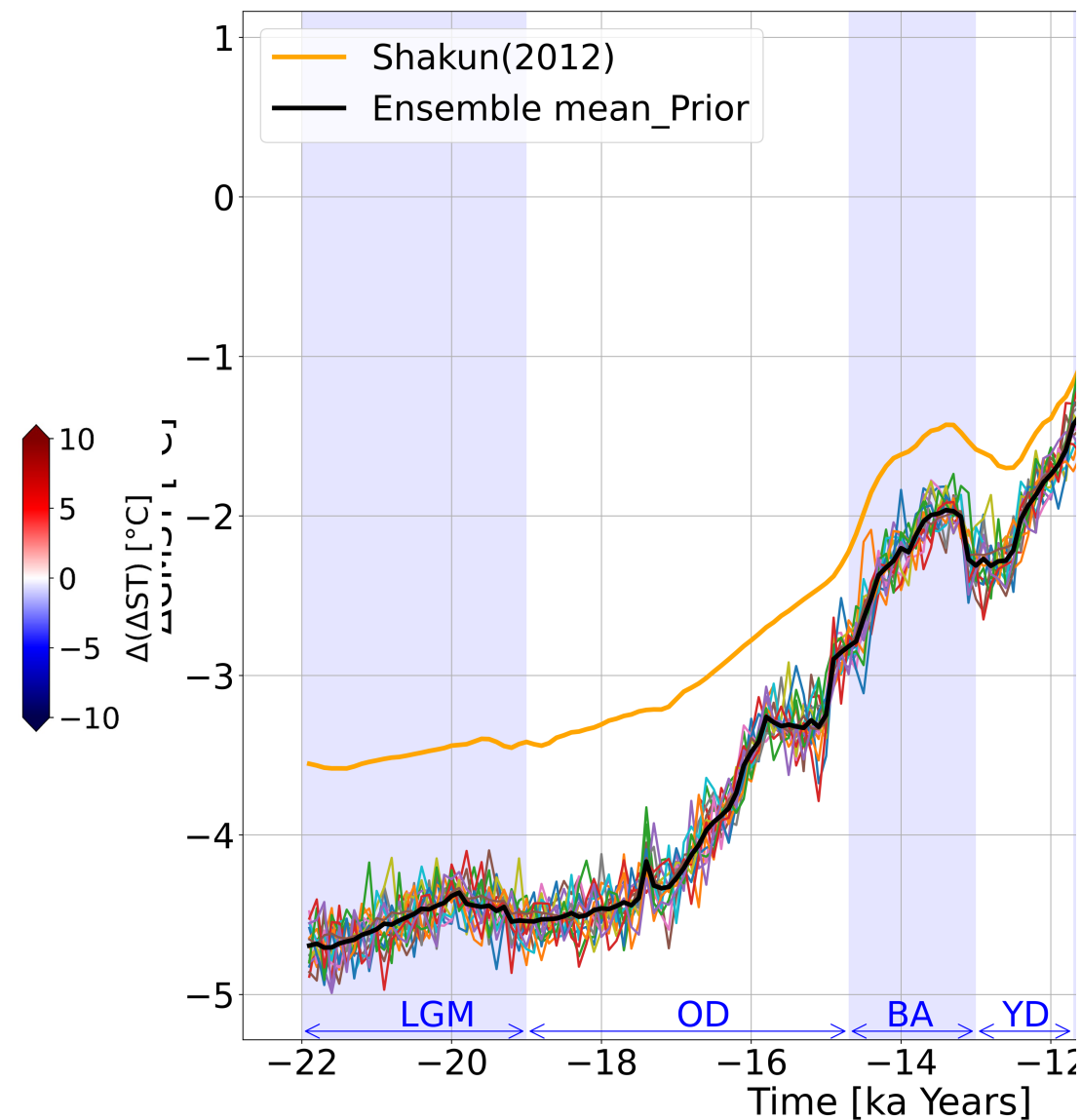
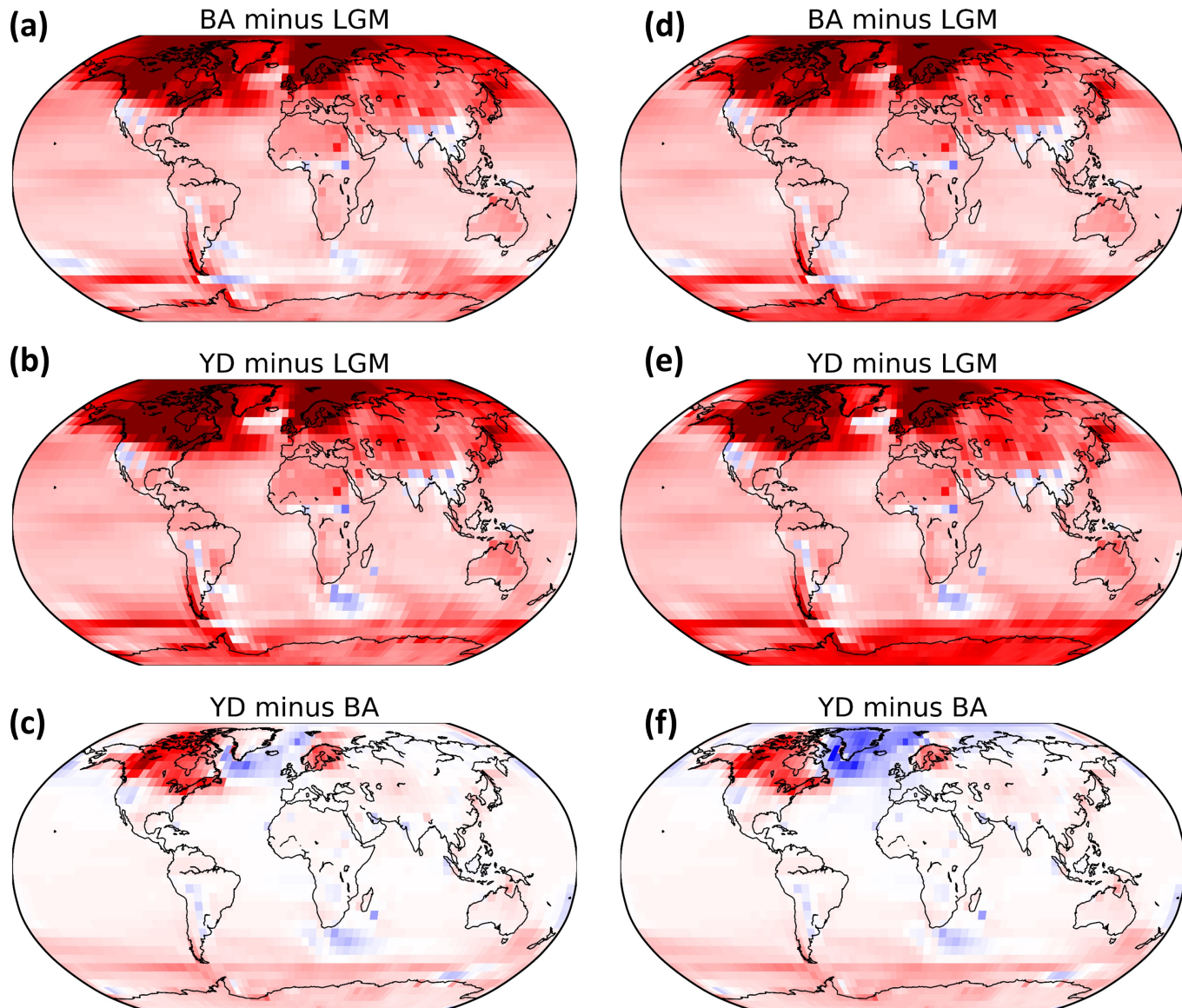
# Data assimilation (Ensemble Kalman Filter)

To make use out of data and models





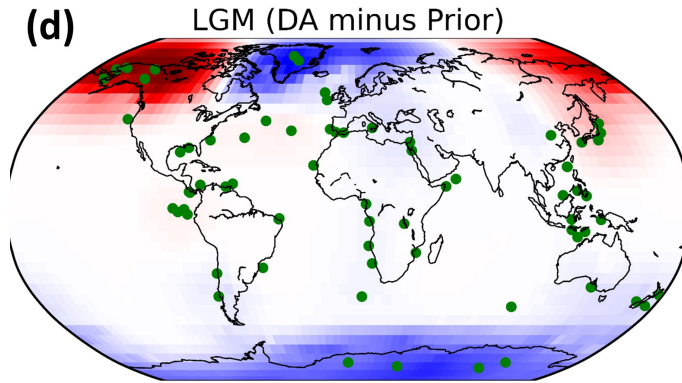
# Prior vs DA



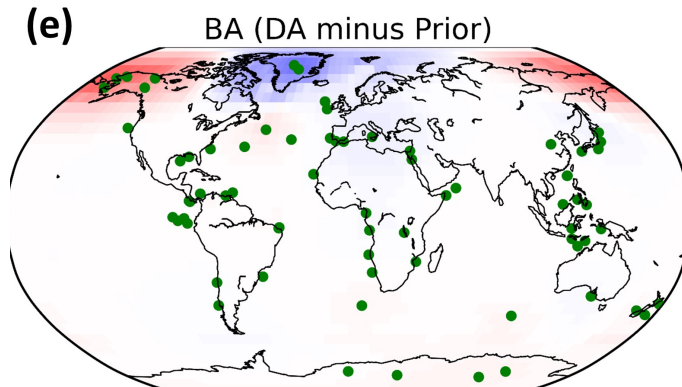
Exp\_PaleoMist

# Effect of Data assimilation

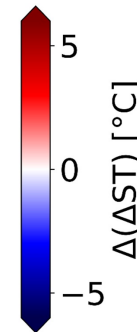
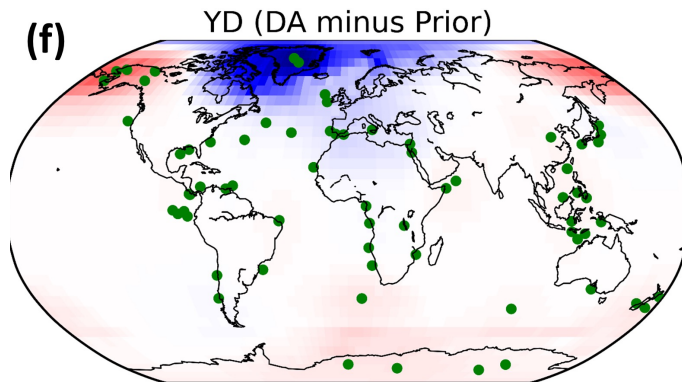
LGM



Boelling/  
Alleroed



Younger  
Dryas

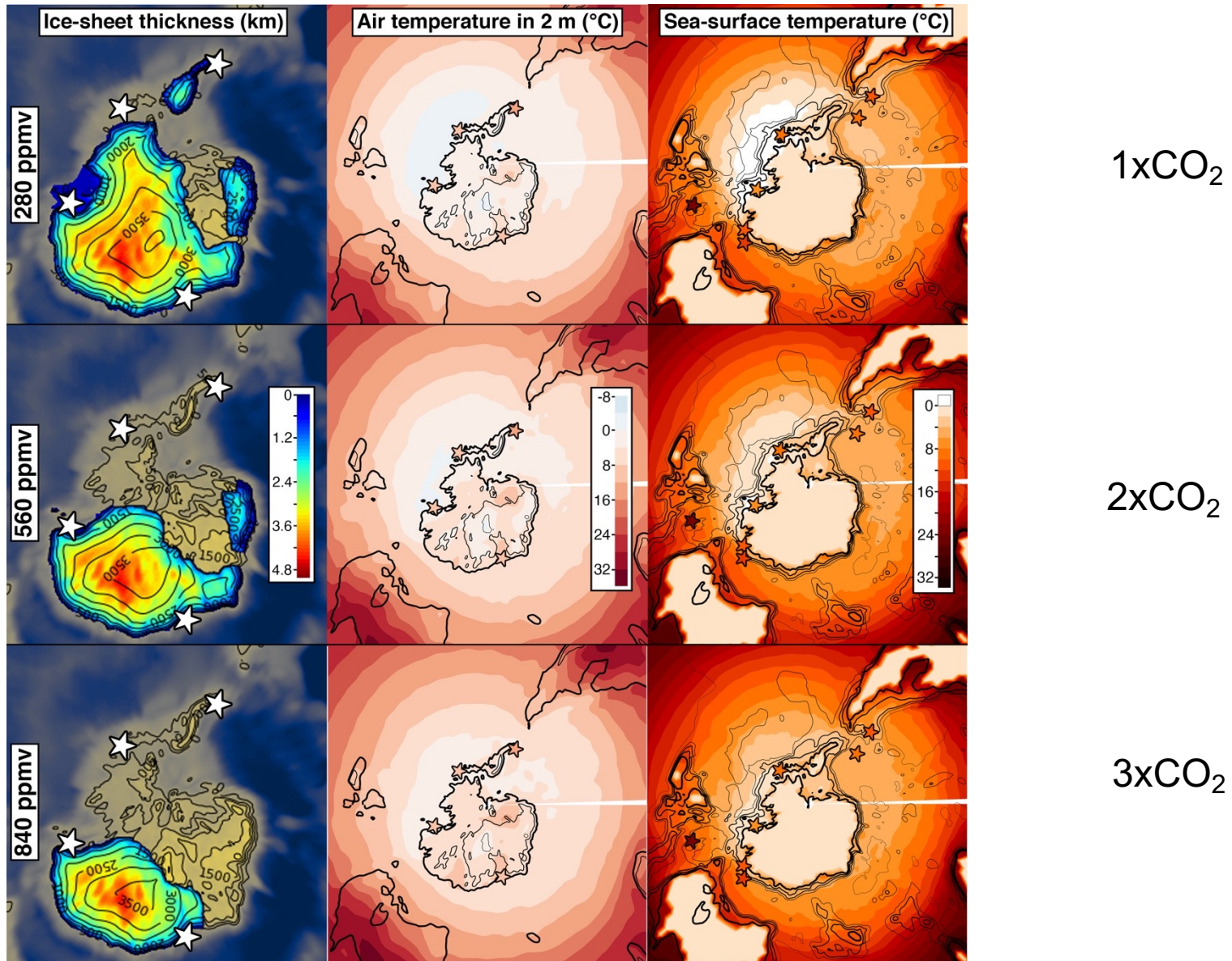


## What can learned:

- Data-model (mis-)fit
- Where do we need more data?
- Problem in the model



# Absent West Antarctic Ice Sheet during the early Oligocene glacial maximum



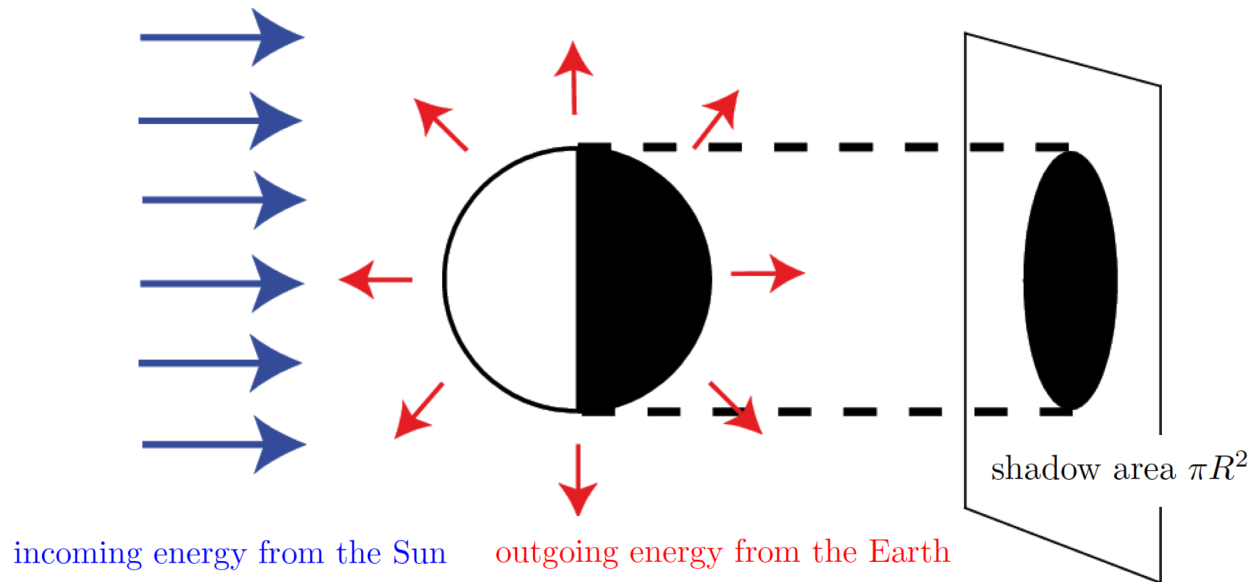




# Energy balance model: Concepts of climate

$$(1 - \alpha)S\pi R^2 = 4\pi R^2 \epsilon \sigma T^4$$

$$T = \sqrt[4]{\frac{(1 - \alpha)S}{4\epsilon\sigma}}$$



Heat capacity of the climate system

Fast rotation