

Climate II

(Winter 2020/2021)

7th lecture:

Dust and Vegetation

(glacial dust increase, iron fertilisation, radiative forcing, vegetation and aridity)

Gerrit Lohmann, Martin Werner

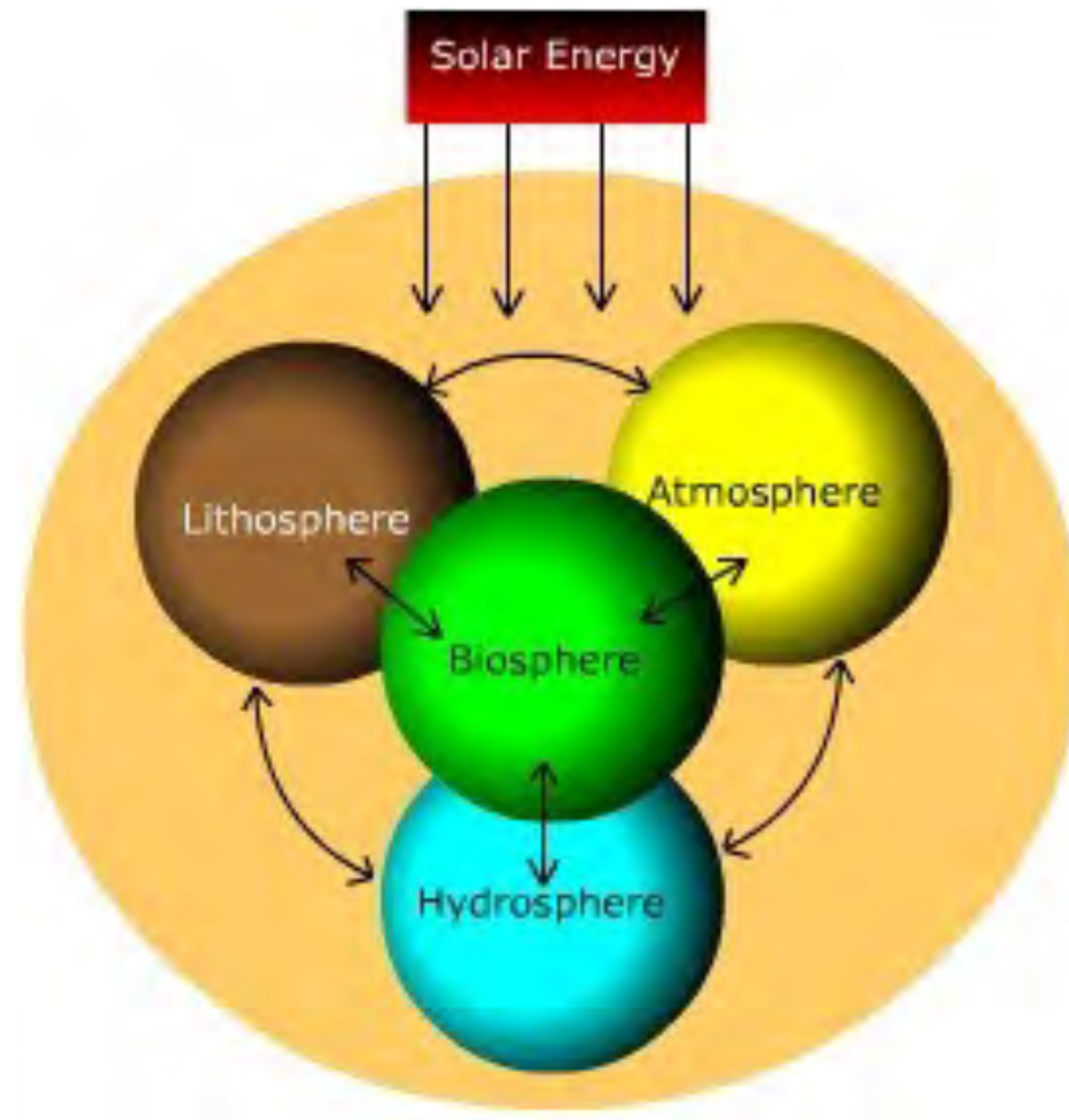
Tuesday, 10:00-11:45

(sometimes shorter, but with some exercises)

https://paleodyn.uni-bremen.de/study/climate2020_21.html

What are biogeochemical cycles?

- **Earth system has four parts**
 - *atmosphere*
 - *hydrosphere*
 - *lithosphere*
 - *biosphere*
- **Biogeochemical cycles**
 - *The chemical interactions (cycles) that exist between the atmosphere, hydrosphere, lithosphere, and biosphere*
- **Abiotic (physio-chemical) and biotic processes drive these cycles**

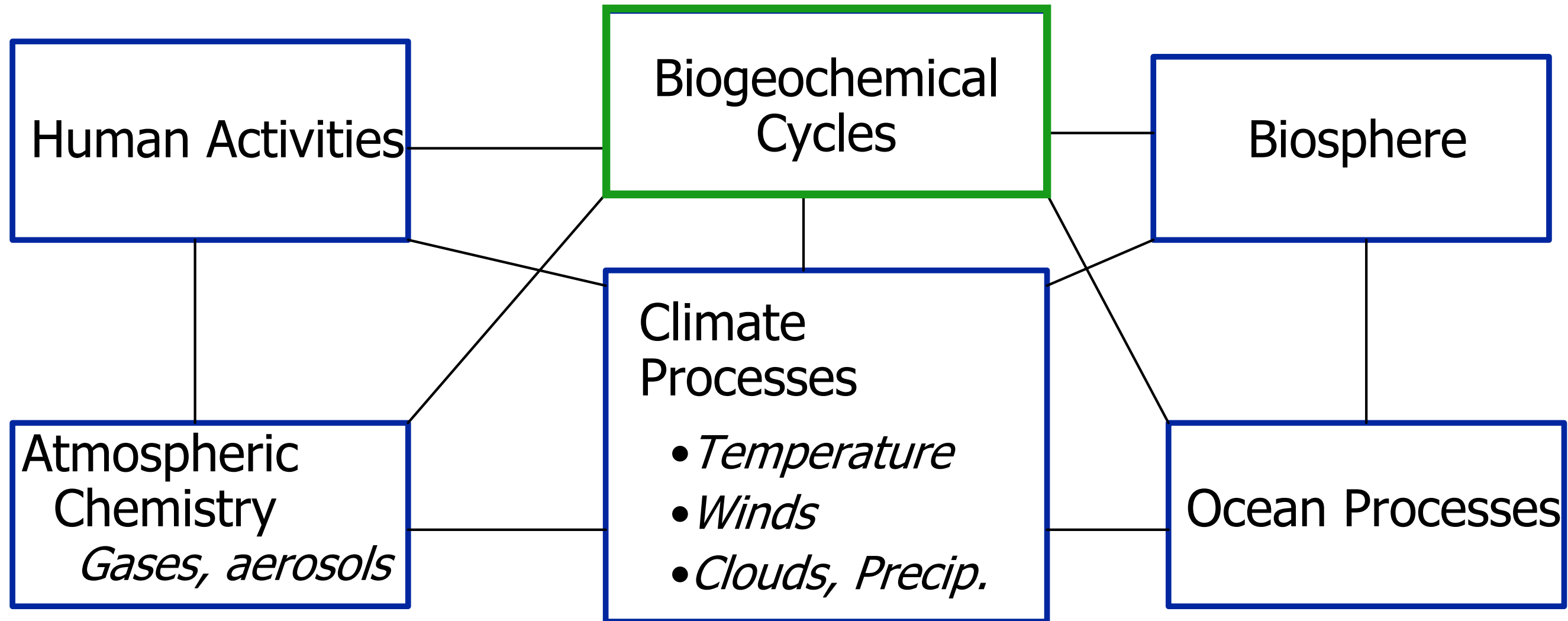


Biogeochemical cycles - key elements

Six nutrient elements make up 95% of the biomass mass on earth and form the biochemical foundation for life.

- **Carbon (CO_2 , CH_4 , CO)**
- **Nitrogen (N_2O , NO , NO_2 , NH_3)**
- **Sulfur (SO_2 , COS , H_2S , H_2SO_4)**
- **Phosphorous**
- **Hydrogen**
- **Oxygen**
- **Water**

Biogeochemical cycles in climate research

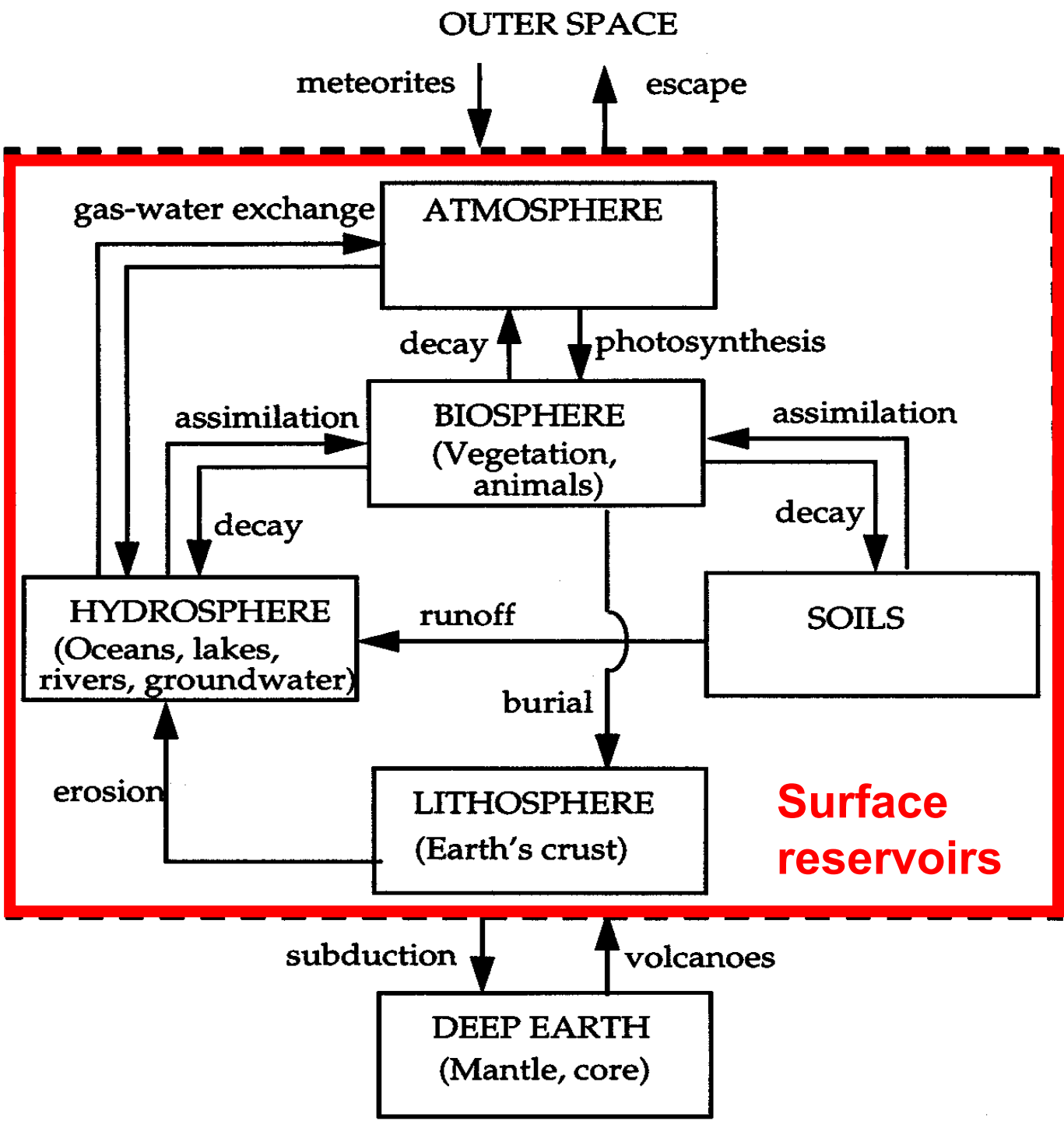
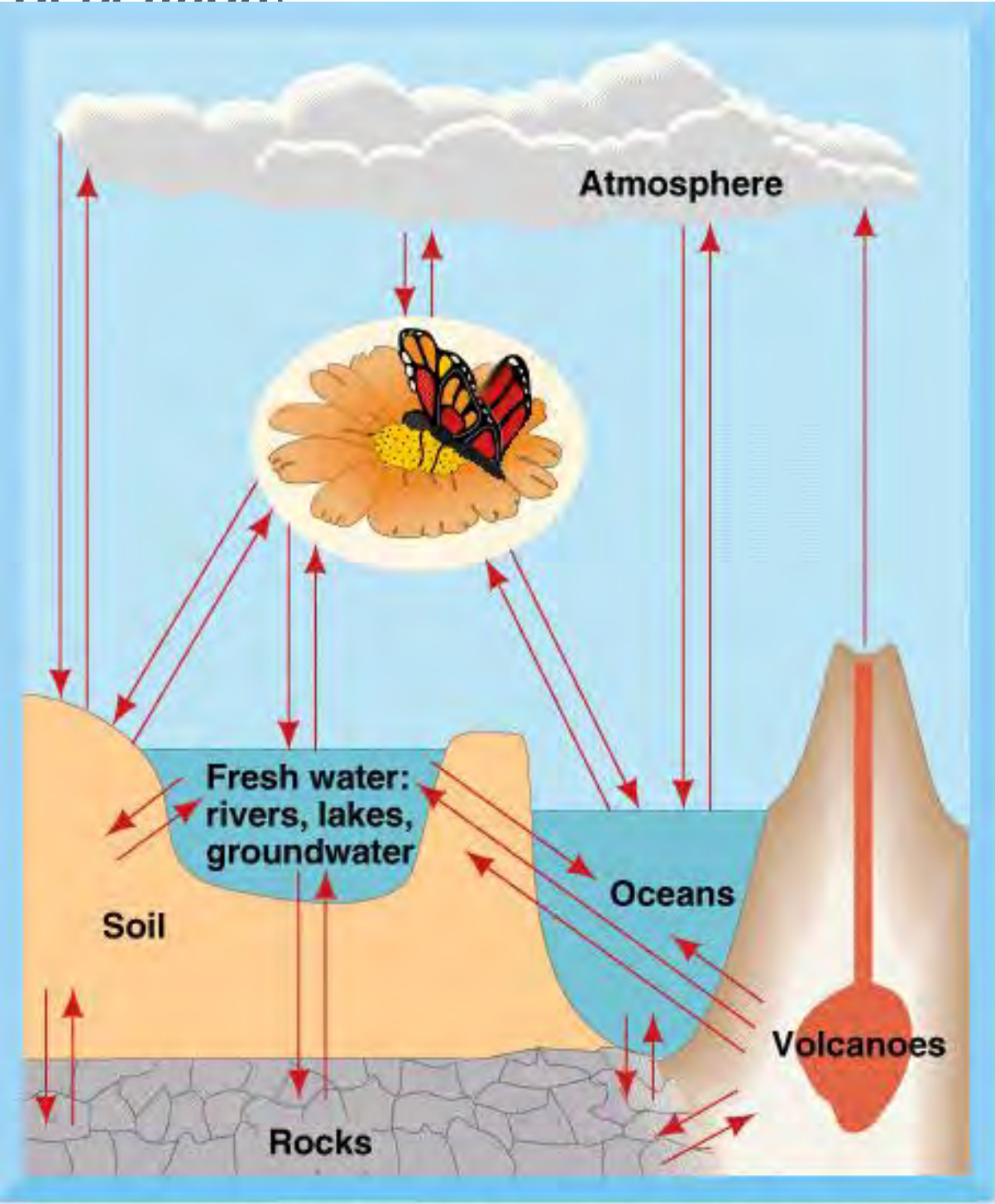


**Biogeochemical cycles are a key element
for understanding our past and present climate!**

Biogeochemical cycles: Common features of all key elements

- each element typically occurs in all four parts of the Earth System (e.g. water, carbon, nitrogen, etc.)
- each biogeochemical cycle can be described by
 - *pools*
 - *fluxes in and out of pools*
 - *chemical or biochemical transformations*

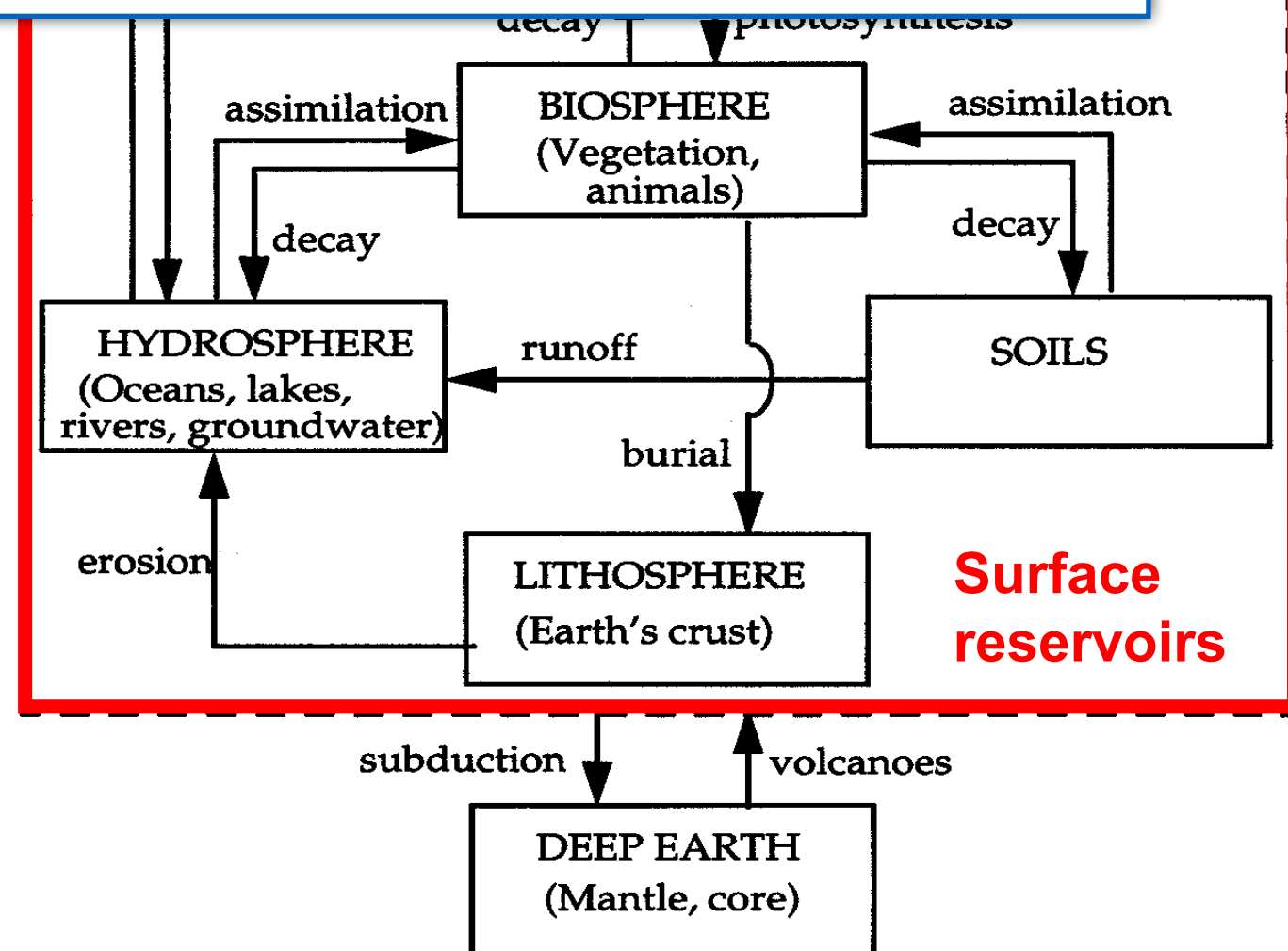
Repentition



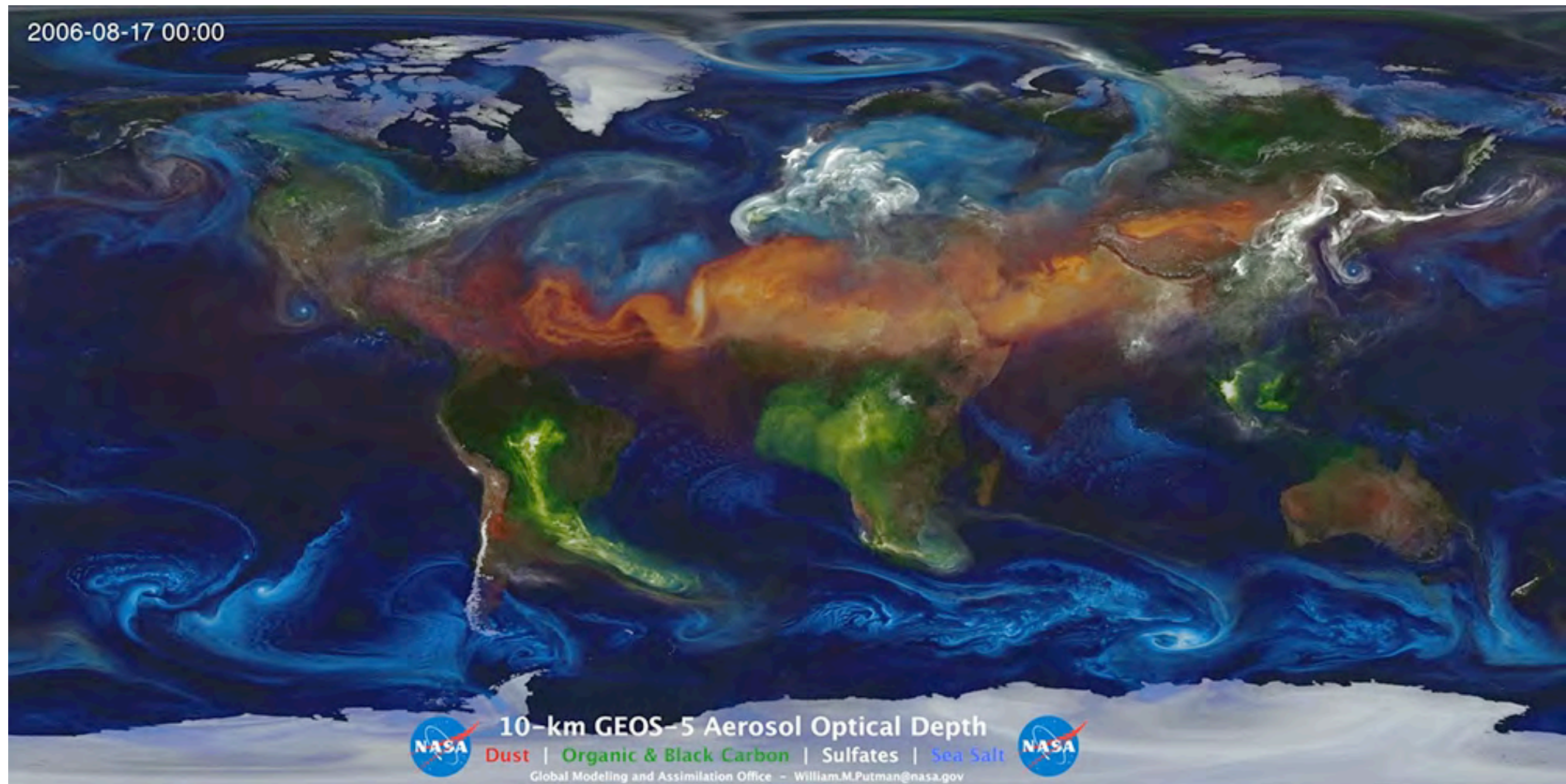
definition

turnover time τ is the size of a reservoir A , divided by the sum of all fluxes into (or out of) it $\sum F$, $\tau = A / \sum F$

- if there are only fluxes out of the reservoir, τ is how long it takes until the reservoir is empty
- if fluxes are balanced, τ is how long it takes until 1/e-th of the original molecules is still in the reservoir

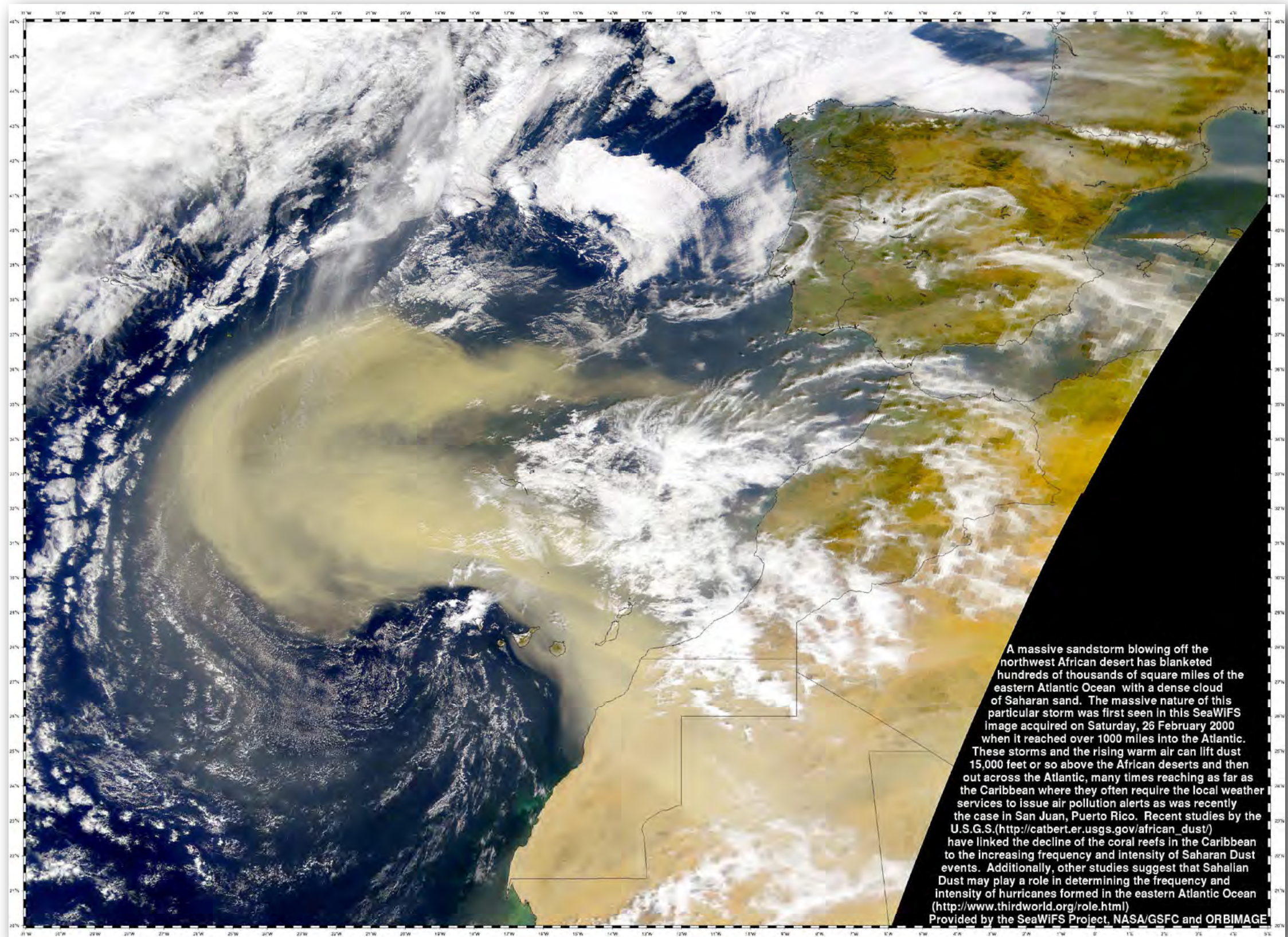


Transport of dust and biogeochemical aerosols



Animation 1. Aerosol optical thickness of black and organic carbon (green), dust (red-orange), sulfates (white), and sea salt (blue) from a 10 km resolution GEOS-5 "nature run" using the GOCART model. The animation shows the emission and transport of key tropospheric aerosols from August 17, 2006 to April 10, 2007.

Dust in the climate system



Dust depositions on glacial-interglacial time scales

Dust–climate couplings over the past 800,000 years from the EPICA Dome C ice core

Vol 452 | 3 April 2008 | doi:10.1038/nature06763

F. Lambert^{1,2}, B. Delmonte³, J. R. Petit⁴, M. Bigler^{1,5}, P. R. Kaufmann^{1,2}, M. A. Hutterli⁶, T. F. Stocker^{1,2}, U. Ruth⁷, J. P. Steffensen⁵ & V. Maggi³

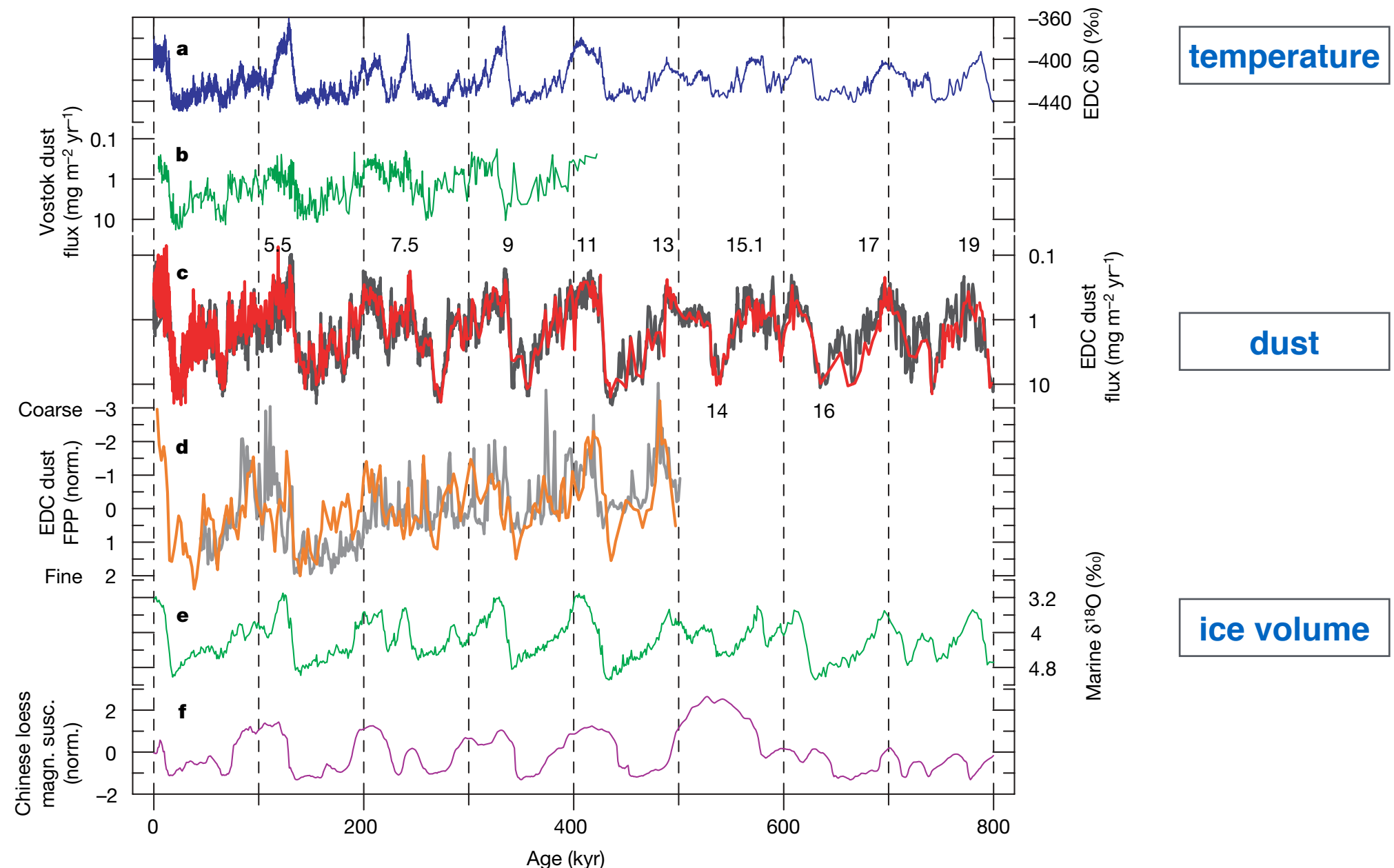


Figure 1 | EDC dust data in comparison with other climatic indicators. **a**, Stable isotope (δD) record from the EPICA Dome C (EDC) ice core⁸ back to Marine Isotopic Stage 20 (EDC3 timescale) showing Quaternary temperature variations in Antarctica. **b**, Vostok dust flux record (Coulter counter) plotted on its original timescale¹¹. **c**, EDC dust flux records. Red and grey lines represent, respectively, Coulter counter (55-cm to 6-m resolution) and laser-scattering data (55-cm mean). Numbers indicate

Marine Isotopic Stages. Note that the vertical extent of the scales of **b** and **c** is larger than for the other records. **d**, EDC dust size data expressed as FPP (see Methods). The orange and grey curves represent measurements by Coulter counter (2-kyr mean) and laser (1-kyr mean), respectively. **e**, Marine sediment $\delta^{18}O$ stack¹⁸, giving the pattern of global ice volume. **f**, Magnetic susceptibility stack record for Chinese loess¹⁷ (normalized).

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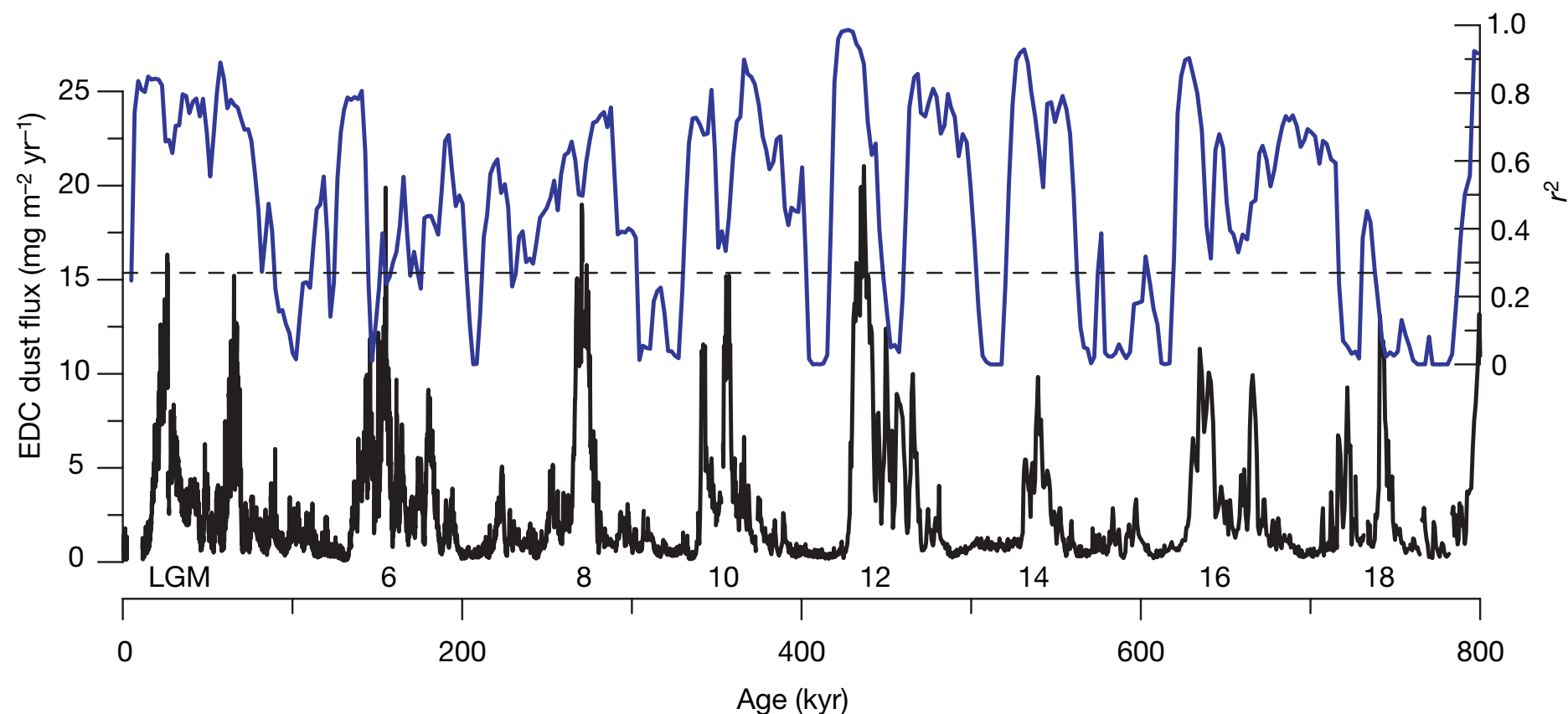


Figure 2 | EDC correlation between dust and temperature. Linear plot of dust flux (black) and the coefficient of determination r^2 (blue) between the high-pass filtered values (18-kyr cut-off) of both the δD and the logarithmic values of dust flux. The correlation was determined using 2-kyr mean values

in both records and a gliding 22-kyr window. Correlations above $r^2 = 0.27$ (dashed line) are significant at a 95% confidence level. Numbers indicate the marine isotopic glacial stages.

- strong correlation between dust and temperature changes

Dust depositions on glacial-interglacial time scales

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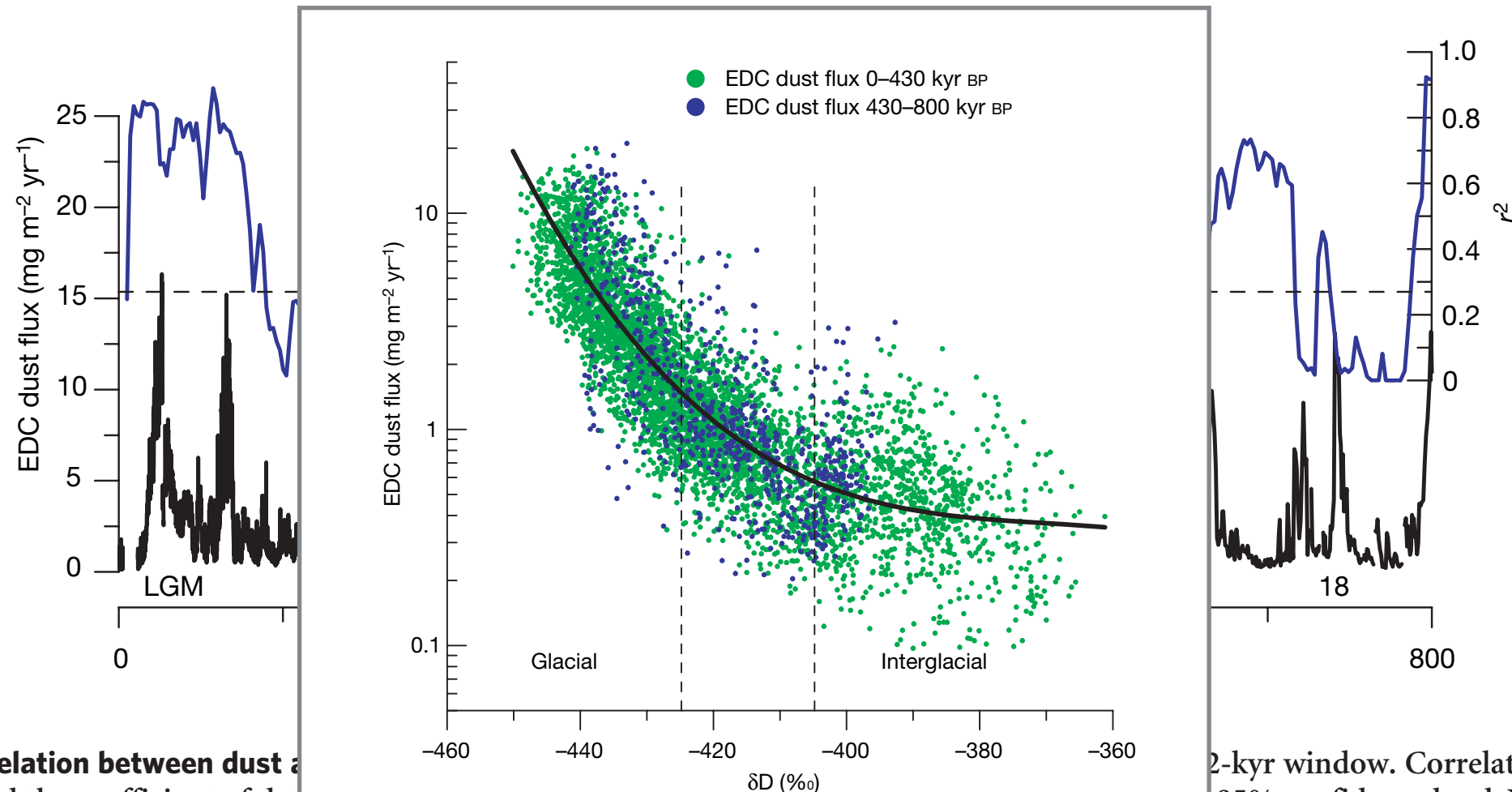


Figure 2 | EDC correlation between dust and temperature. EDC dust flux (black) and the coefficient of determination (blue) for a 2-kyr window. The correlation was determined for high-pass filtered values (18-kyr cut-off) of dust flux. The correlation was determined for high-pass filtered values (18-kyr cut-off) of dust flux. The correlation was determined for high-pass filtered values (18-kyr cut-off) of dust flux.

Figure 3 | EDC dust–temperature relationship. Values of δD (ref. 8) are plotted against dust flux (both at 55-cm resolution). Green and blue dots represent data from 0–430 kyr BP and 430–800 kyr BP, respectively. Superposed is a cubic polynomial fit, $\log_{10}(f) = -3.737 \times 10^{-6}(\delta D)^3 - 4.239 \times 10^{-3}(\delta D)^2 - 1.607(\delta D) - 204$, where f is the dust flux (mg m⁻² yr⁻¹), and δD is in ‰ ($r^2 = 0.73$, $N = 5,164$).

2-kyr window. Correlations above $r^2 = 0.27$ are above the 95% confidence level. Numbers indicate the

- strong correlation between dust and temperature changes

Dust depositions on glacial-interglacial time scales

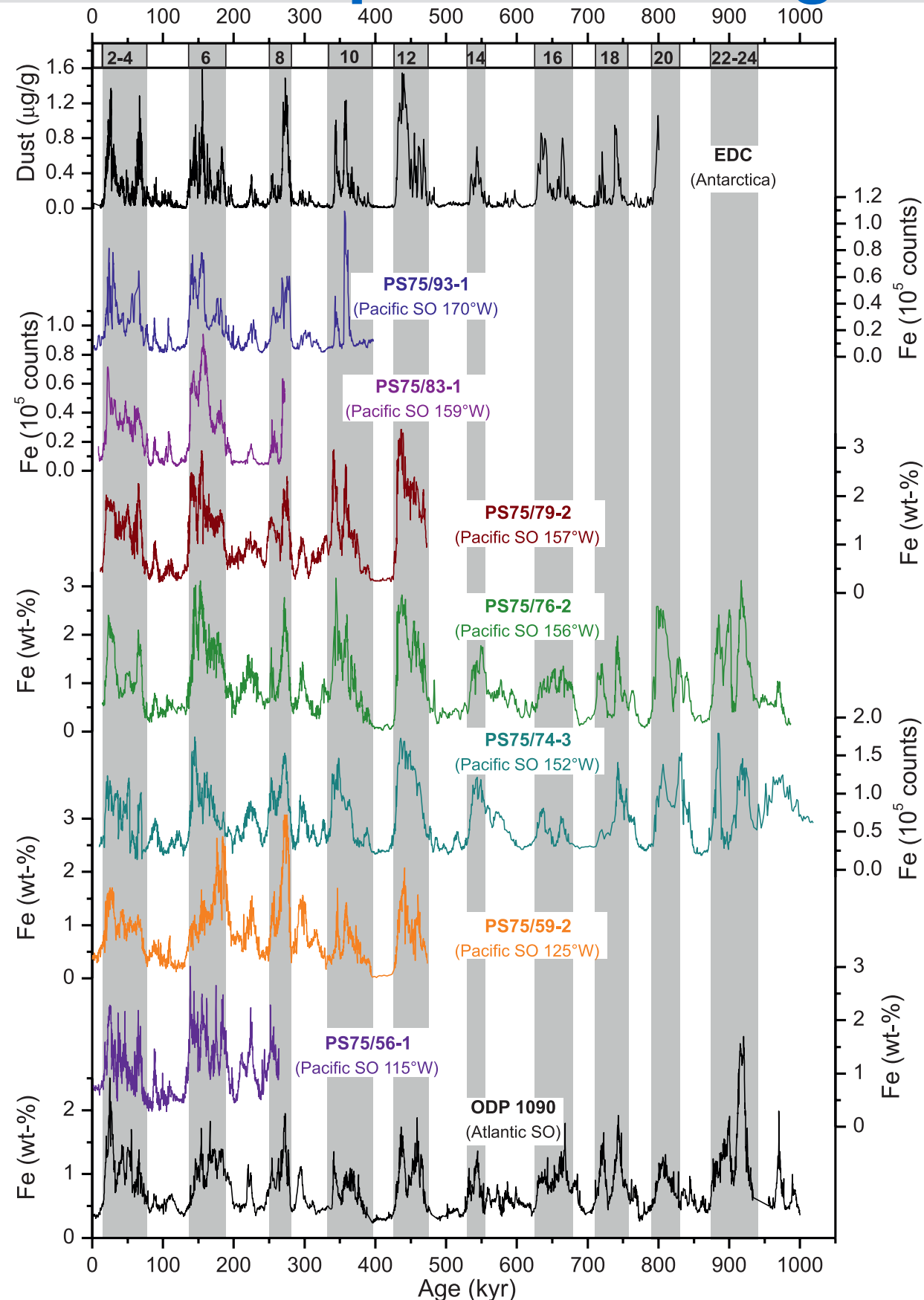


Fig. 2. Iron content fluctuations across the Pacific and Atlantic SO (7) compared to dust content changes in the EDC ice core (1).

SCIENCE VOL 343 24 JANUARY 2014

Increased Dust Deposition in the Pacific Southern Ocean During Glacial Periods

F. Lamy,^{1,2*} R. Gersonde,^{1,2} G. Winckler,^{3,4} O. Esper,¹ A. Jaeschke,^{1,2} G. Kuhn,¹ J. Ullermann,¹ A. Martinez-Garcia,⁵ F. Lambert,⁶ R. Kilian⁷

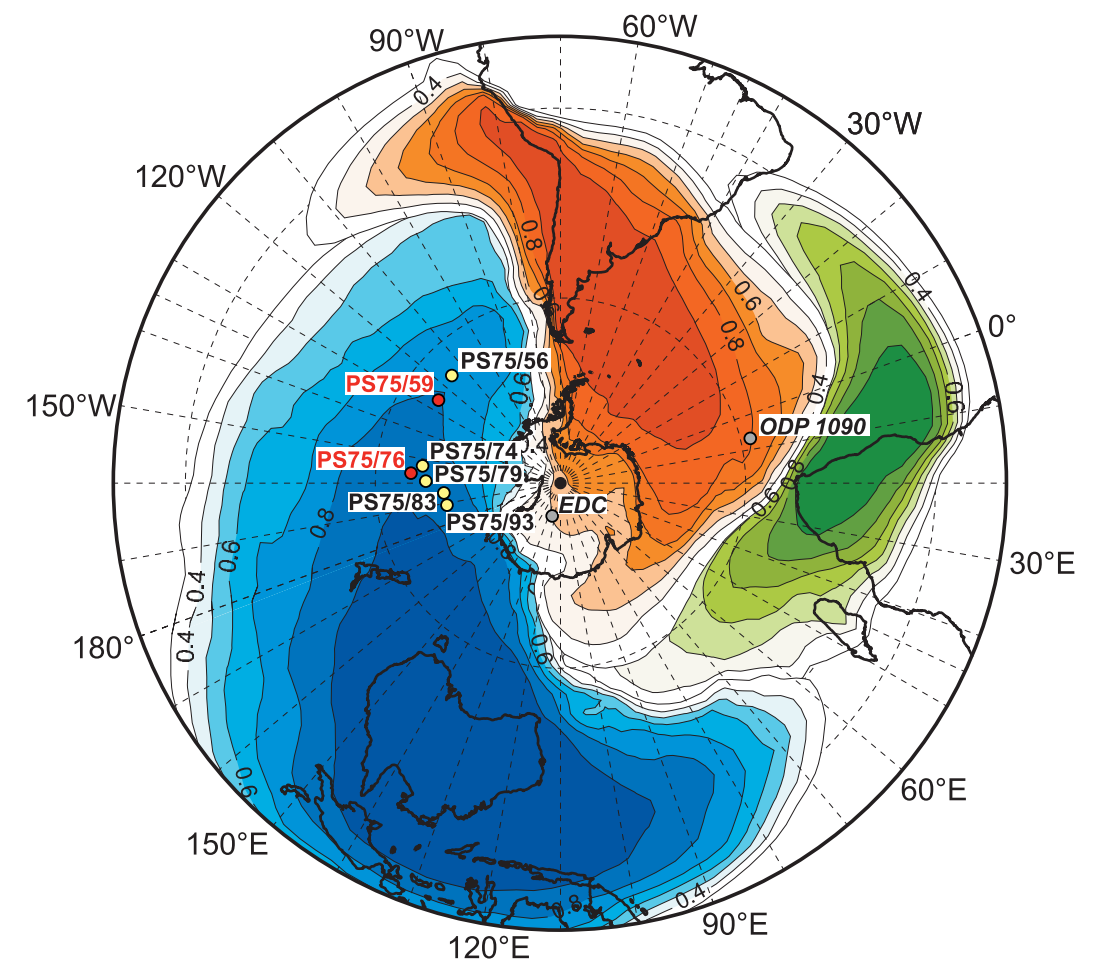


Fig. 1. Map showing the modern relative contributions of the three major dust sources in the Southern Hemisphere (blue, Australia; red, South America; green, South Africa), based on model data (20). Red dots mark primary core locations; yellow dots indicate additional cores; gray dots denote the location of published reference records (1, 4, 7).

LGM dust cycle: fertilisation of the marine biosphere



John Martin

The Iron Hypothesis

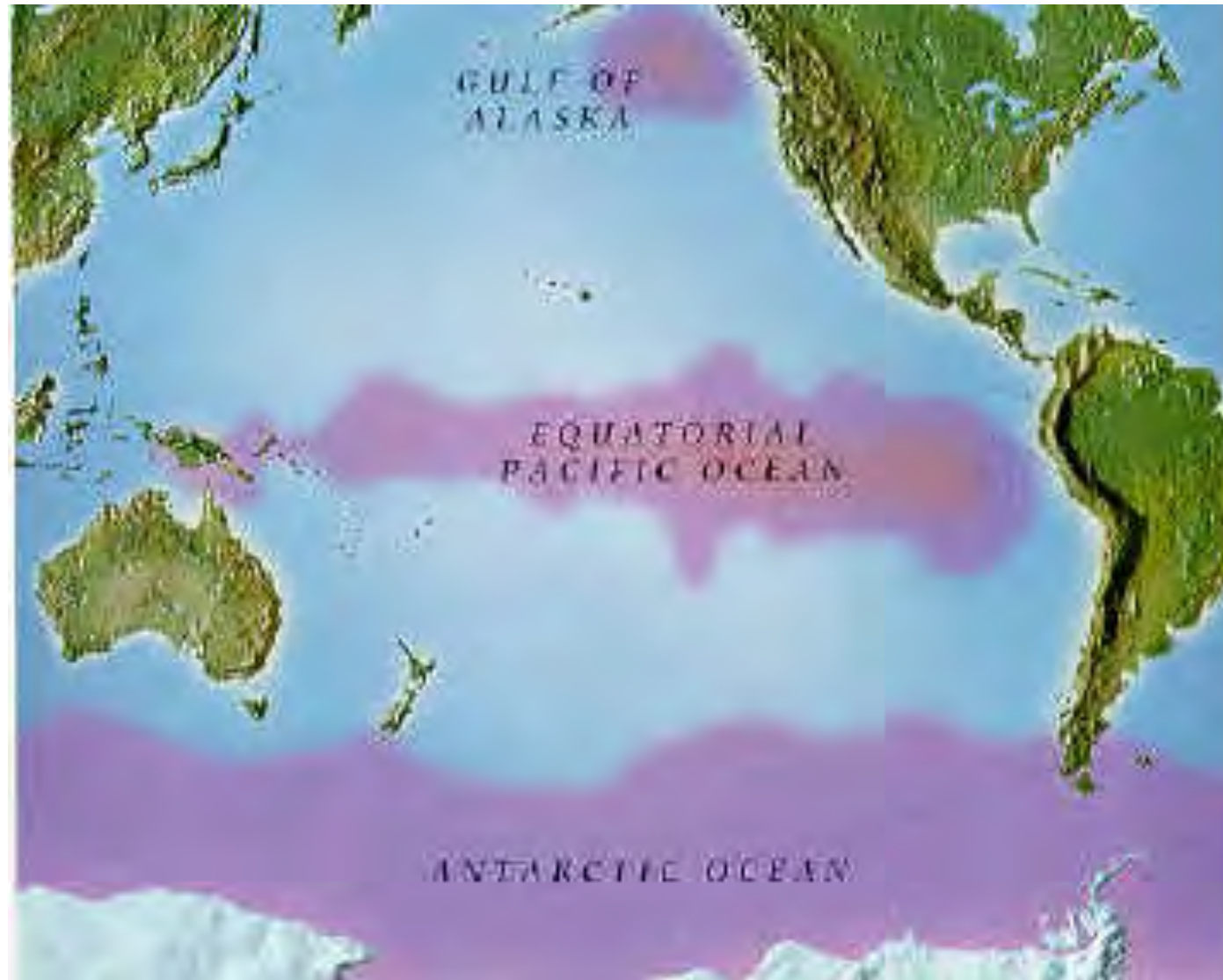


Figure 1. In the equatorial Pacific Ocean and Gulf of Alaska, phytoplankton populations are relatively low (purple shaded areas on map), despite adequate sunlight and nutrients. John Martin set out to prove that a lack of dissolved iron in the water in these areas keeps populations of marine algae lower than normal.

- **Dust contains iron which is a key micro-nutrient in the Southern Ocean**

LGM dust cycle: fertilisation of the marine biosphere



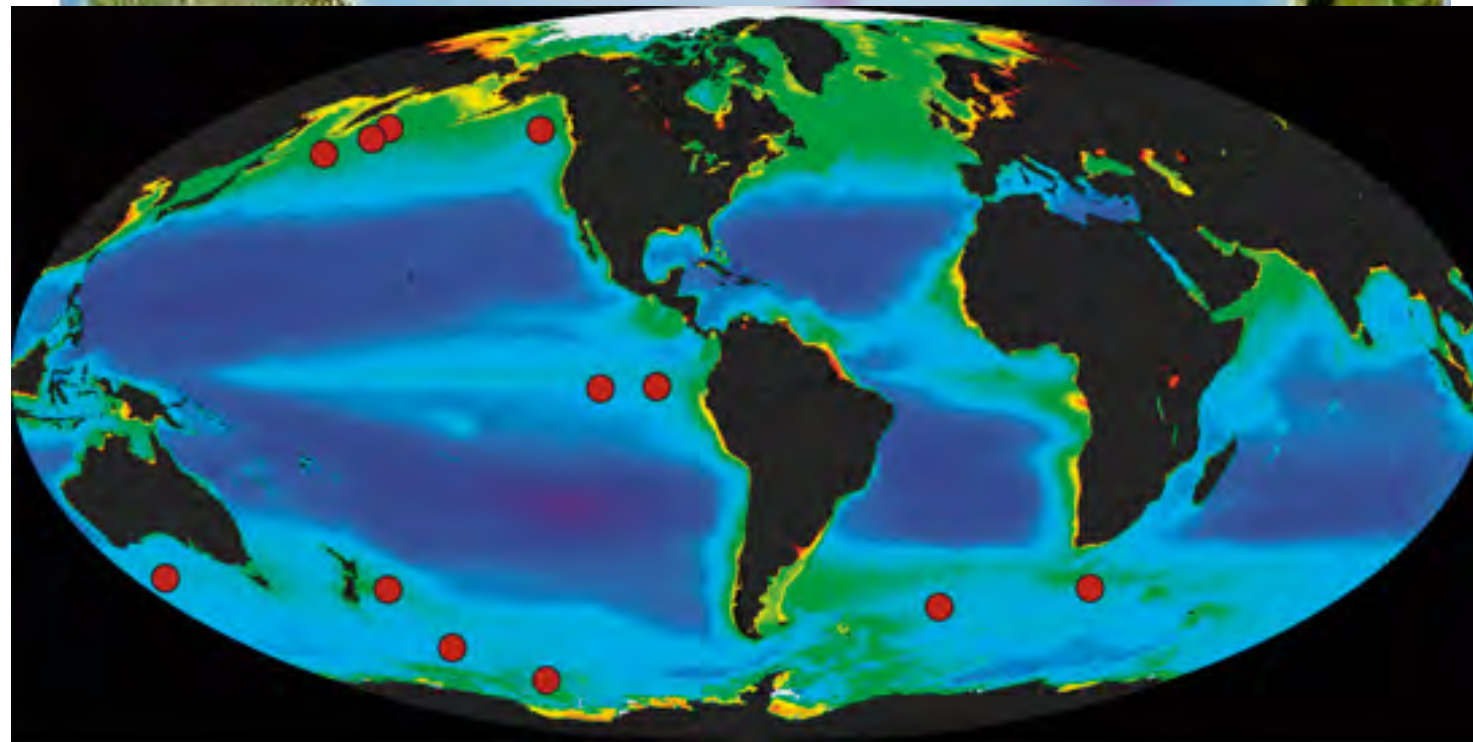
John Martin

The Iron Hypothesis



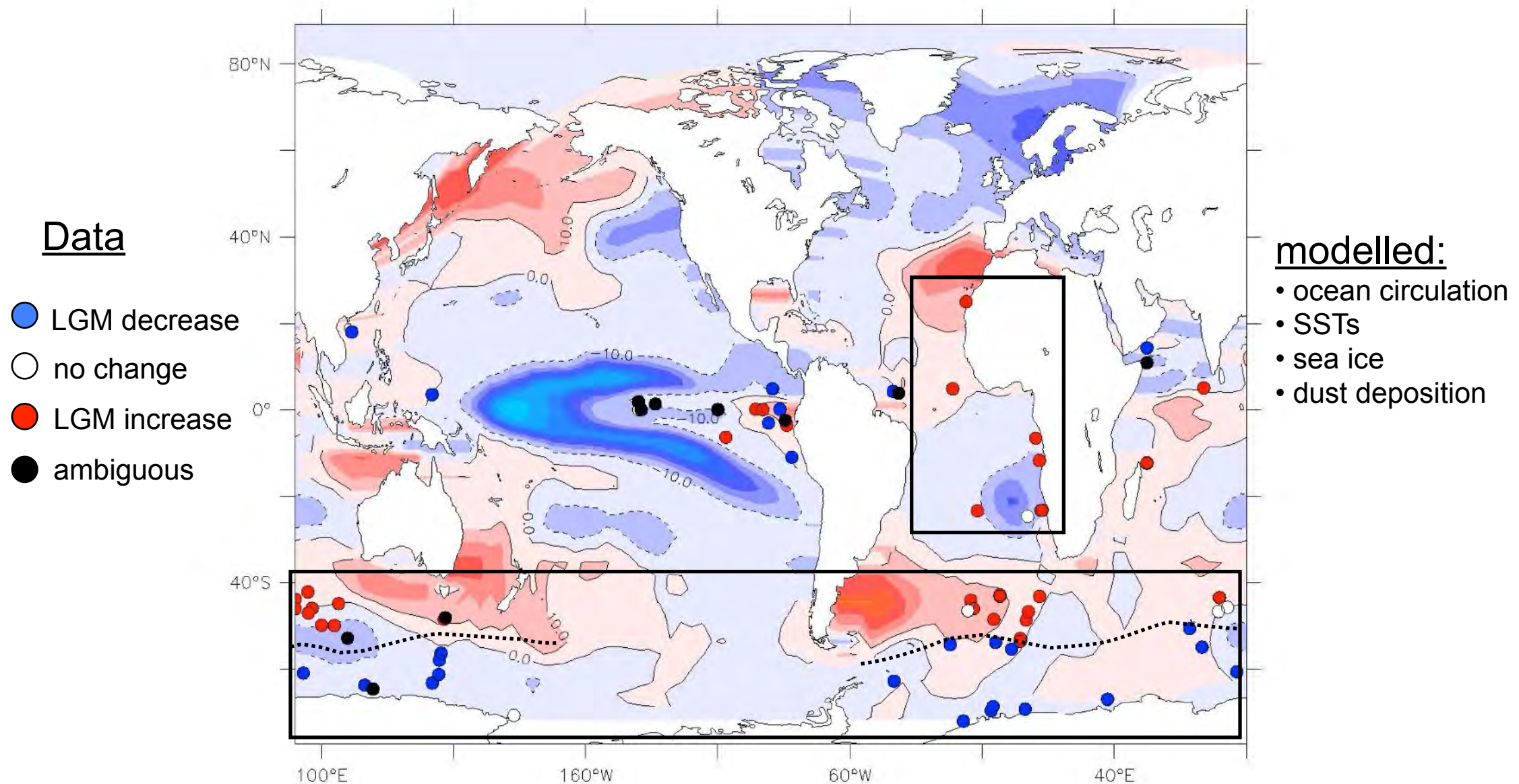
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Since 1993, these 12 small-scale open ocean experiments (red dots) have shown that **iron additions do indeed result in phytoplankton blooms, thereby drawing carbon dioxide out of the atmosphere and into the ocean**



LGM dust cycle: fertilisation of the marine biosphere

Ocean Carbon Cycle: LGM-Modern Changes in Export Production



Data Compilation

(proxies : C-org, CaCO₃, Ba, assemblages...)

/ LGM simulation with PISCES-OPA model

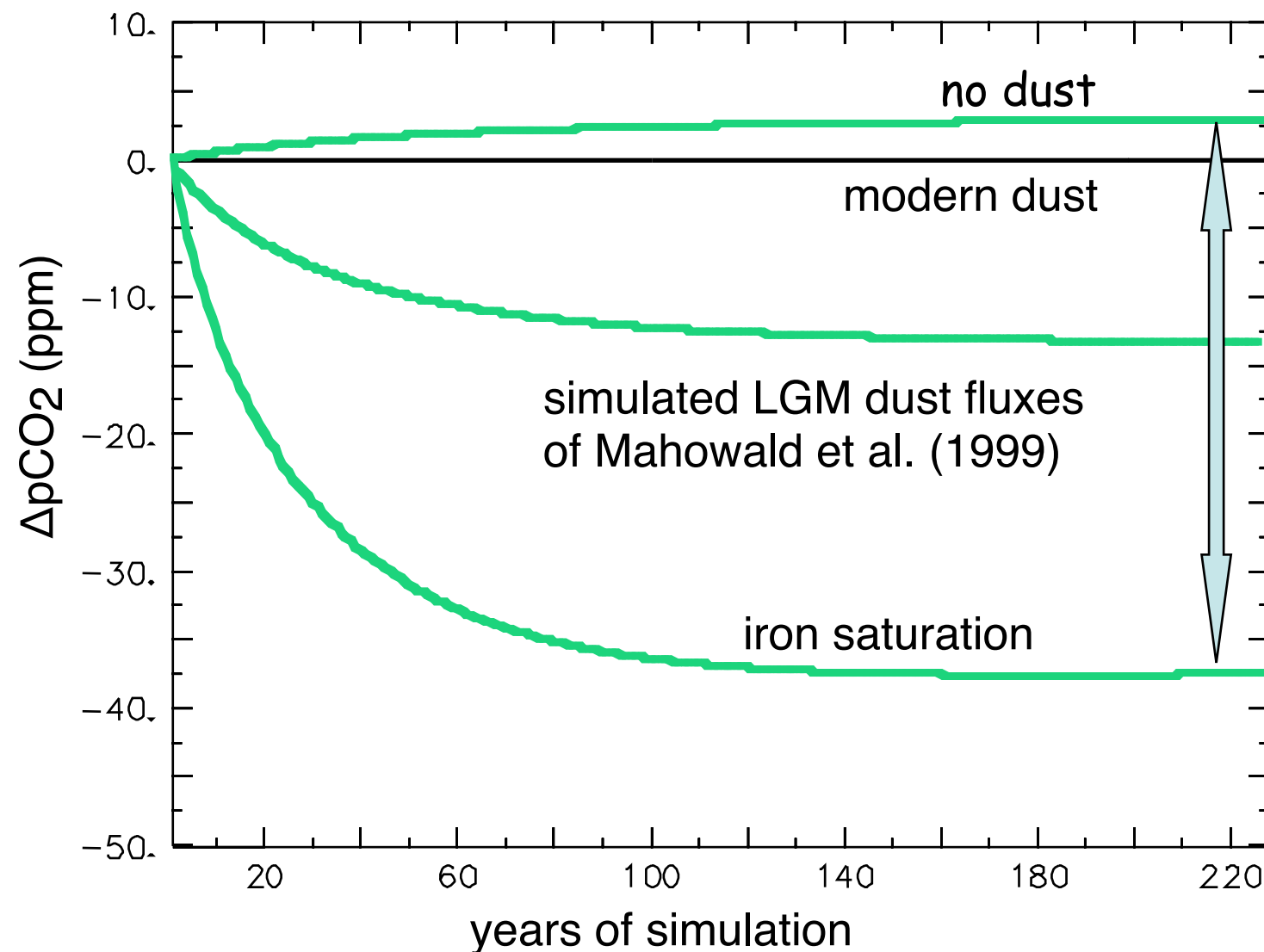
(Export, LGM-present, gC m⁻² yr⁻¹)

(Quelle: Bopp et al., 2003)

LGM dust cycle: fertilisation of the marine biosphere

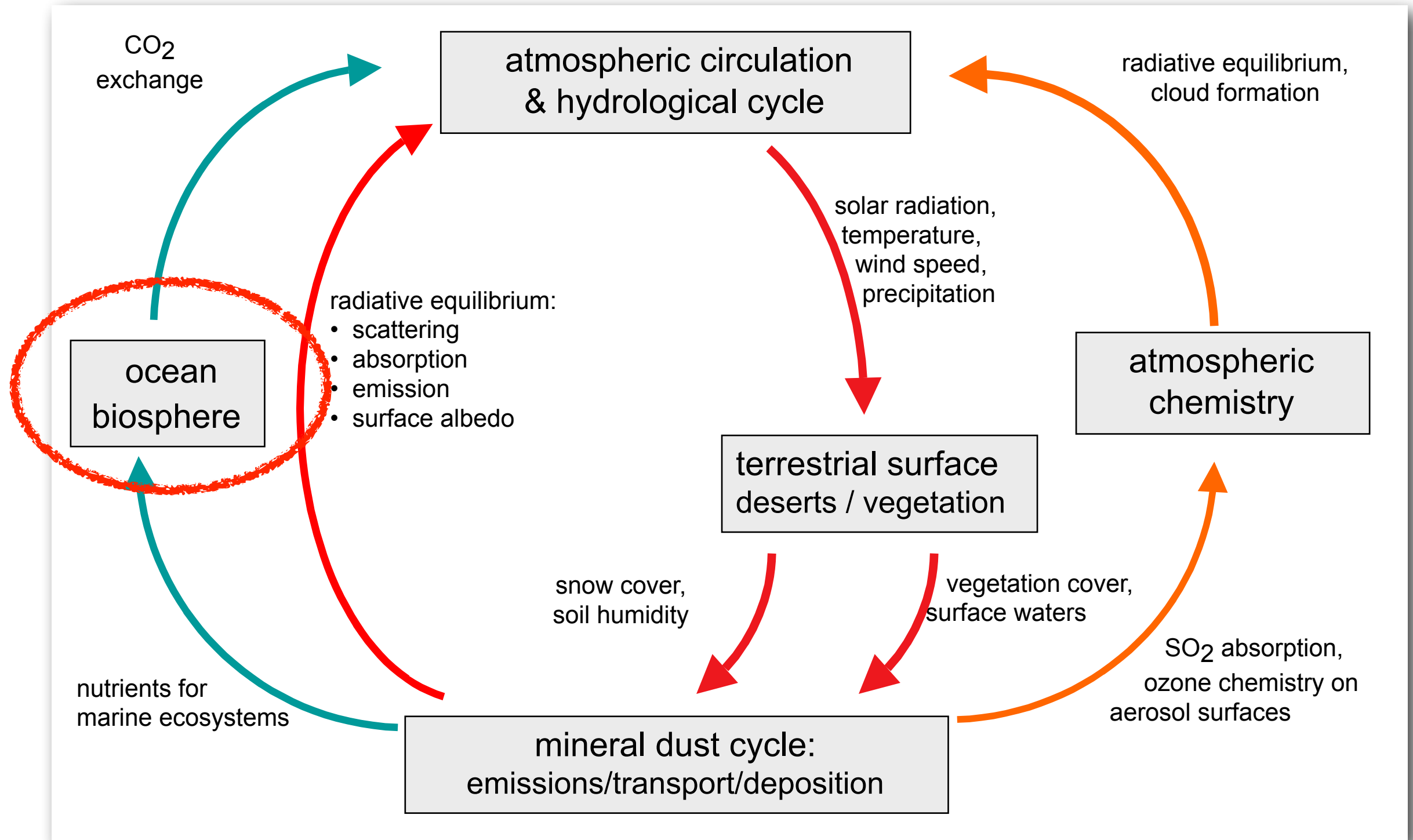
Ocean Carbon Cycle: Sensitivity Studies for different LGM conditions

- Hypothesis of Martin (1988) : Iron Fertilisation of the Glacial Ocean

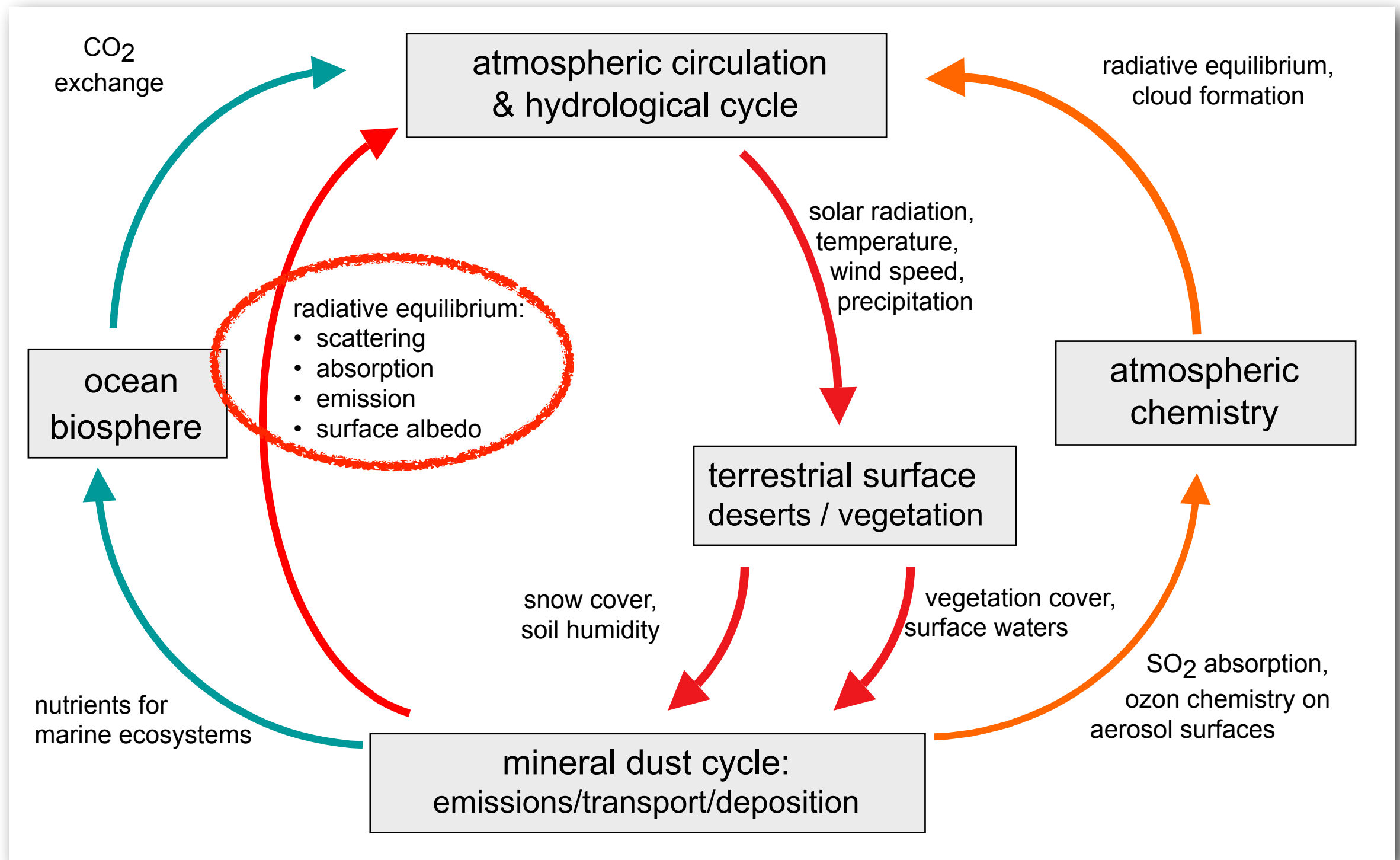


- increased LGM iron deposition: -15 ppm
- mechanisms
 - export of organic material
 - export of CaCO₃ (change of alkalinity)
- sensitivity
 - no dust deposition: +3 ppm
 - iron saturation: -39 ppm

Dust in the climate system

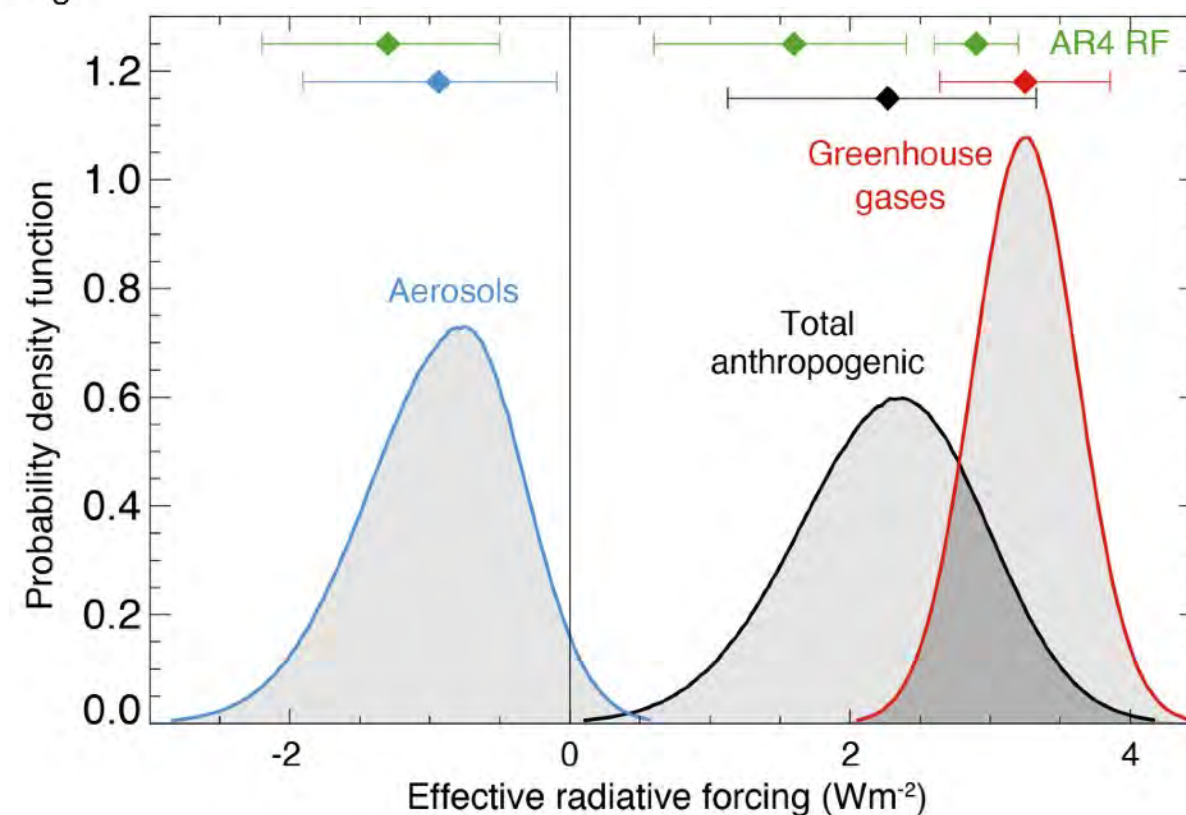
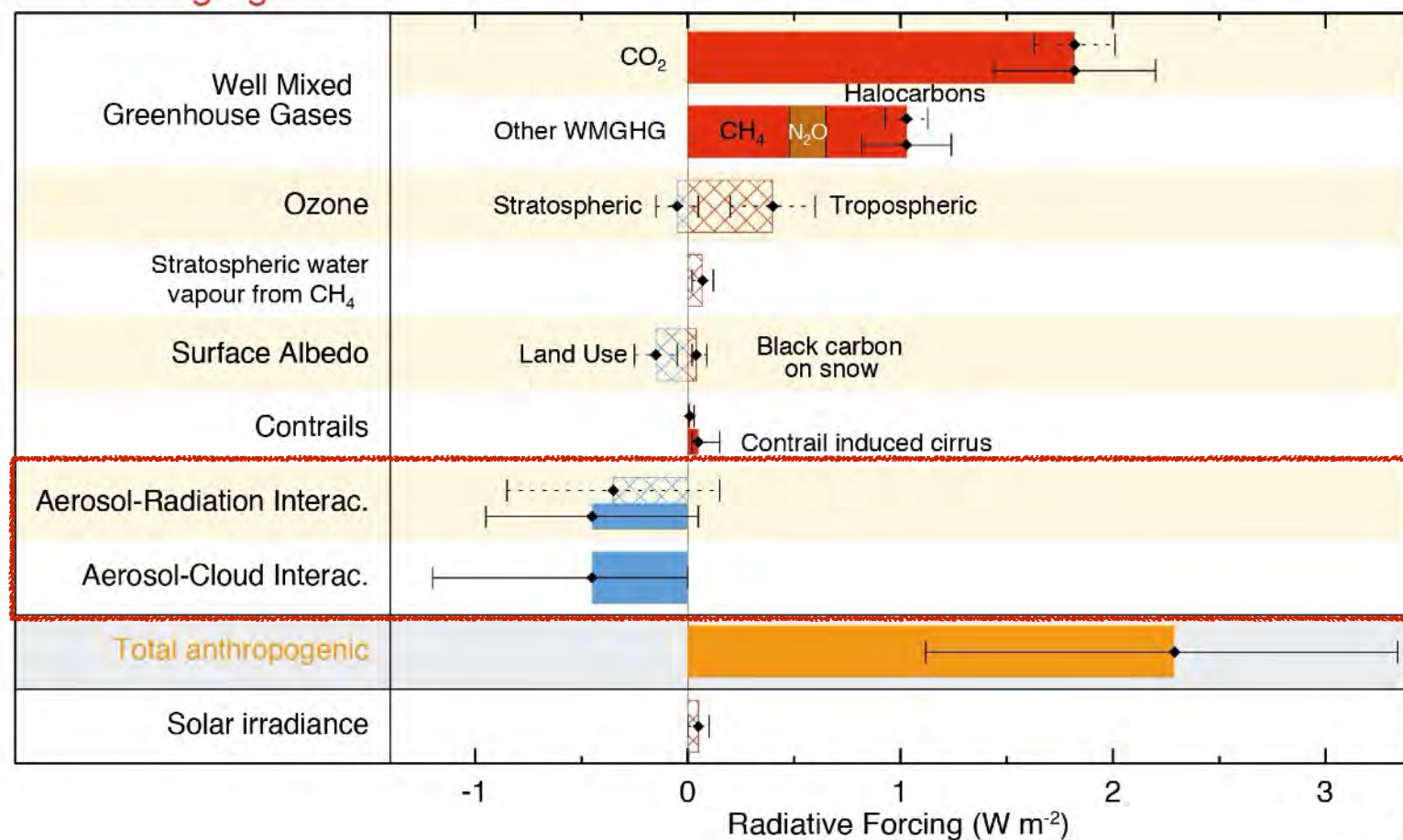


Dust in the climate system

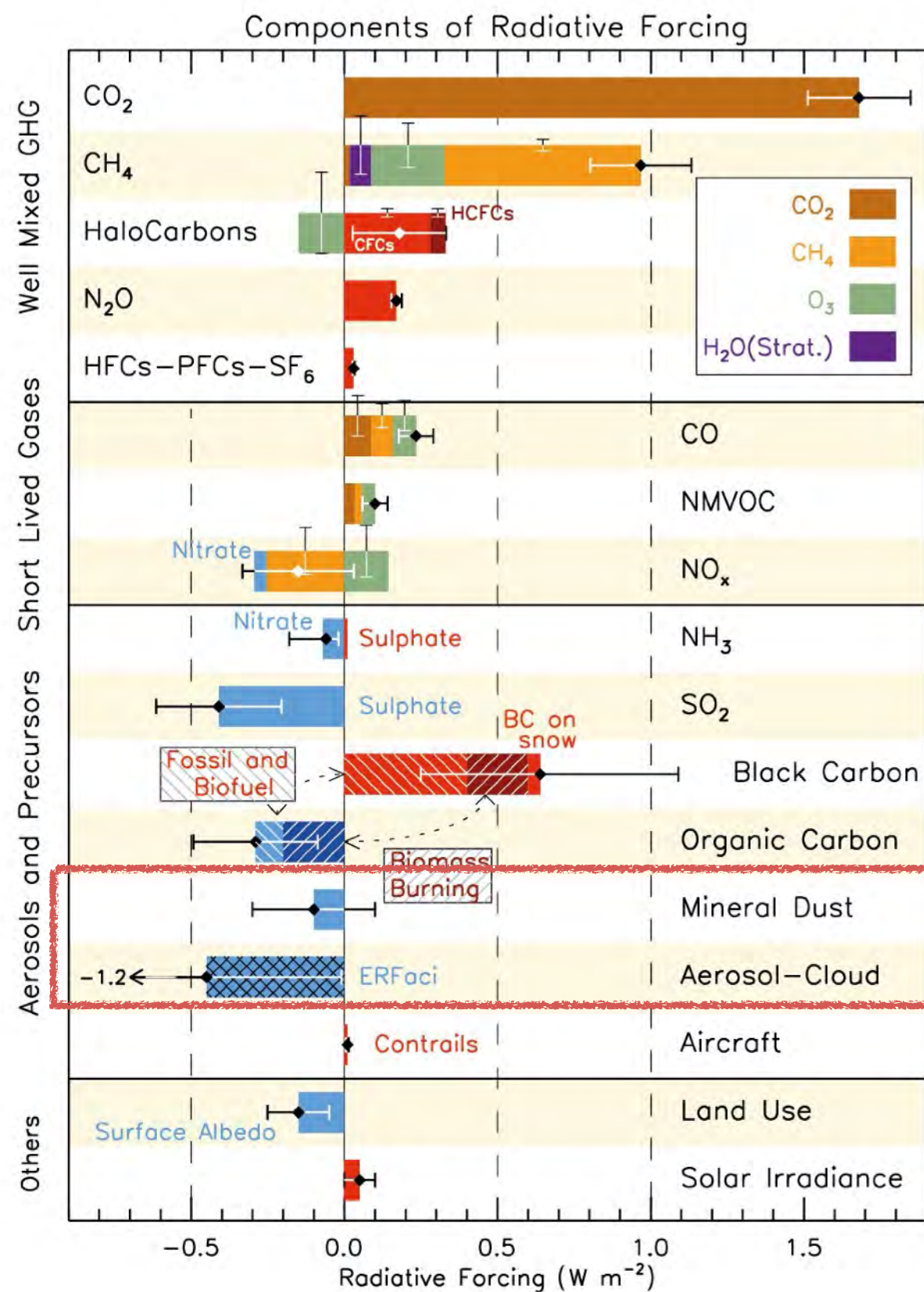


Present-day radiative forcing

Radiative forcing of climate between 1750 and 2011
Forcing agent

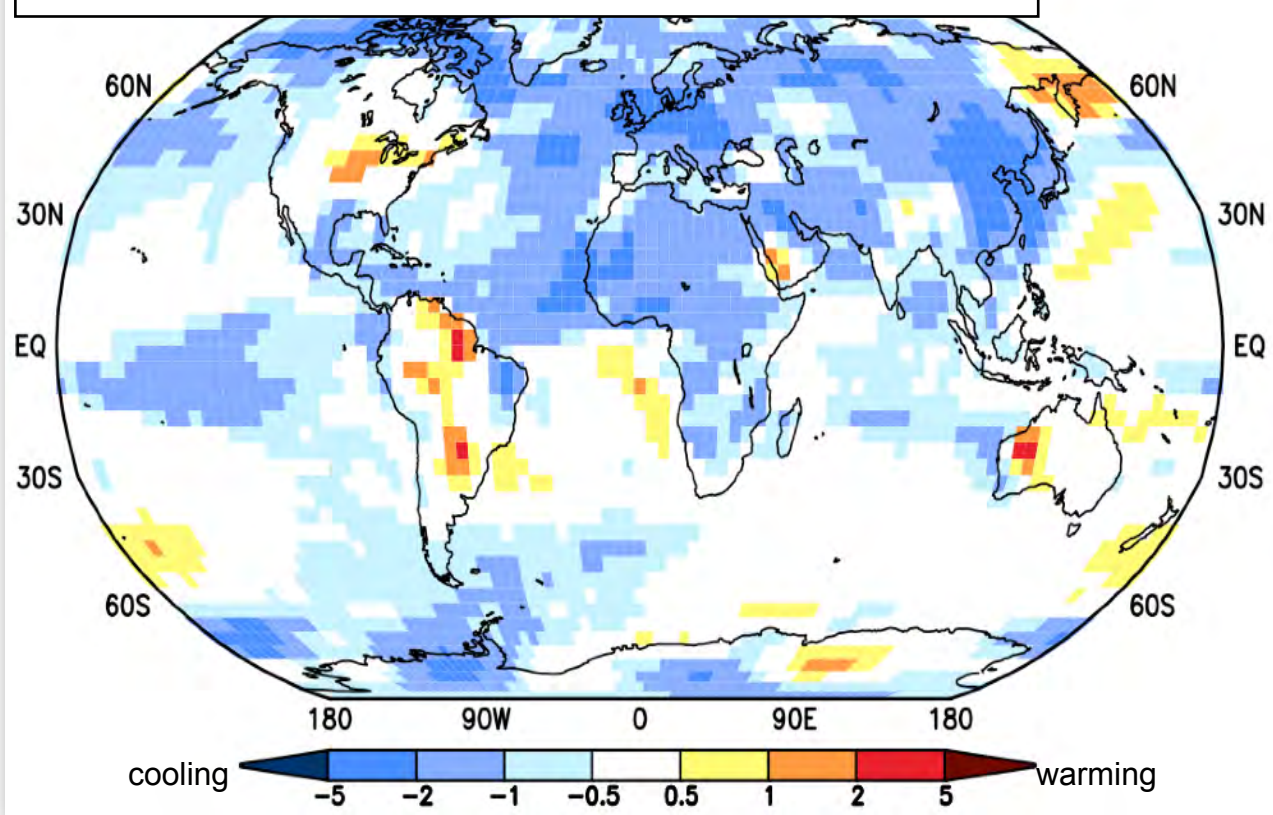


The radiative effect of mineral dust



Changes of the LGM radiative budget by dust

LGM temperature changes by dust radiative forcing [K]



- largest radiative cooling effect near the source regions (between 45°S und 45°N)
- no *linear* relation between radiative effect and temperature changes
- feedback processes of radiative changes to the dust cycle (emission/transport/deposition) are relatively small

optical properties of mineral aerosols:

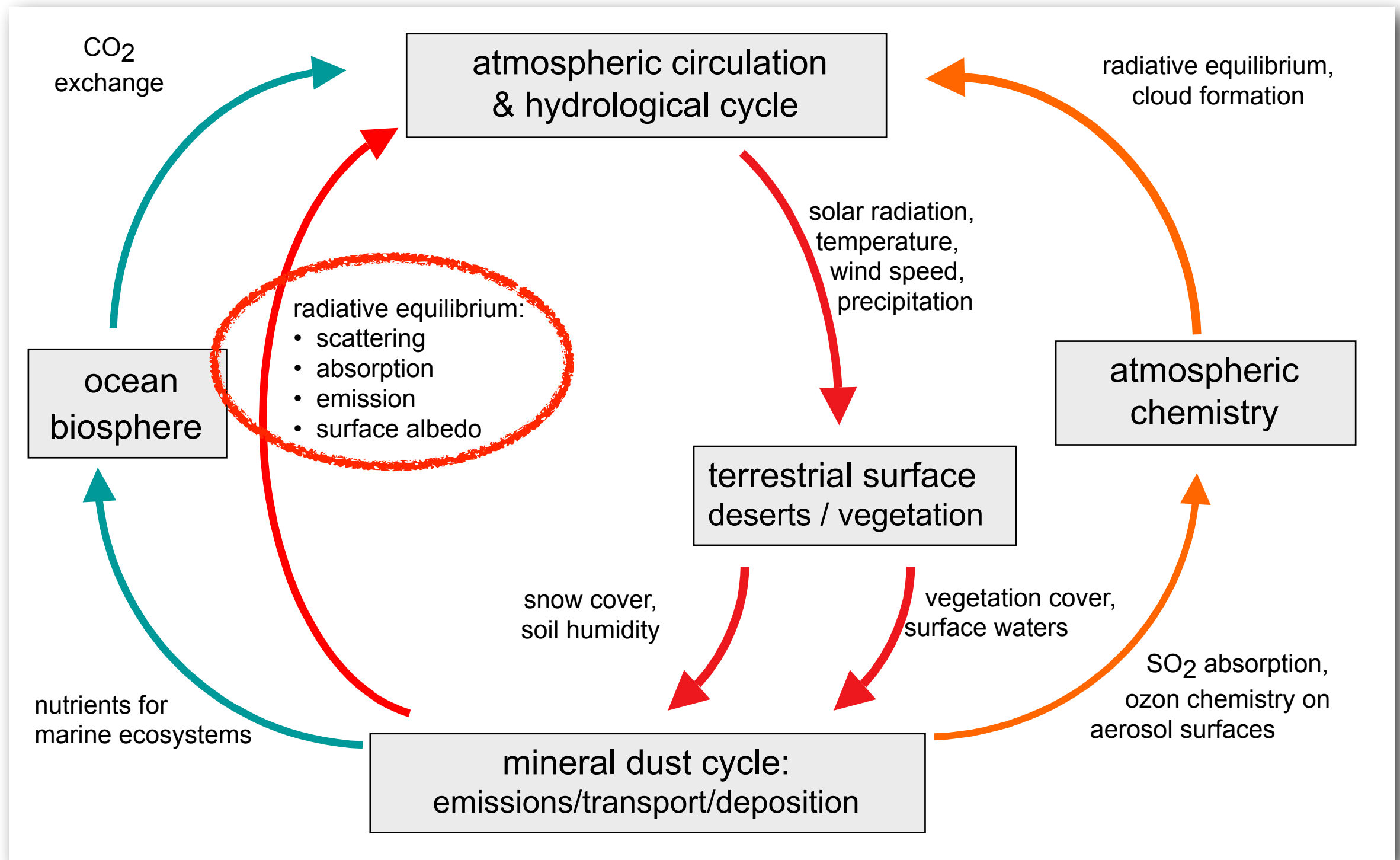
- only aerosol particles between 0.3 - 4 μm have been considered
- uniform mineralogy: 98% clay, 2% hematite, intern mixing
- refractive indices based on Sokolik und Toon (1999)

Zonal changes of radiative forcing and temperature:

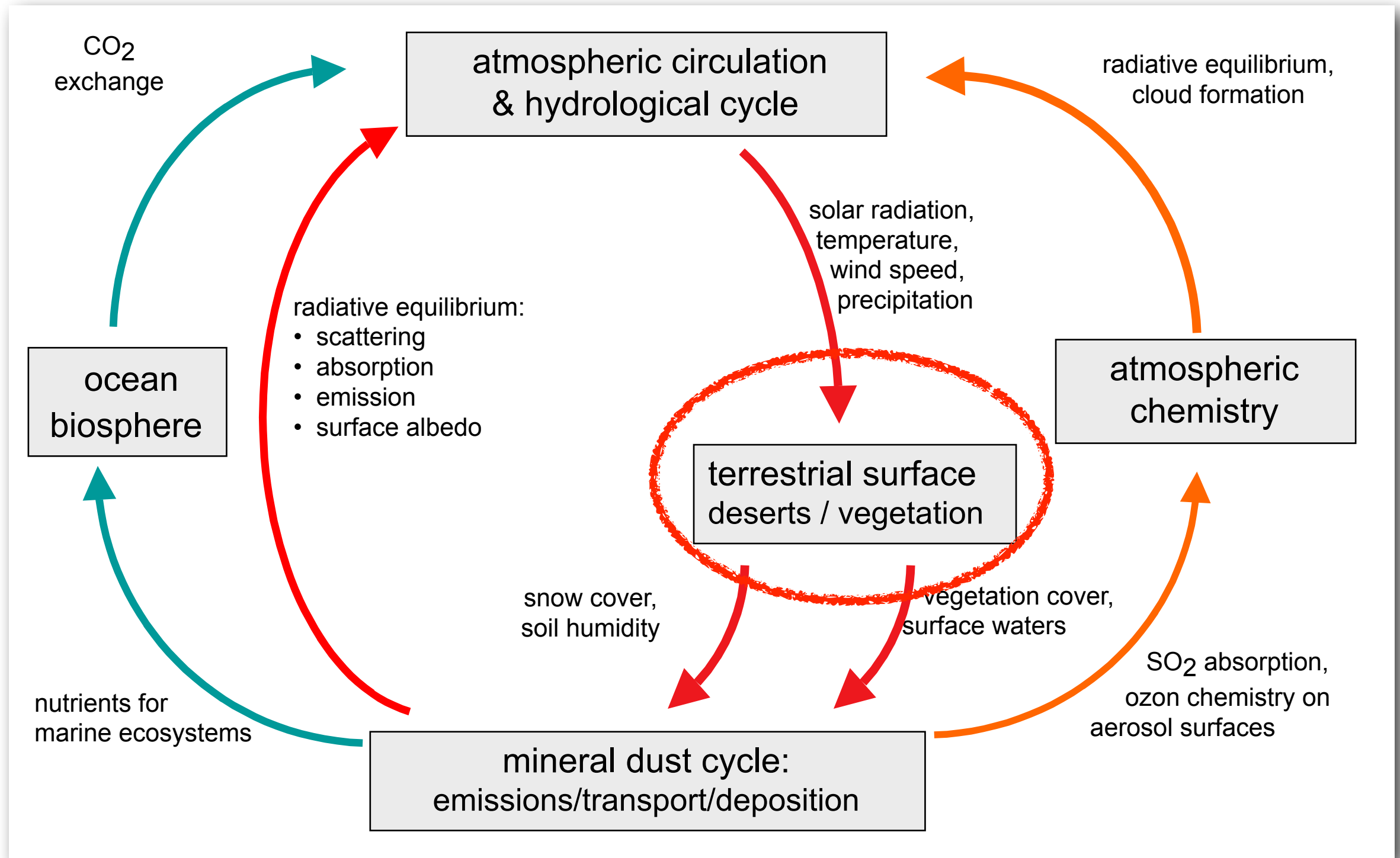
	Claquin et al. [2003]		ECHAM5-HAM	
	TOA forcing [W/m ²]	TOA forcing [W/m ²]	ΔT_{surf} [K] (by dust)	ΔT_{surf} [K] (CO ₂ , albedo*, ice)
Global	-1	-1.4	-0.6	-4.3
90°N-45°N	-0.3	-0.6	-1.1	-14.6
45°N-45°S	-1.6	-1.8	-0.5	-1.7
45°S-90°S	+0.2	-0.8	-0.4	-6.3

(*change of albedo by dust deposition has been neglected)

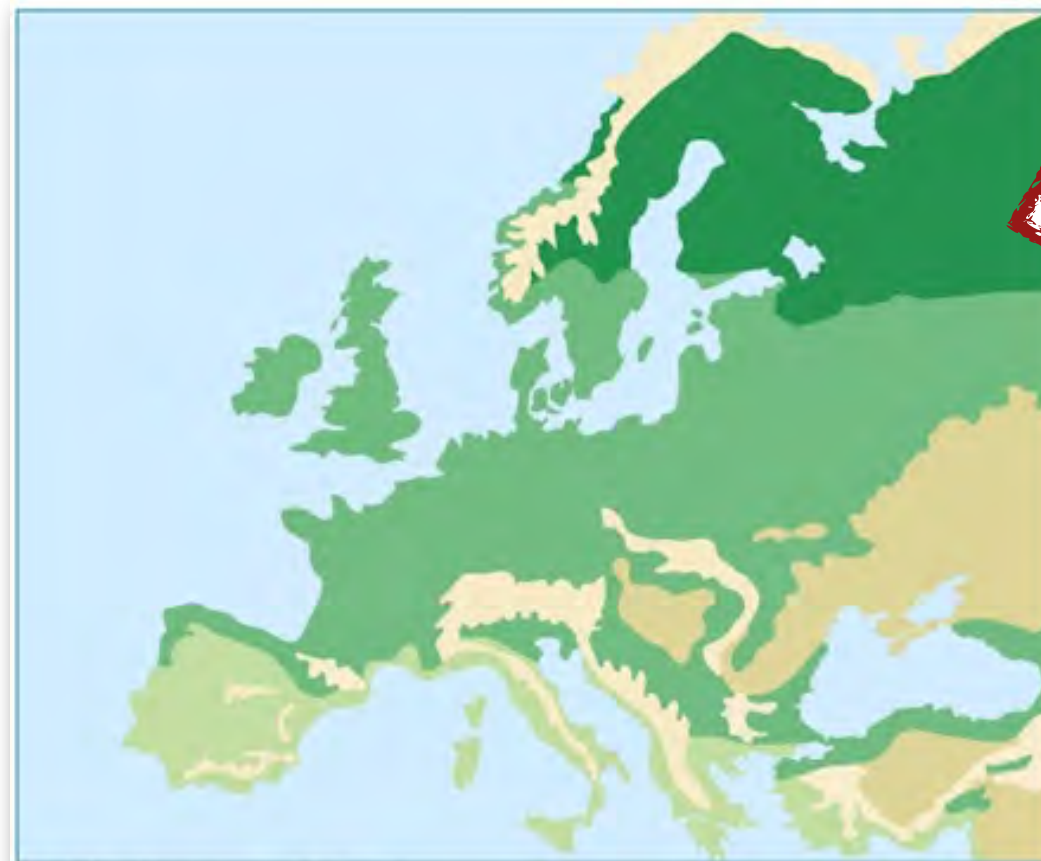
Dust in the climate system



Dust in the climate system



LGM climate of North America and Europe



A Modern vegetation

□ Ice	■ Boreal forest	■ Mediterranean scrub
■ Tundra and mountain	■ Deciduous and conifer forest	■ Prairie-steppe

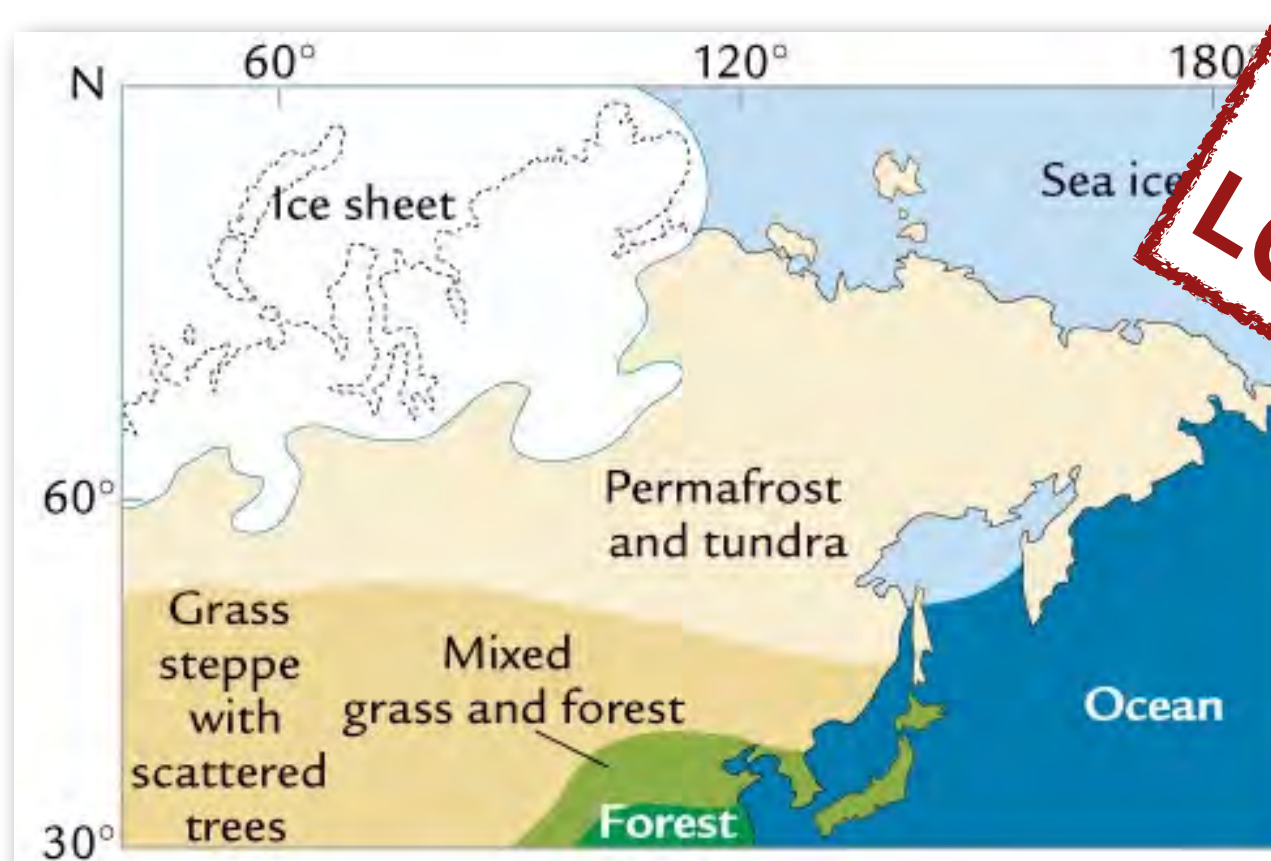


B Glacial vegetation

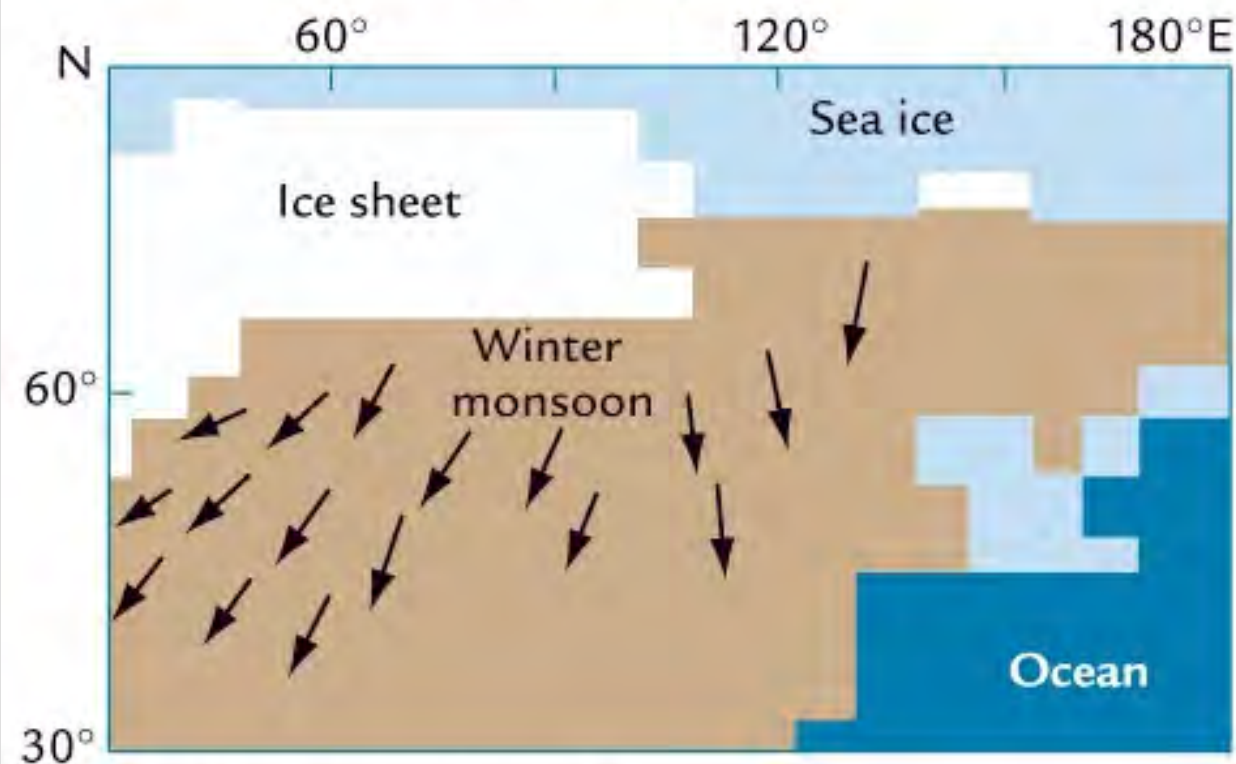
[from: Ruddiman, 2008]

remember
LGM lecture!

LGM climate of Asia



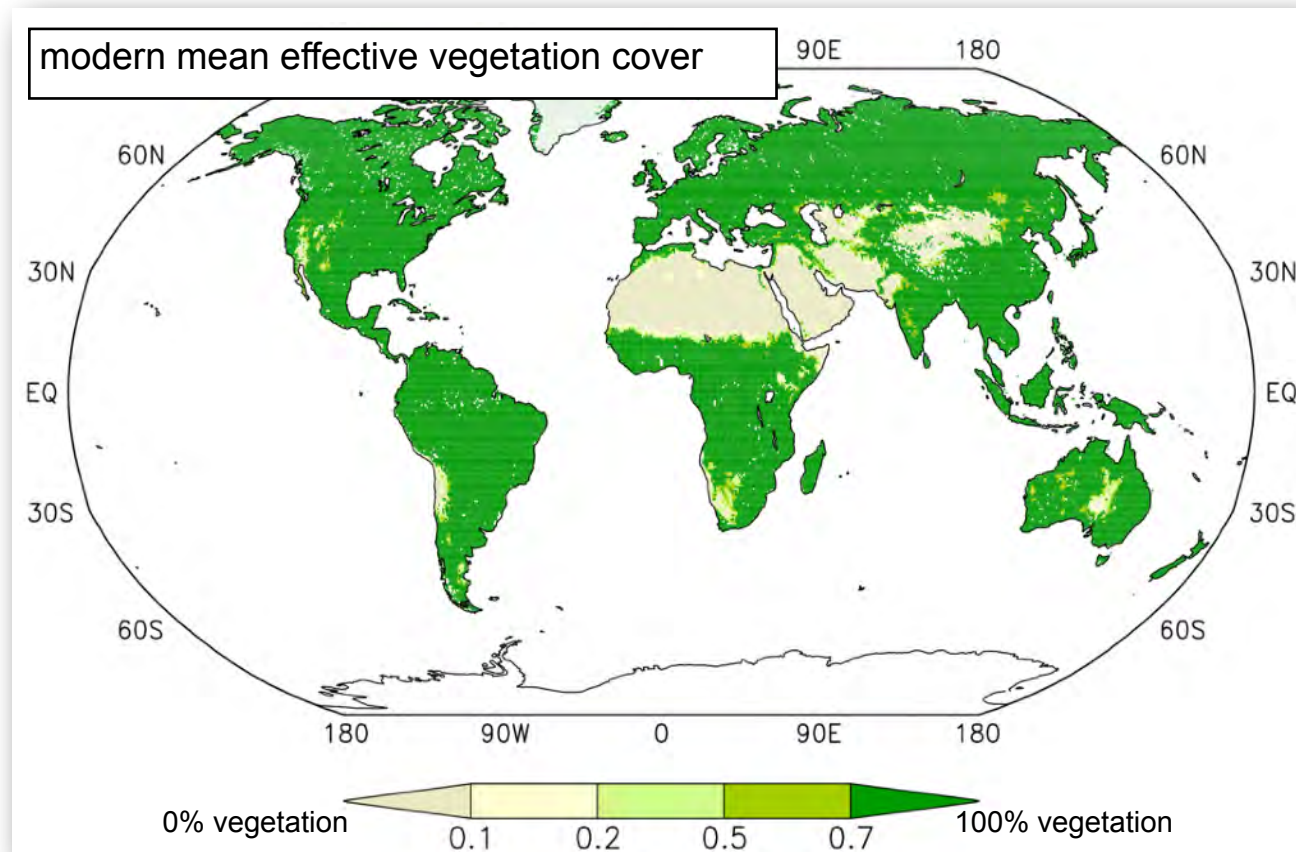
A Glacial maximum (observed)



B Glacial maximum (model)

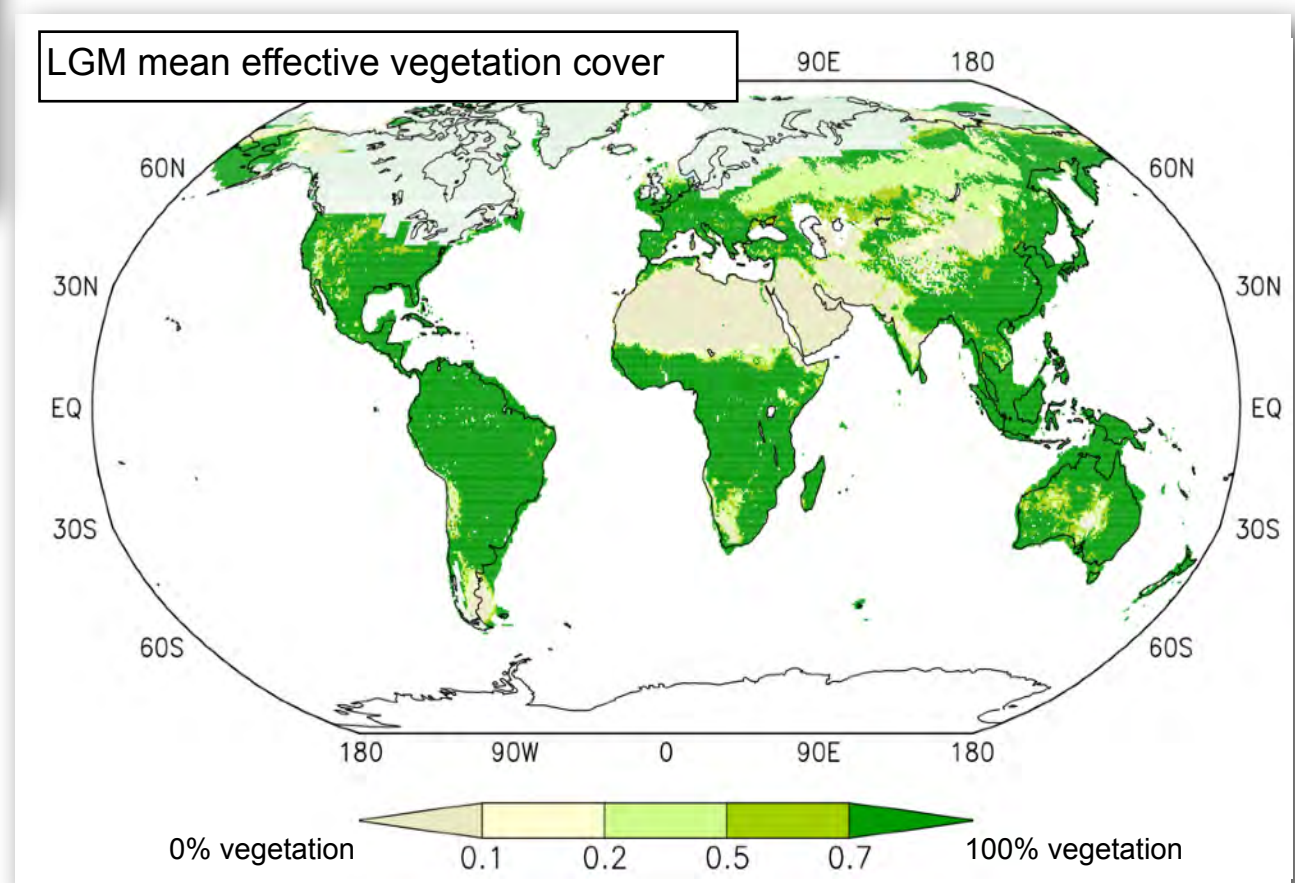
**remember
LGM lecture!**

Simulation of glacial vegetation changes

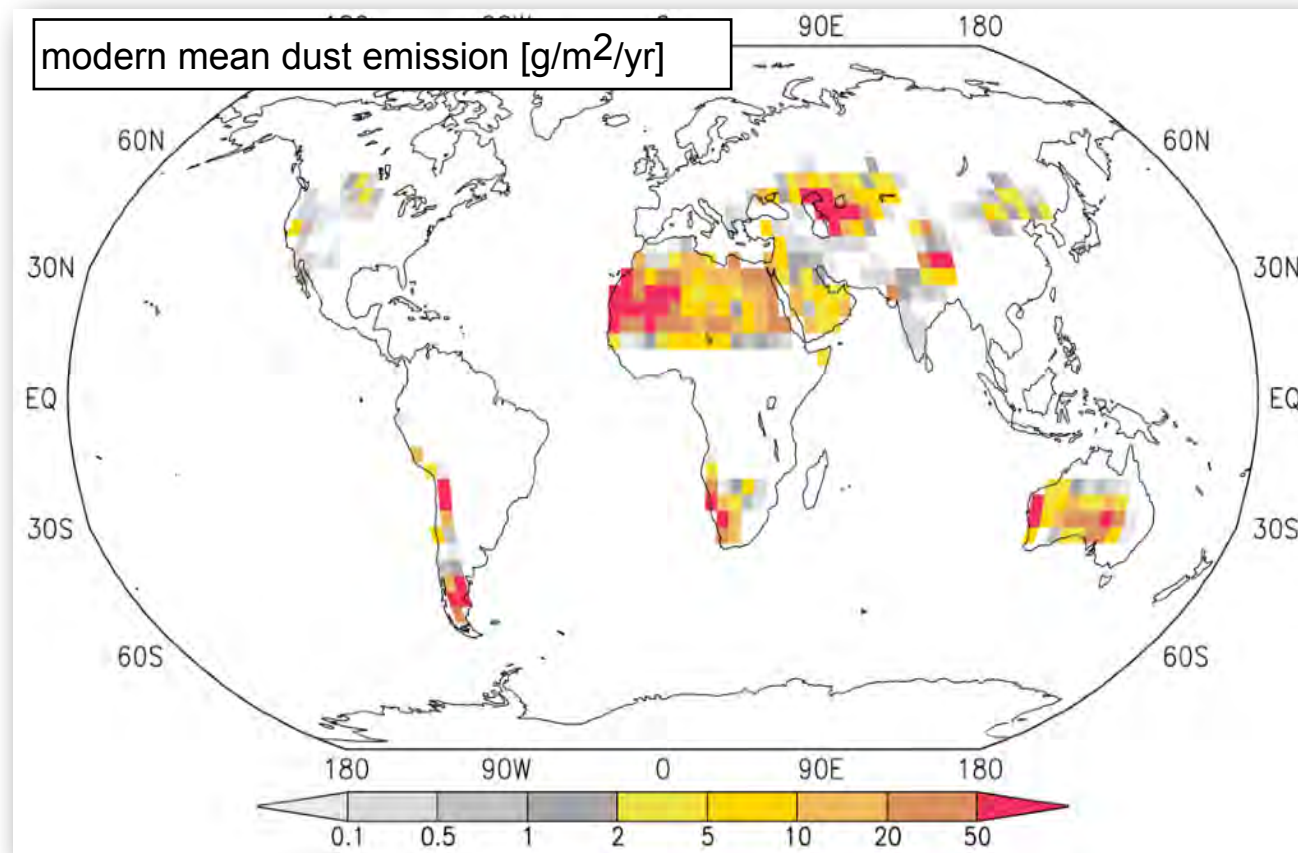


changes in glacial vegetation cover:

- major decrease in vegetation cover in central and northern Asia, South America and lesser changes in Australia
- (almost) no change of vegetation cover in present-day desert regions, e.g. Sahara

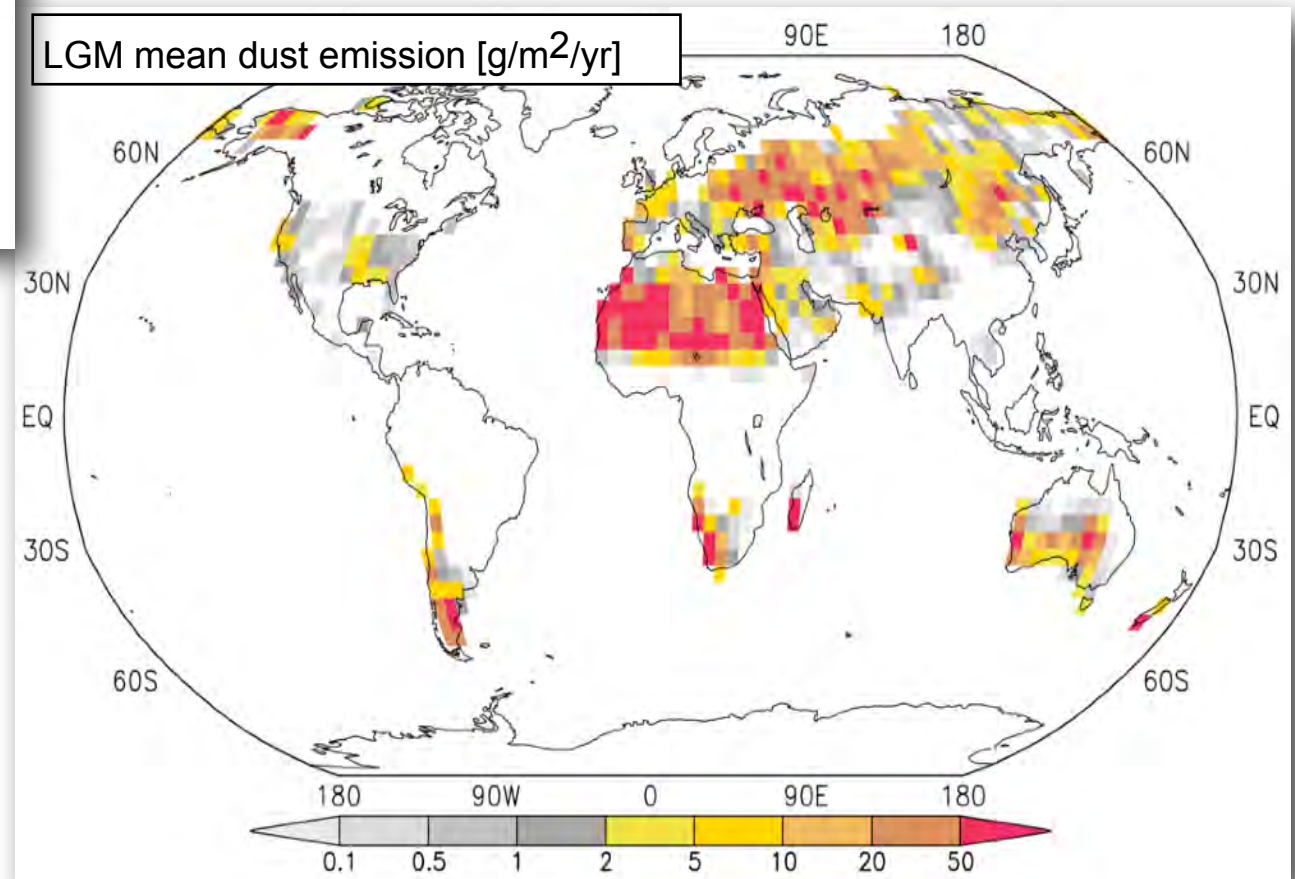


Simulation of glacial dust emissions



changes of glacial dust emissions:

- global emission increase of factor $\sim 2-3$
- glacial increase is not uniform, but varies for the different source regions

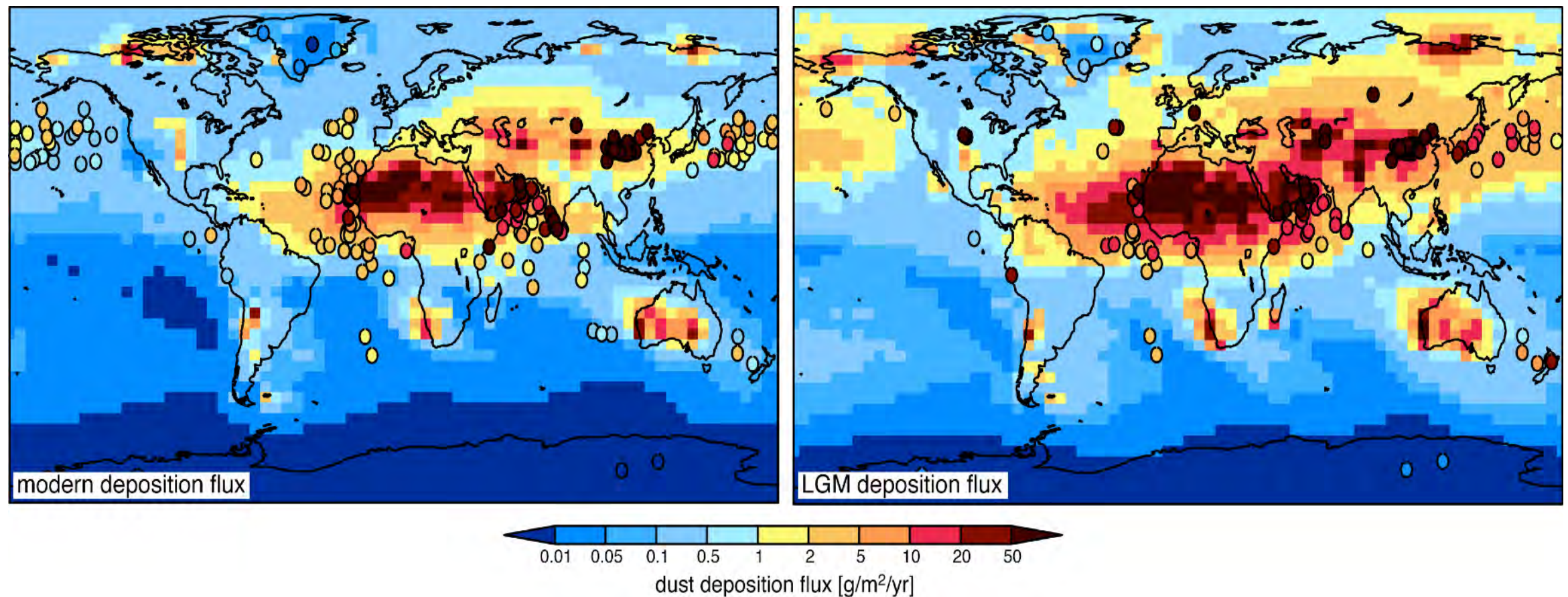


- **Sahara:** increased dust emissions by increased number and intensity of dust storms
- **Asia:** increased dust emissions by increase of (potential) source areas
- seasonal changes of LGM dust emissions are mainly related to changes in vegetation cover

Modern and LGM dust deposition fluxes

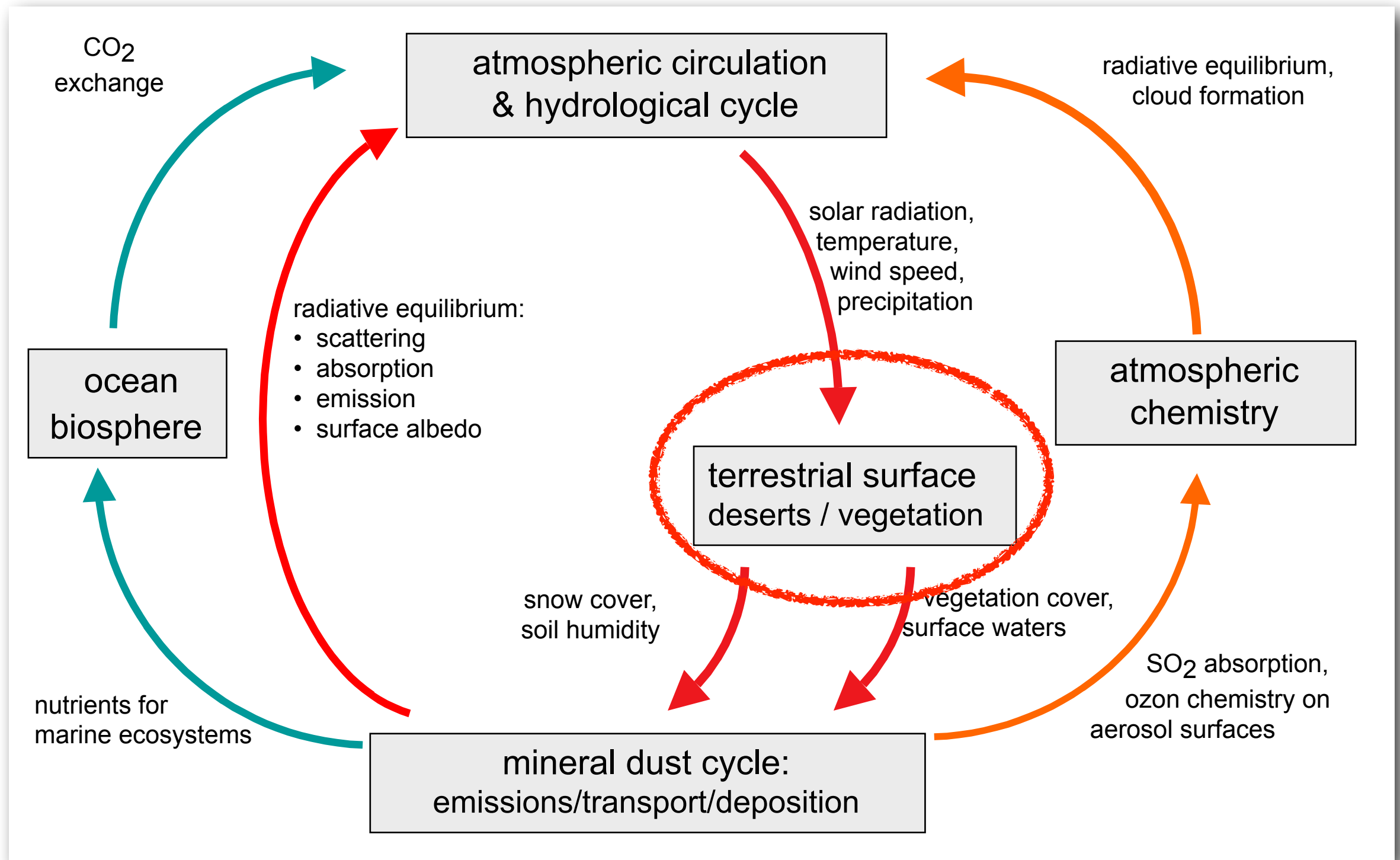
simulation: ECHAM-4 GCM climatology (modern and LGM, respectively)

reconstructions: DIRTMAP Datenbank

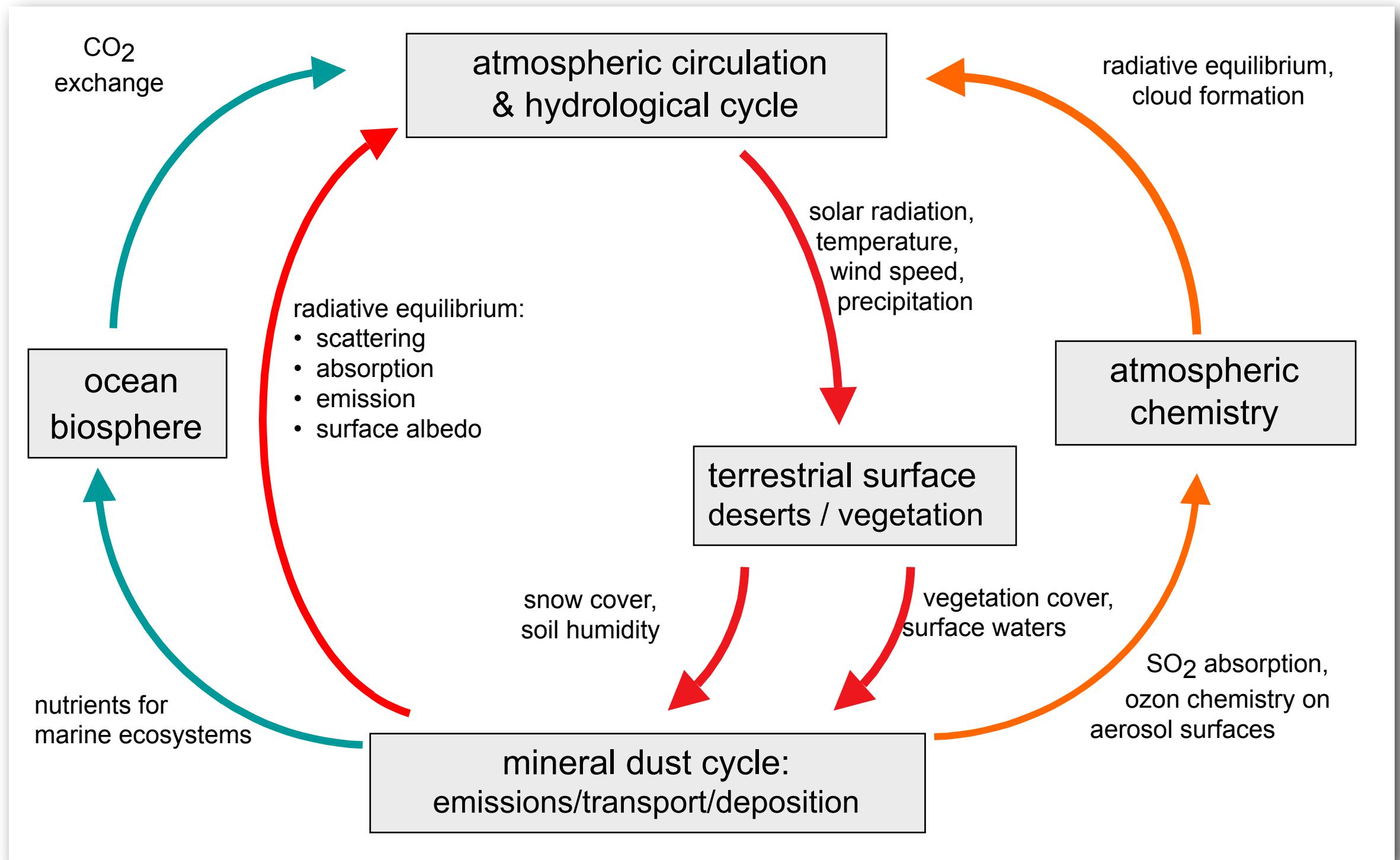


(DIRTMAP: Dust Indicator and Records of Terrestrial and Marine Paleoenvironments)

Dust in the climate system



Dust in the climate system



The glacial dust cycle: summary

- during LGM: 2- to 3-fold increase of dust cycle intensity (global value)
 - about 1/2 - 2/3 of glacial increase in dust emissions might be caused by the increased wind strength during the LGM
 - about 1/3 - 1/2 of the glacial dust emissions might stem from glacial-only source regions (change of glacial vegetation cover)
- mineral dust aerosol might be responsible for about 1/4 of the glacial temperature cooling in the (sub)tropical low latitudes
 - in higher northern latitudes these radiative effects are minor as compared to radiative changes caused by the glacial ice sheet
- model simulations with marine biogeochemistry models reveal that the glacial atmospheric CO₂ concentration may have been decreased by up to 40-50ppm
 - it seems unlikely that glacial dust input into the oceans is the only reason for the observed total glacial CO₂-reduction of ~80 ppm

Climate II

(Winter 2020/2021)

7th lecture:

Dust and Vegetation

(glacial dust increase, iron fertilisation, radiative forcing, vegetation and aridity)

End of lecture.

Slides available at:

https://paleodyn.uni-bremen.de/study/climate2020_21.html