

Climate warming `backward`

The last 100 Million years

Transitions from Greenhouse to Icehouse Climate

Climate System II

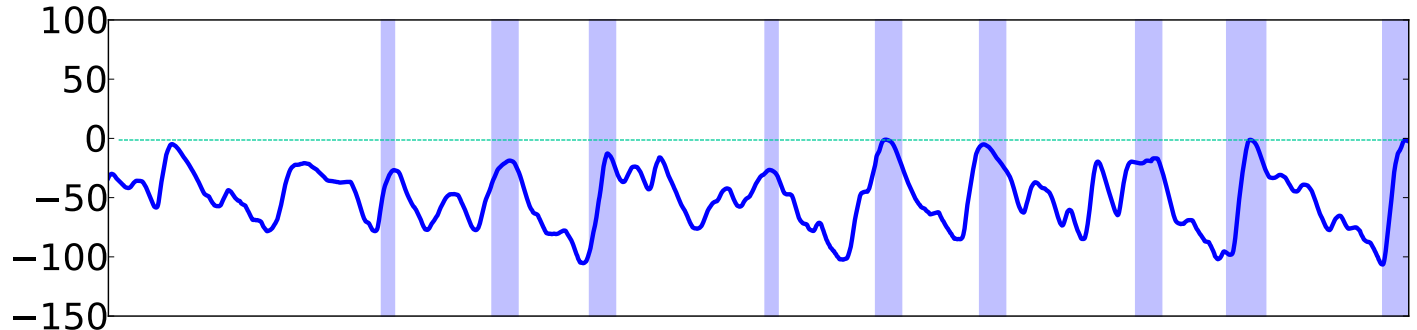
Gerrit Lohmann

with Christian Stepanek

Glacial-Interglacial variability

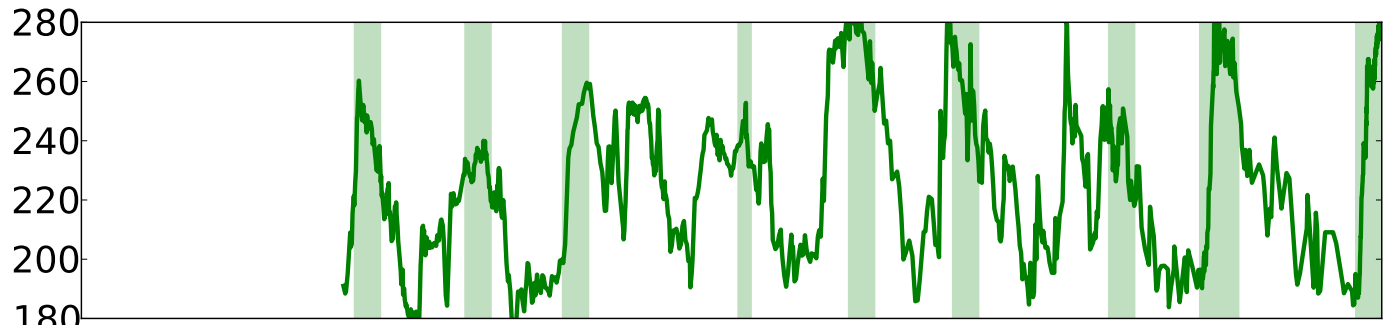
Global Sea Level [m]

(Bintanja et al., 2005)

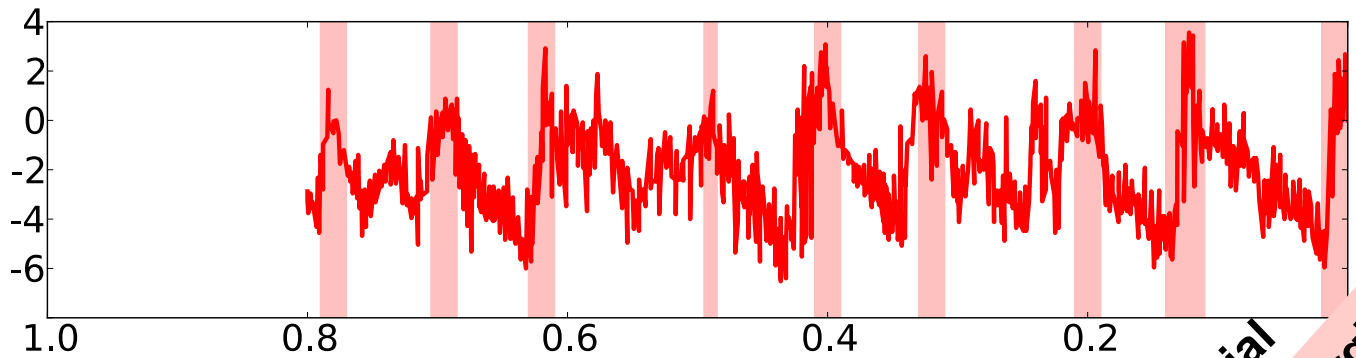


CO₂ [ppmv]

From ice cores (EPICA, 2009)



Temp. anomaly "O-18" [° C]



Million years

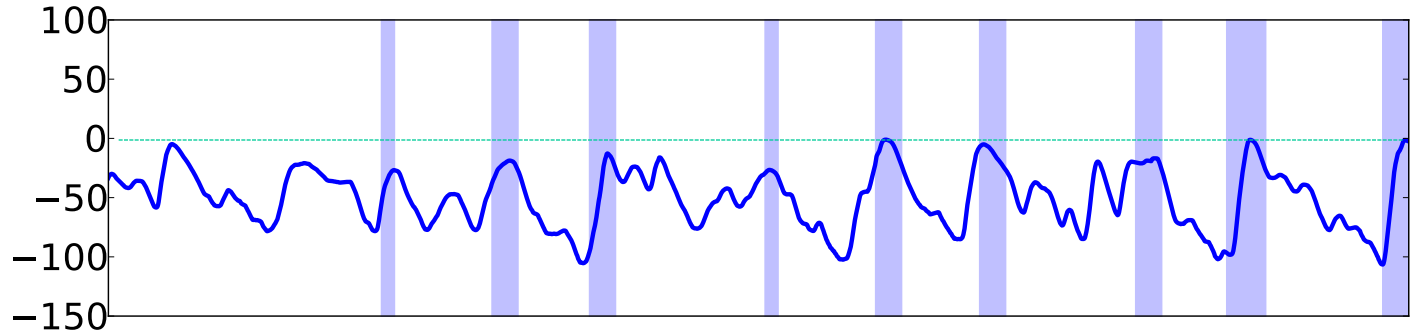
Glacial

Interglacial

Glacial-Interglacial variability

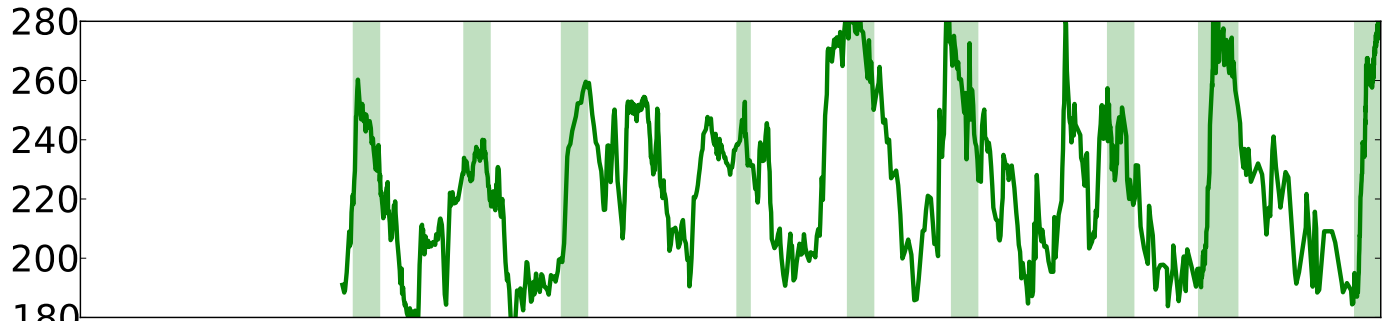
Global Sea Level [m]

(Rohling et al., 2009)

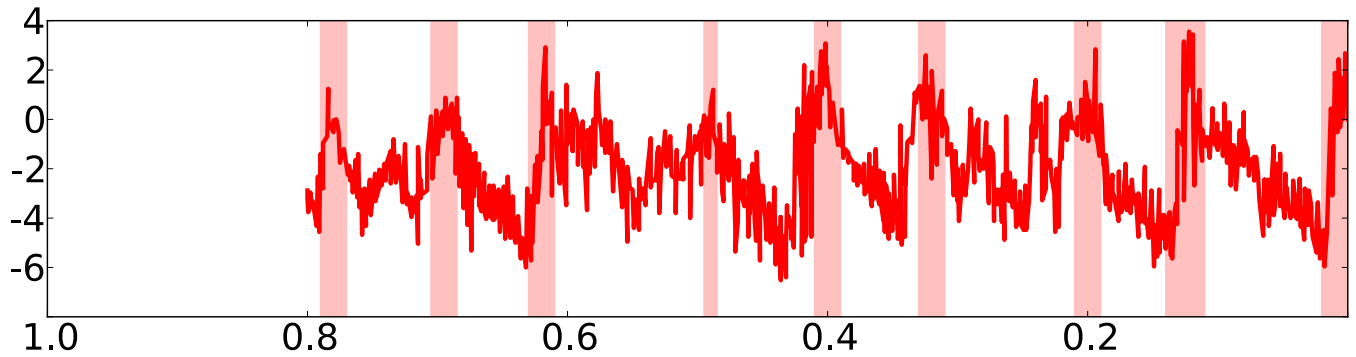


CO₂ [ppmv]

From ice cores (EPICA, 2009)

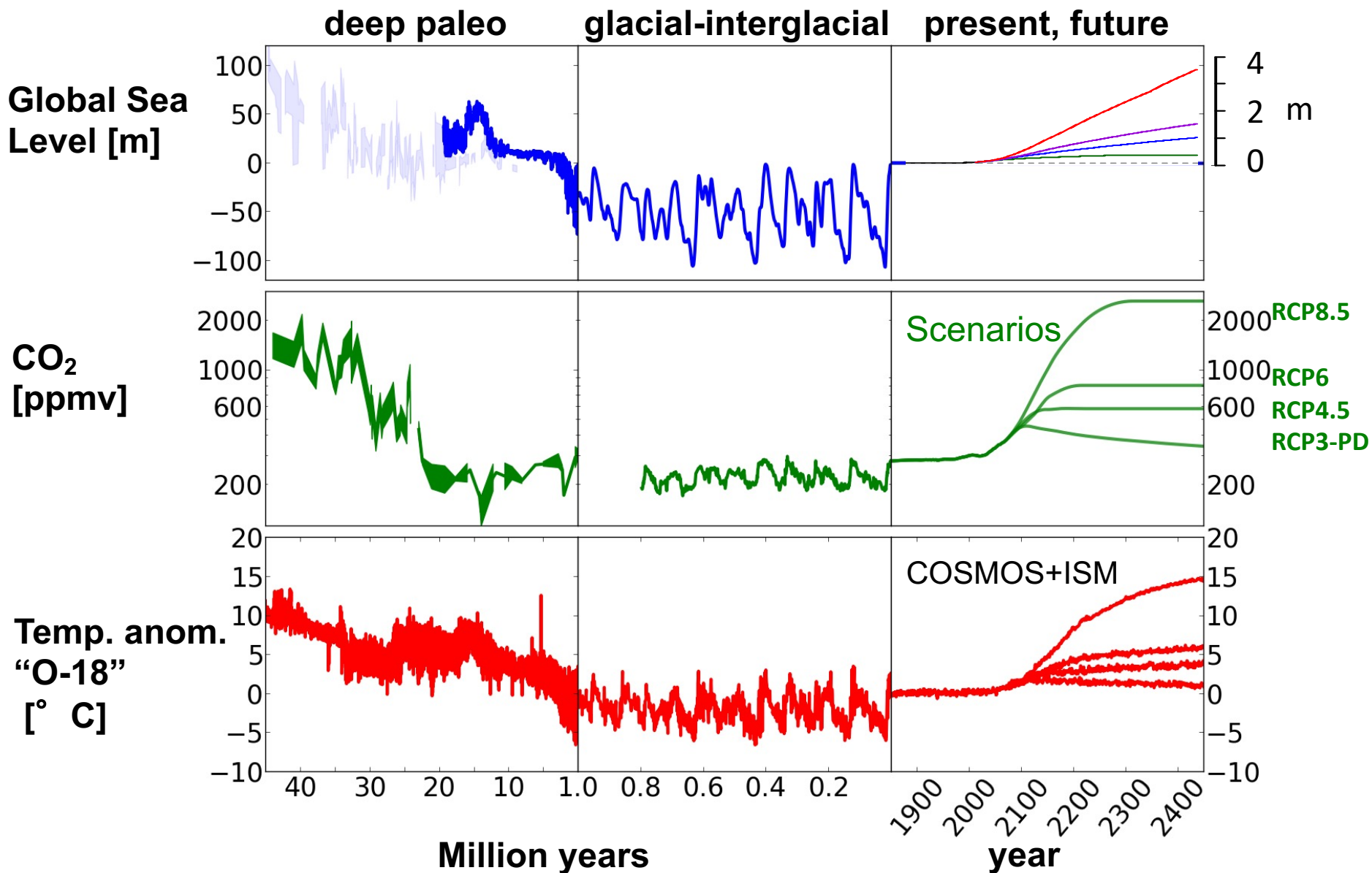


Temp. anomaly "O-18" [° C]



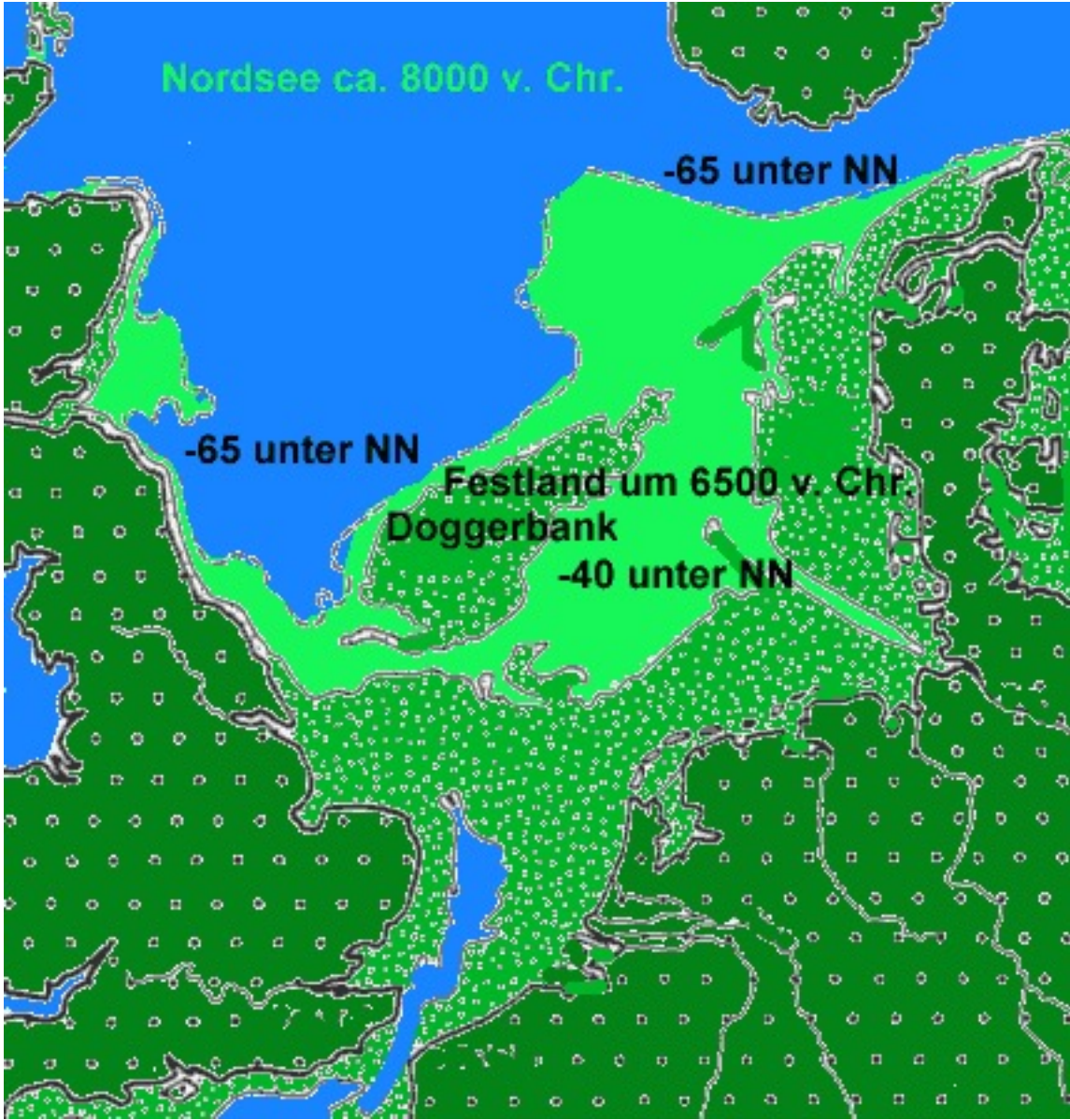
Million years

Natural variability and perturbed climate

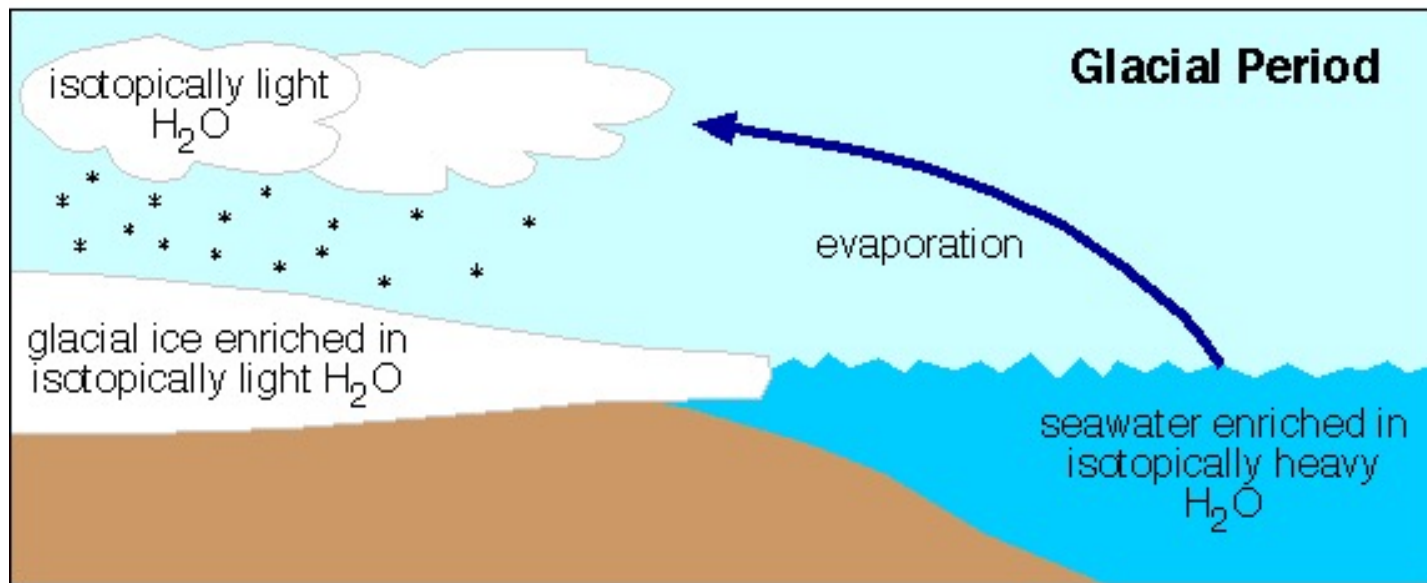
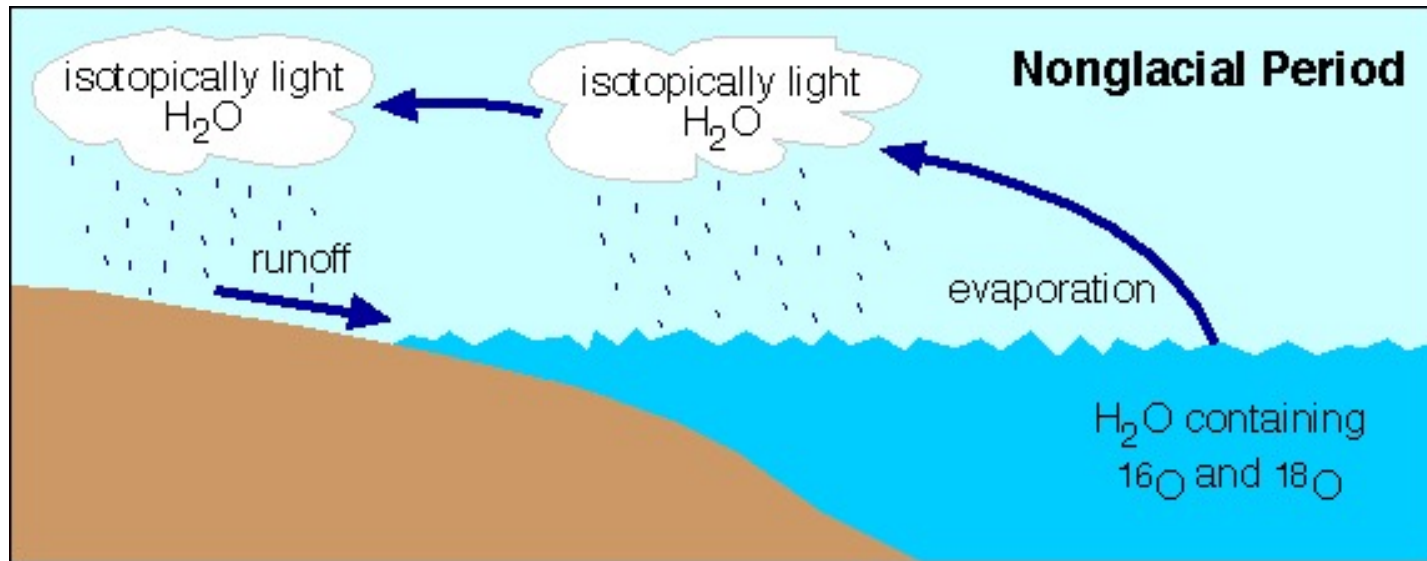


(Kominz et al., 2008; Pagani et al., 2009; Kramer et al., 2011; Crowley & Kim 1995, Wei & Lohmann,)

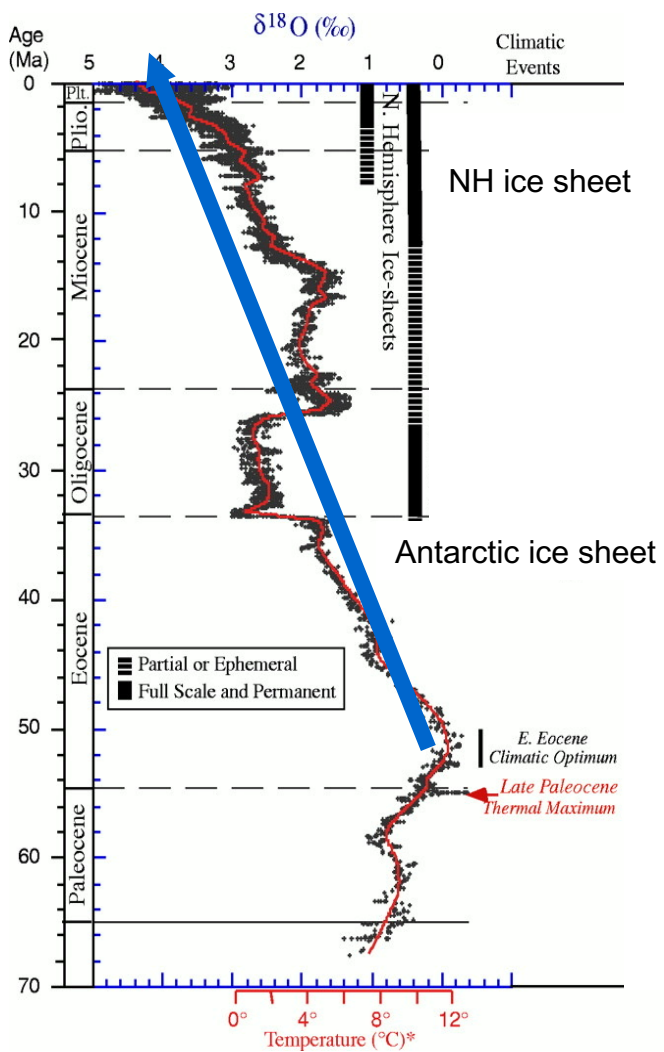
Sea level: 10 ky BP



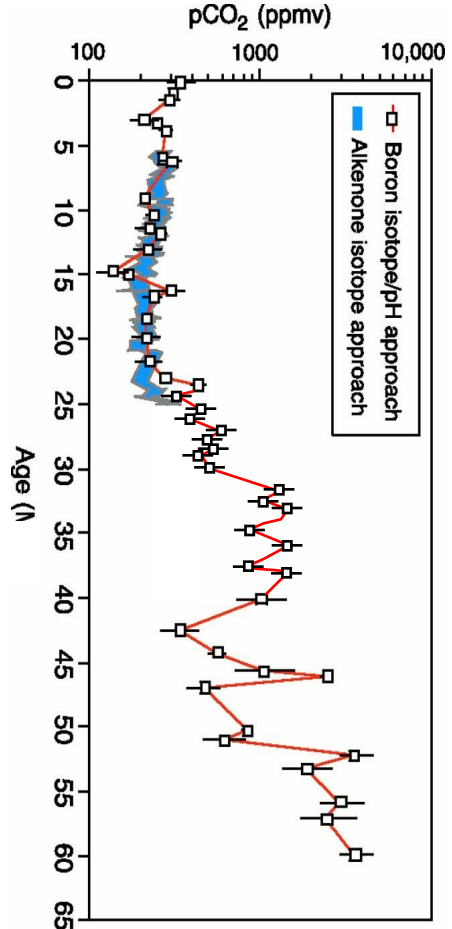
O-18 and sea level



Transitions from Greenhouse to Icehouse Climate: Evidence from Marine Sediments



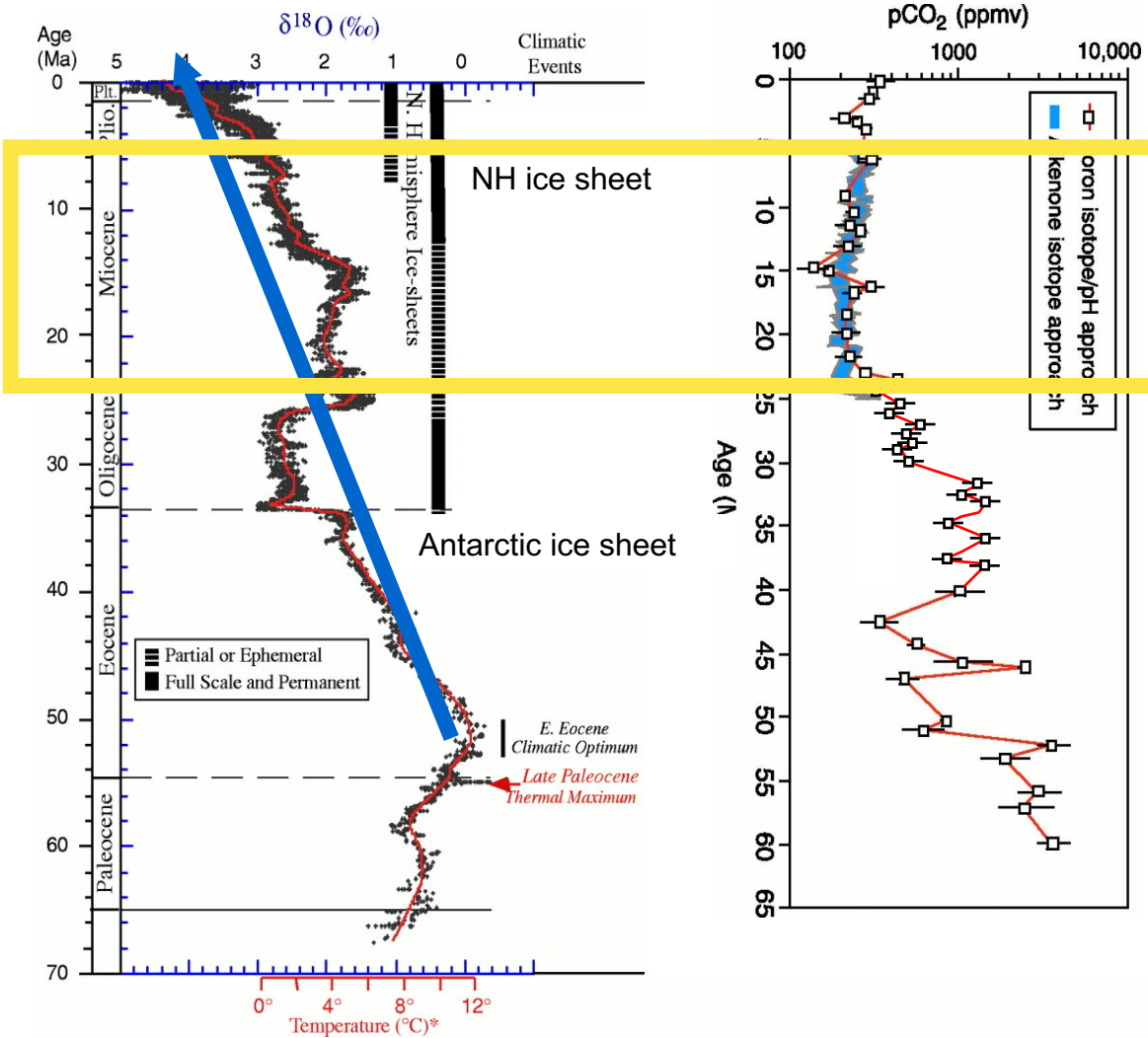
Global deep-sea O-18
(Zachos et al. 2001)



Proxy estimates of atmospheric pCO₂ (Pearson & Palmer 2000; Pagani et al. 1999, 2005)

Integrative approach
Data-Modelling

Transitions from Greenhouse to Icehouse Climate: Evidence from Marine Sediments

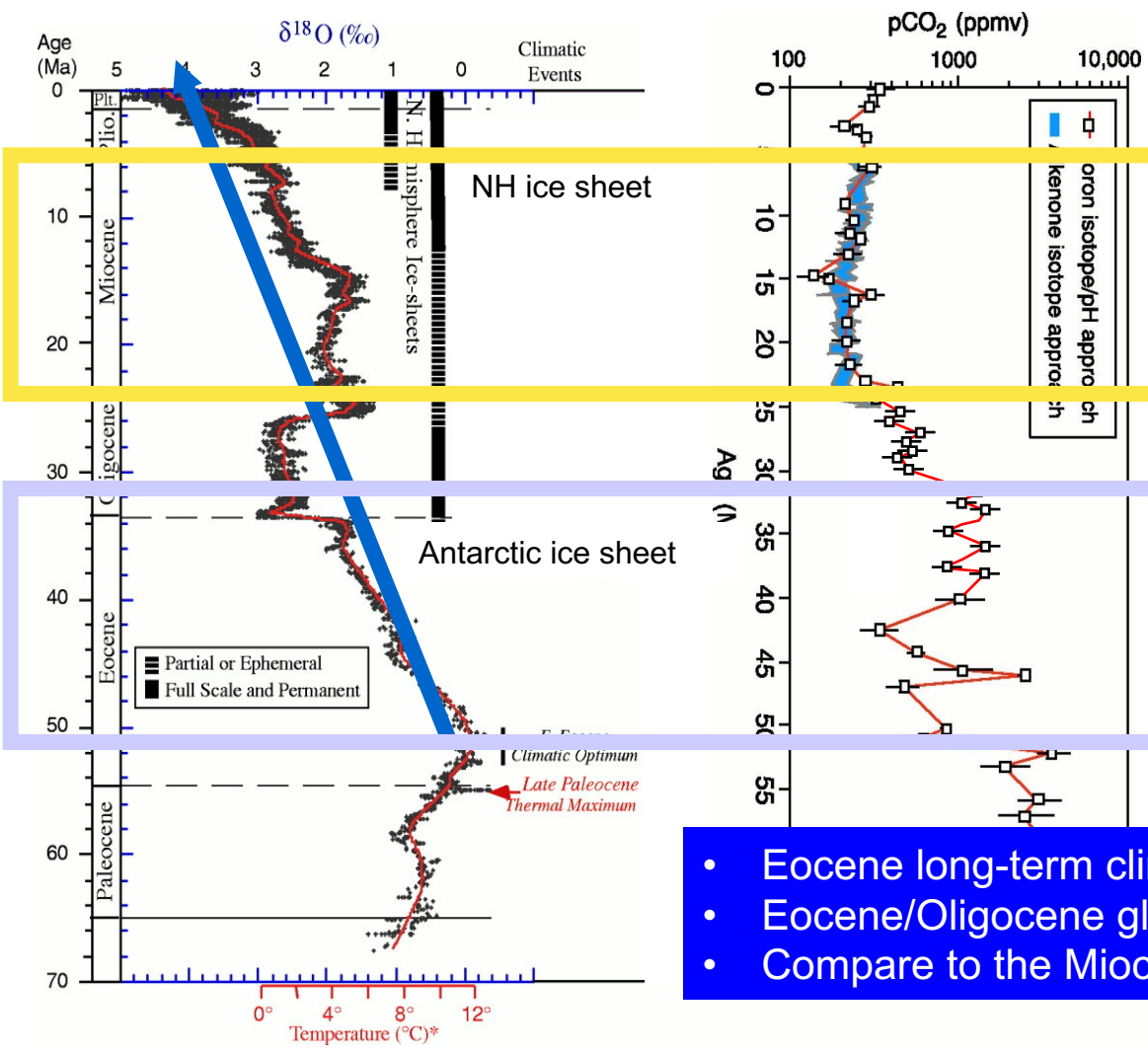


Miocene

Global deep-sea O-18
(Zachos et al. 2001)

Proxy estimates of atmospheric pCO_2 (Pearson & Palmer 2000; Pagani et al. 1999, 2005)

Transitions from Greenhouse to Icehouse Climate: Evidences from Marine Sediments



Miocene

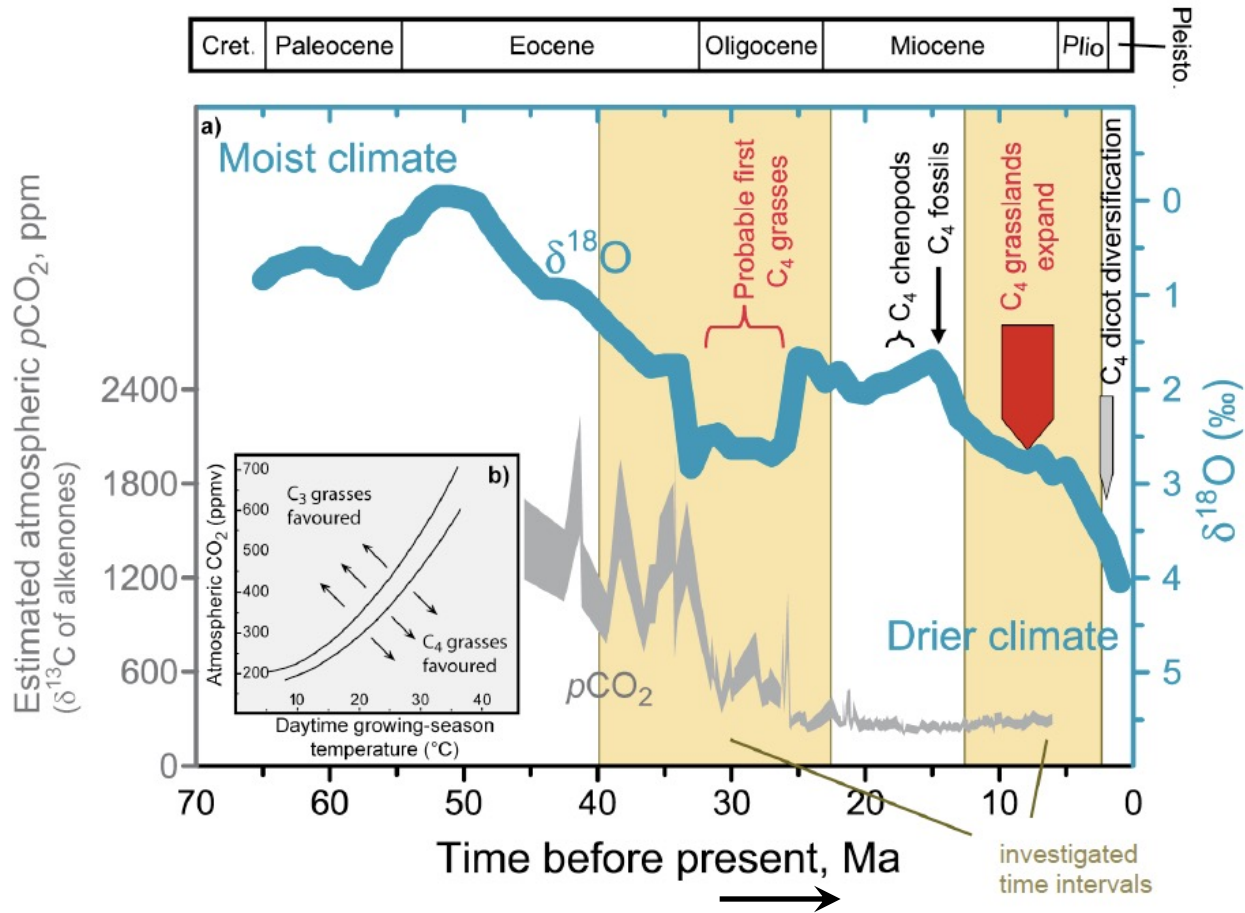
Eocene/Oligocene

- Eocene long-term climate cooling
- Eocene/Oligocene glaciation of Antarctica; drop in pCO₂
- Compare to the Miocene/Pliocene cooling; low pCO₂

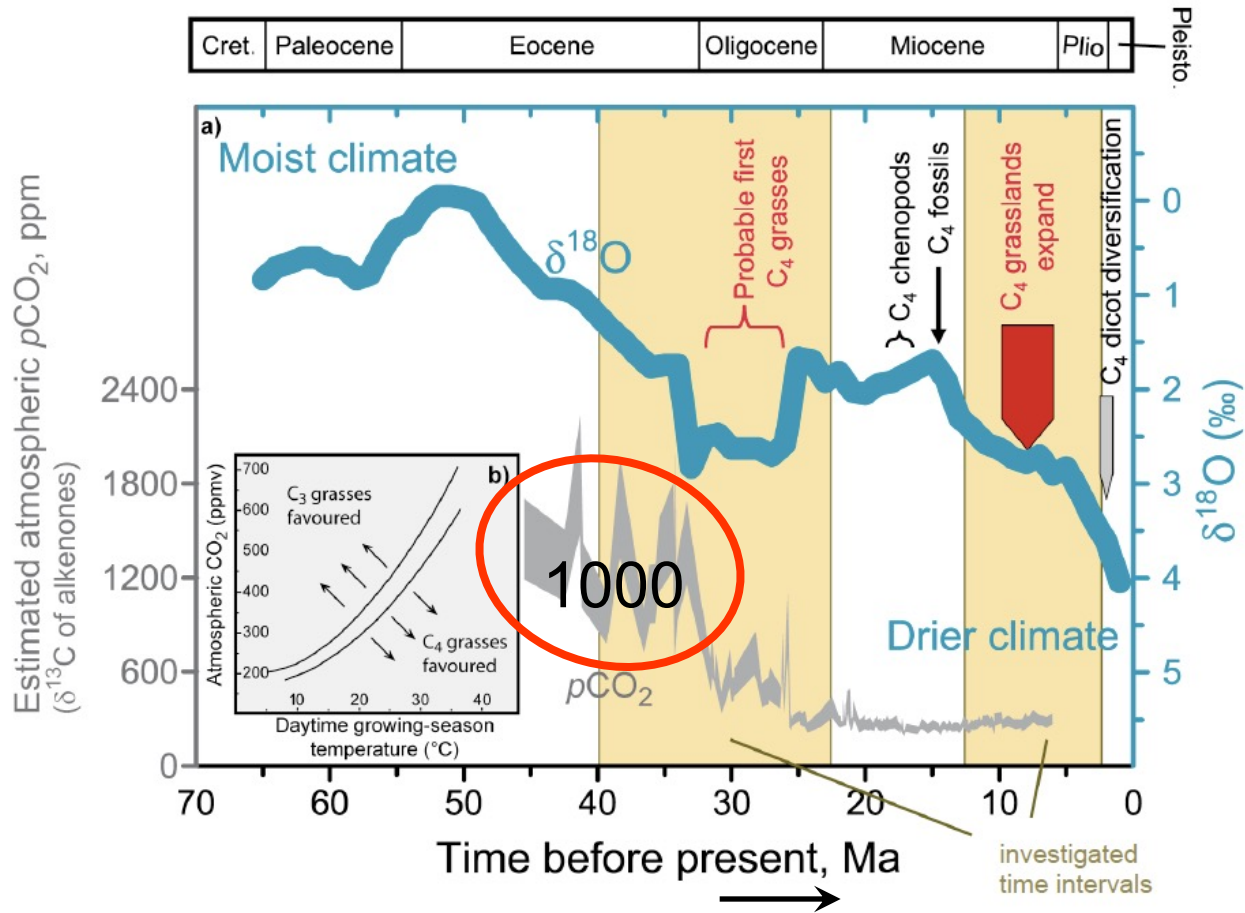
Global deep-sea O-18
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Proxy estimates of atmospheric pCO₂ (Pearson & Palmer 2000; Pagani et al. 1999, 2005)

Climate warming 'backward'

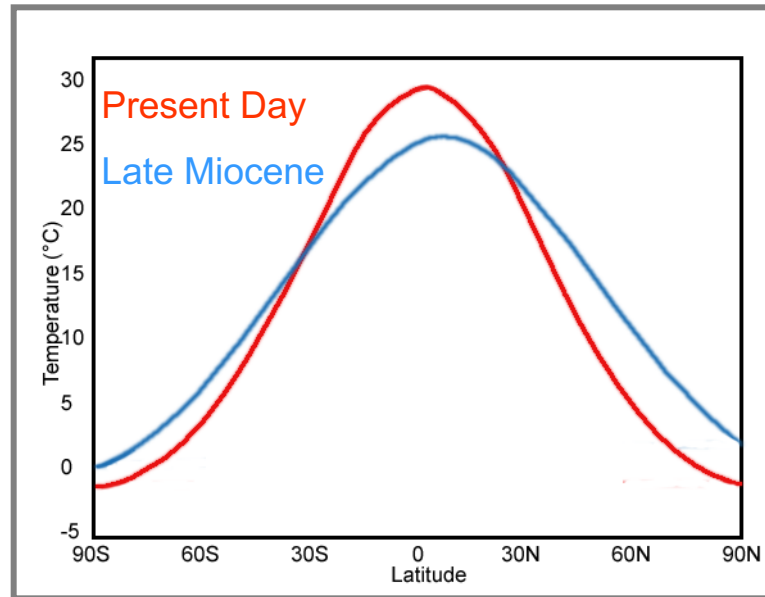


Climate warming 'backward'



Flat Temperature Gradient

Reason?



Crowley

Steppuhn et al., 2006

Many authors

The international journal of science / 2 April 2020

nature

POLAR OPPOSITE

Evidence for a temperate
rainforest near the South
Pole 90 million years ago

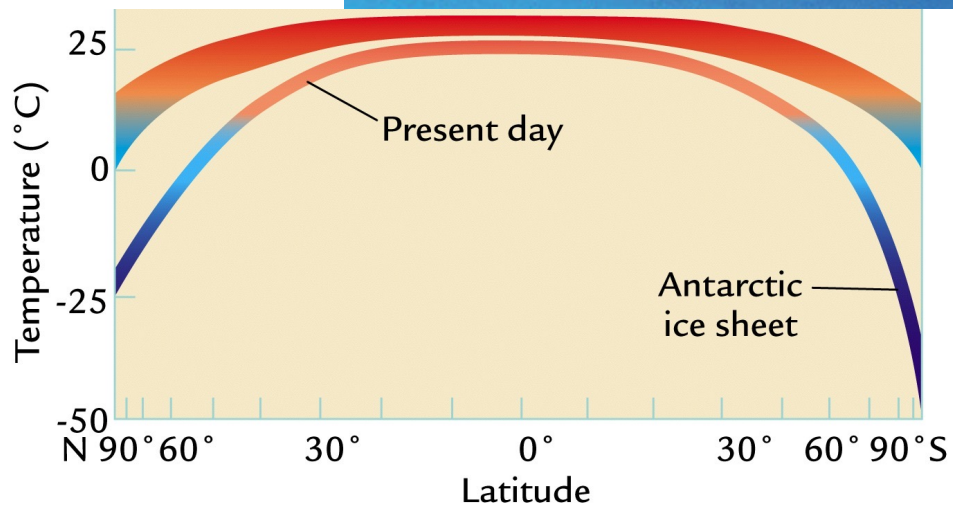
High stakes at sea
Will treaty to protect
ocean resources
hamper research?

Chemical switch
Light turns neutral
radical into powerful
reducing agent

Behind your back
Visualization of the
clock that governs
spine formation

WILEY
DISCOVER SOMETHING GREAT





OFFSHORE

Evidence for a temperate rainforest near the South Pole 90 million years ago

High stakes at sea
Will treaty to protect ocean resources hamper research?

Chemical switch
Light turns neutral radical into powerful reducing agent

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Visualization of the clock that governs spine formation

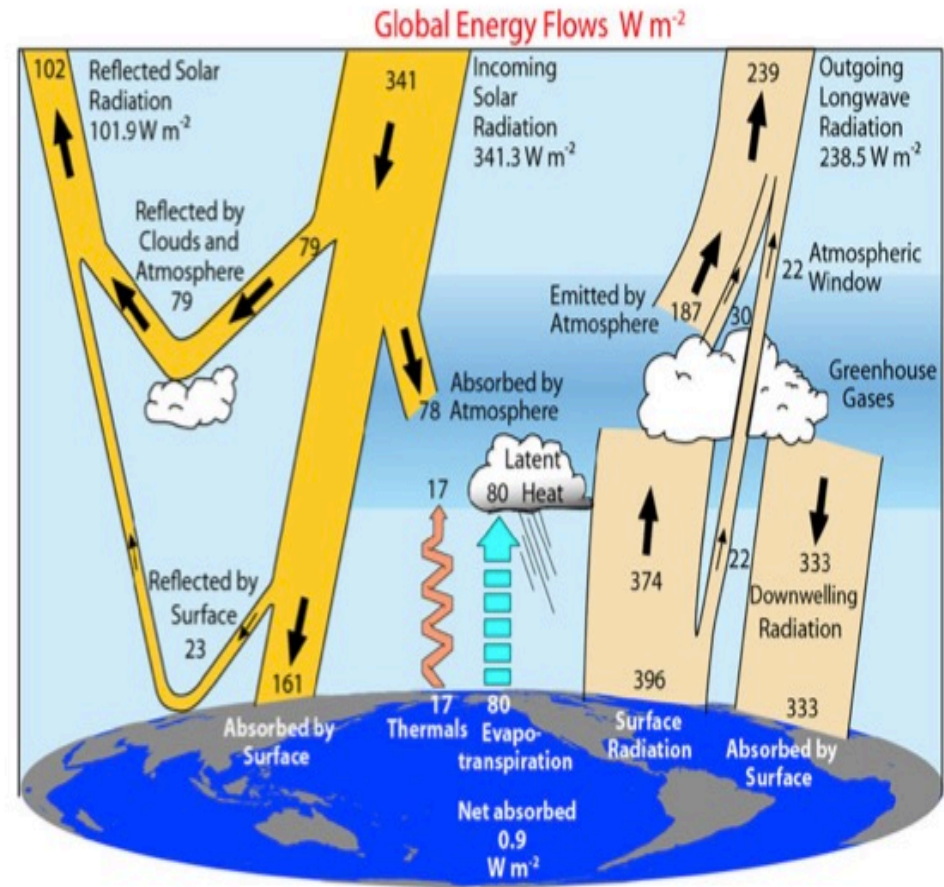
9 770026 033075 148

Energy Budget

- In steady state, energy follows energy balance model:

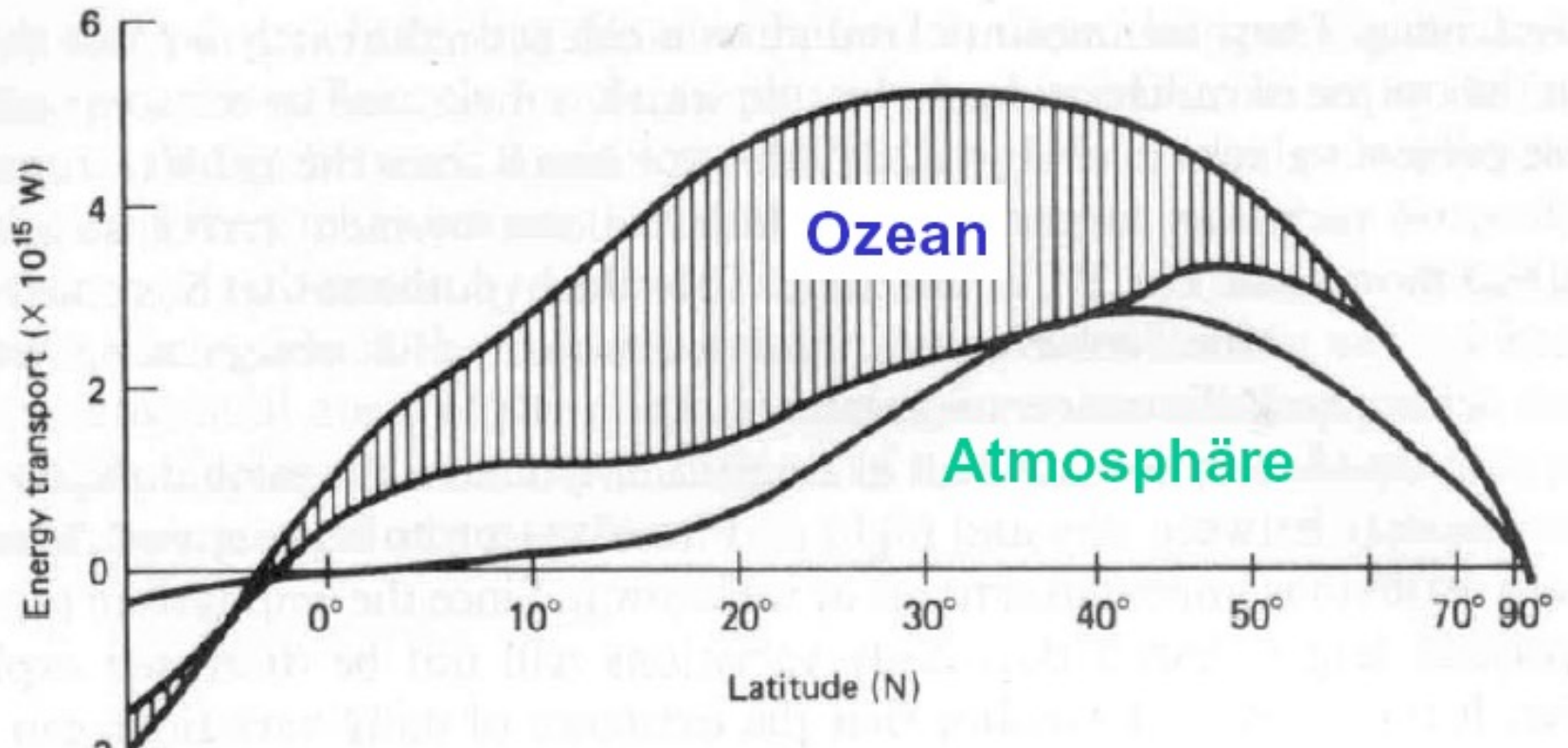
CHANGE IN STORAGE = IN – OUT

- many papers discuss an imbalance in this equation, which results in missing energy



(Trenberth & Fasullo, 2012)

Northward Heat Transport

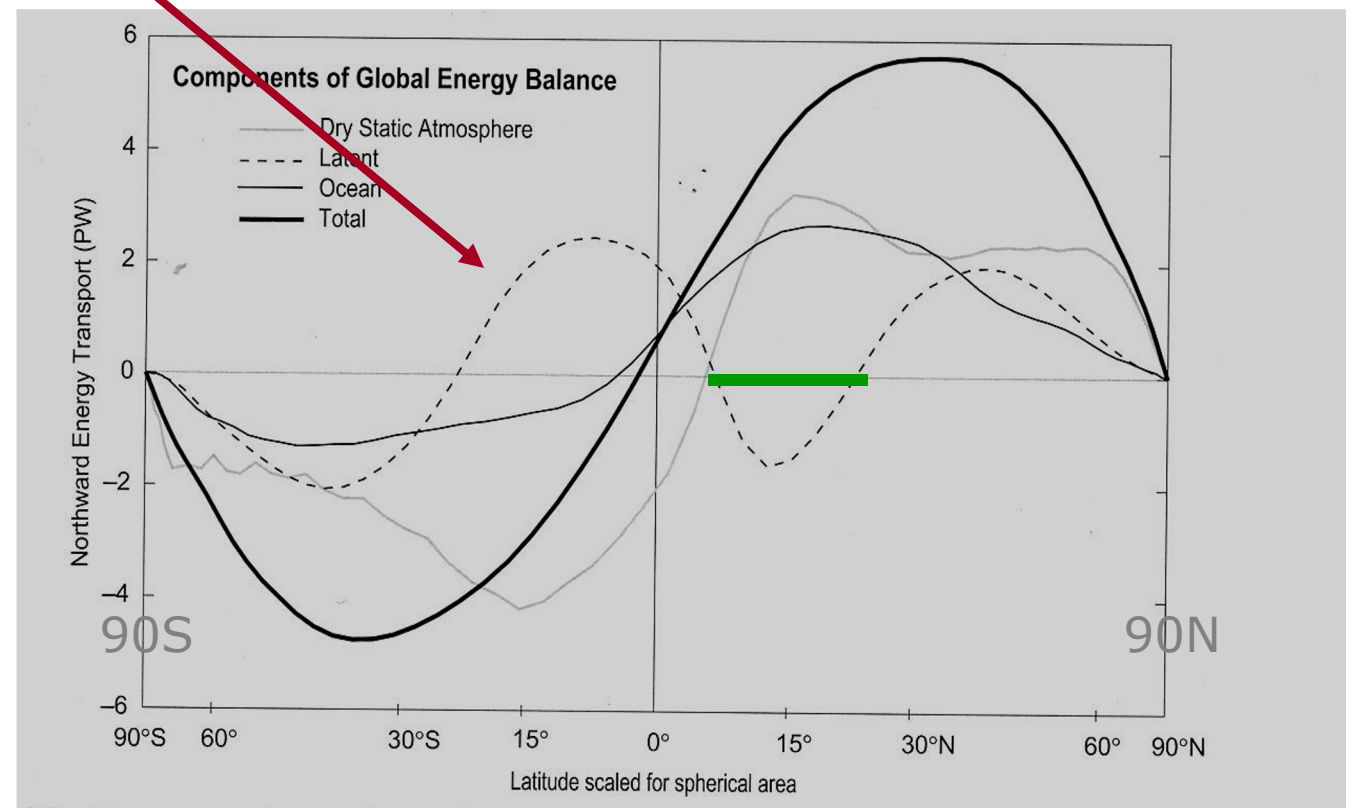


nach Von der Haar & Ort; Quelle: Gill

Global meridional heat transport divides roughly equally into 3 modes:

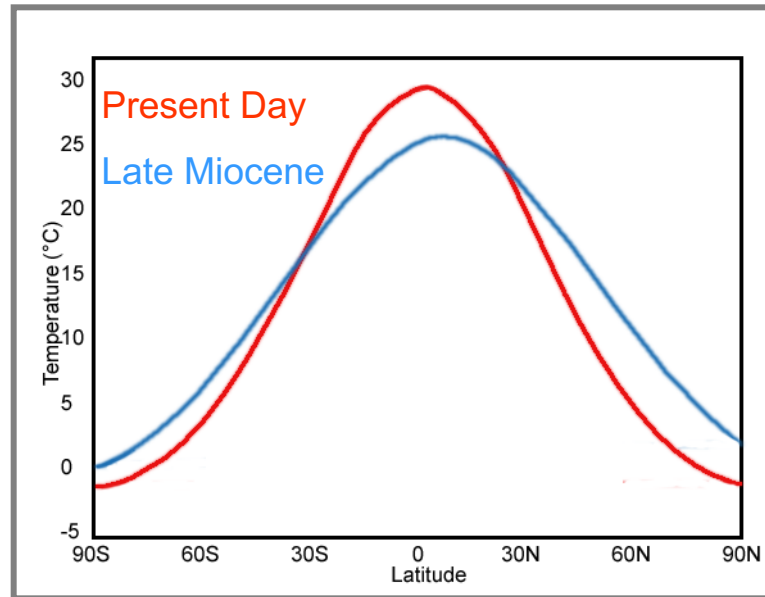
1. atmosphere (dry static energy)
2. ocean (sensible heat)
3. water vapor/latent heat transport

The three modes of poleward transport are comparable in amplitude, and distinct in character (sensible heat flux divergence focused in tropics, latent heat flux divergence focus in the subtropics)



(residual method,
TOA radiation
1985-89 and
ECMWF/NMC
atmos obs)

Flat Temperature Gradient



Crowley

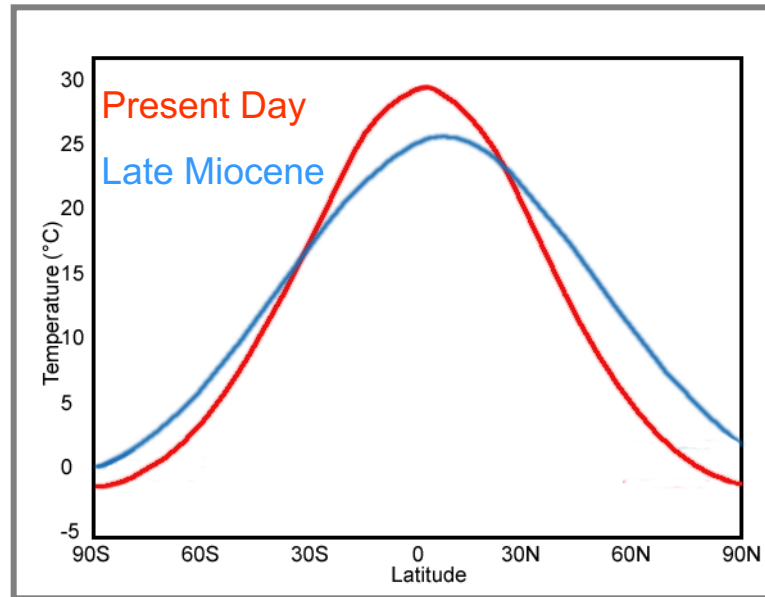
Steppuhn et al., 2006

Many authors

Reason?

- **Sensible heat transport**
- **Latent heat transport**
- **Ocean heat transport**
- **Orography** → **Greenland: high latitude warming**
- **Changes in the land surface cover**
- **Other effects?**

Flat Temperature Gradient






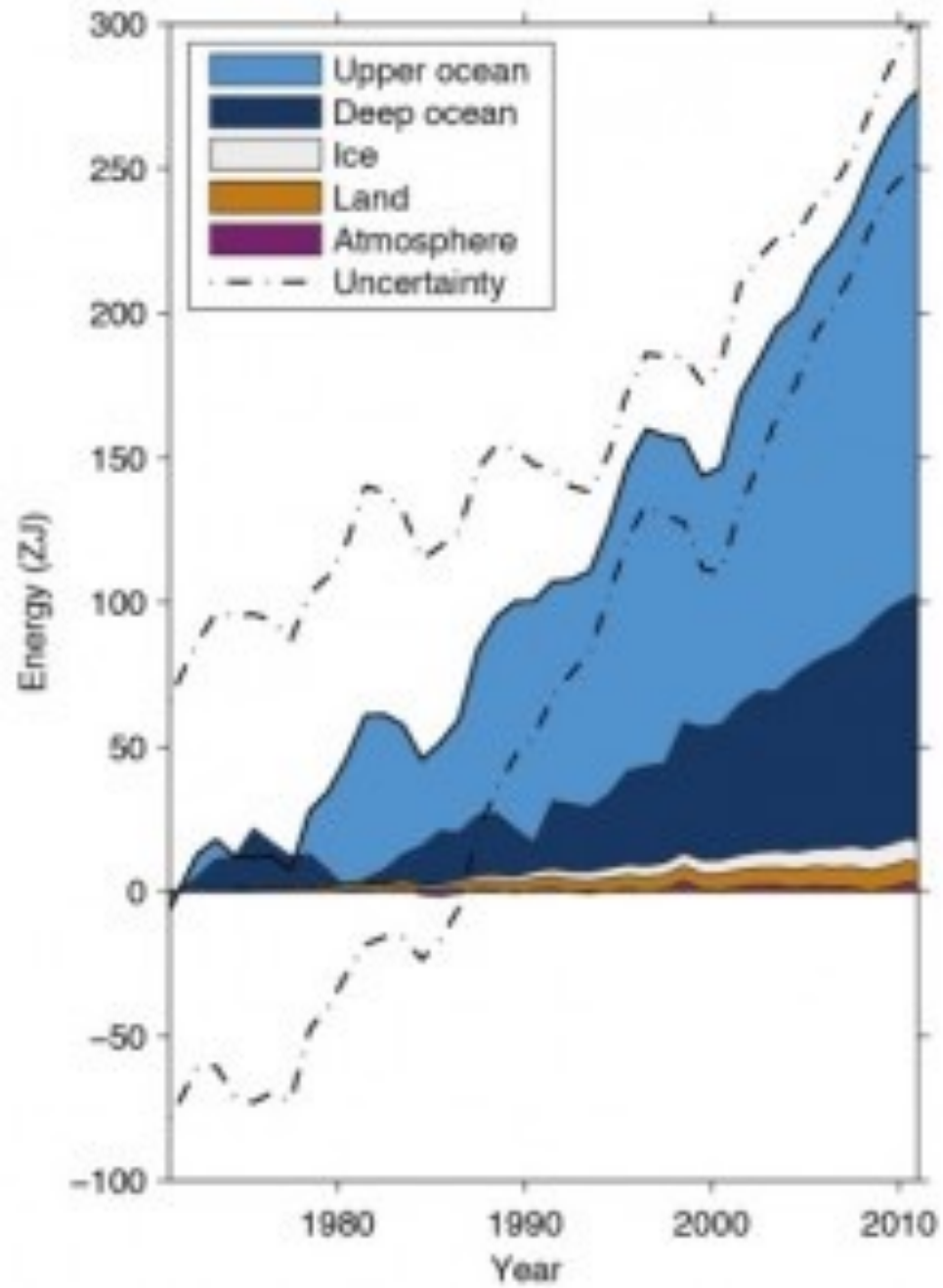
Crowley

Steppuhn et al., 2006

Many authors

Reason?

- Sensible heat transport $\sim \Delta T$  EBMs,
- Latent heat transport $\sim \Delta T \frac{\partial q}{\partial T}$  Caballero & Langen
- Ocean heat transport  Panama Gateway, Atlantic salinity reduction
- Orography → Greenland: high latitude warming
- Changes in the land surface cover
- Other effects?



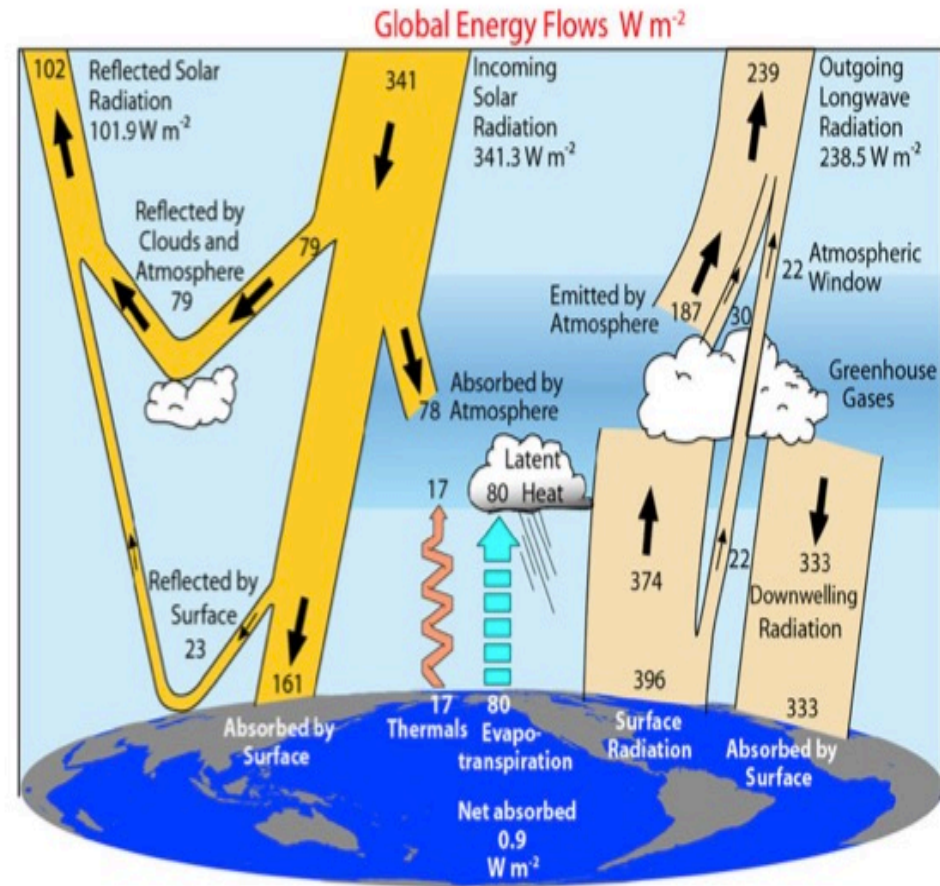
Our current
warming

boring for the hundredth time

Energy Budget

CHANGE IN STORAGE = IN – OUT

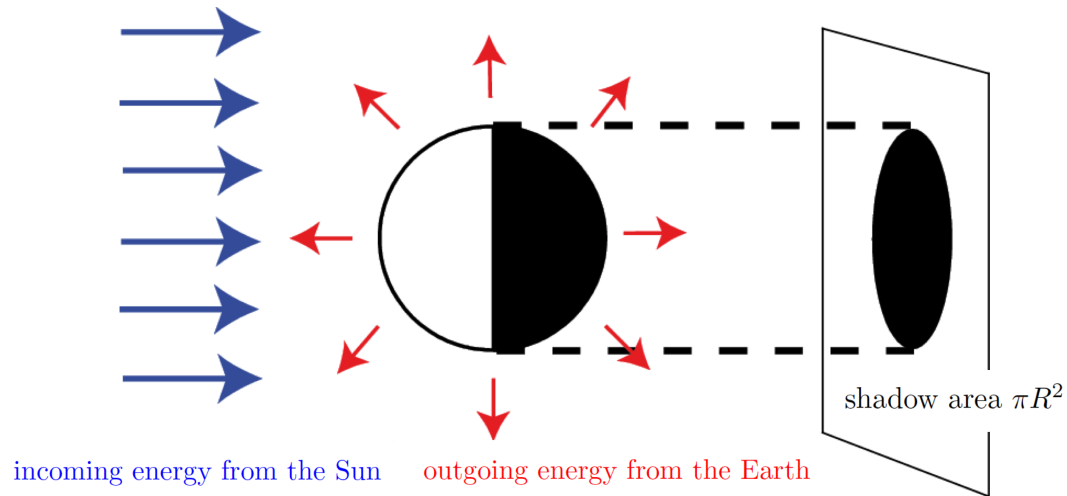
- many papers discuss an imbalance in this equation, which results in missing energy



(Trenberth & Fasullo, 2012)

Energy balance model

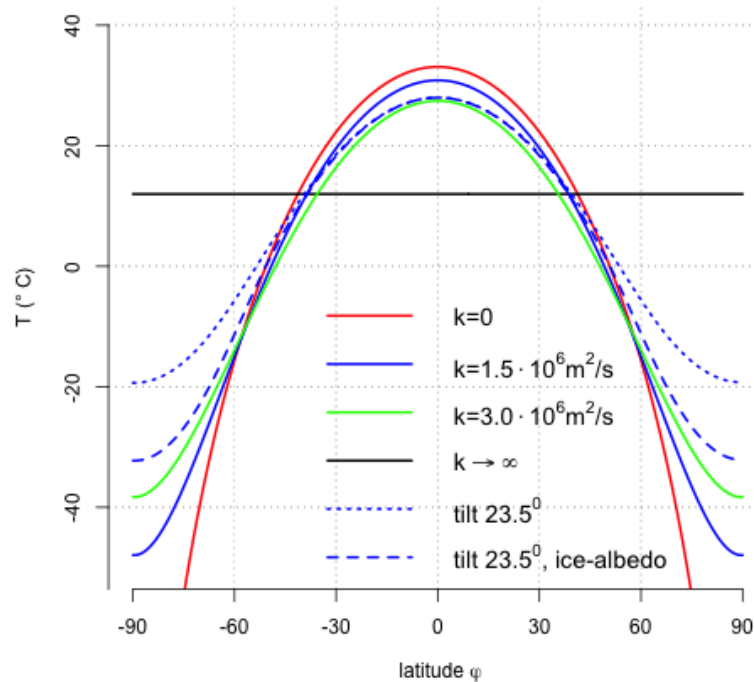
$$(1 - \alpha)S\pi R^2 = 4\pi R^2\epsilon\sigma T^4$$



$$T = \sqrt[4]{\frac{(1 - \alpha)S}{4\epsilon\sigma}}$$

boring for the hundredth time, but ...

$$C_p \partial_t T = \nabla \cdot HT + (1 - \alpha)S(\varphi, t) - \epsilon\sigma T^4$$

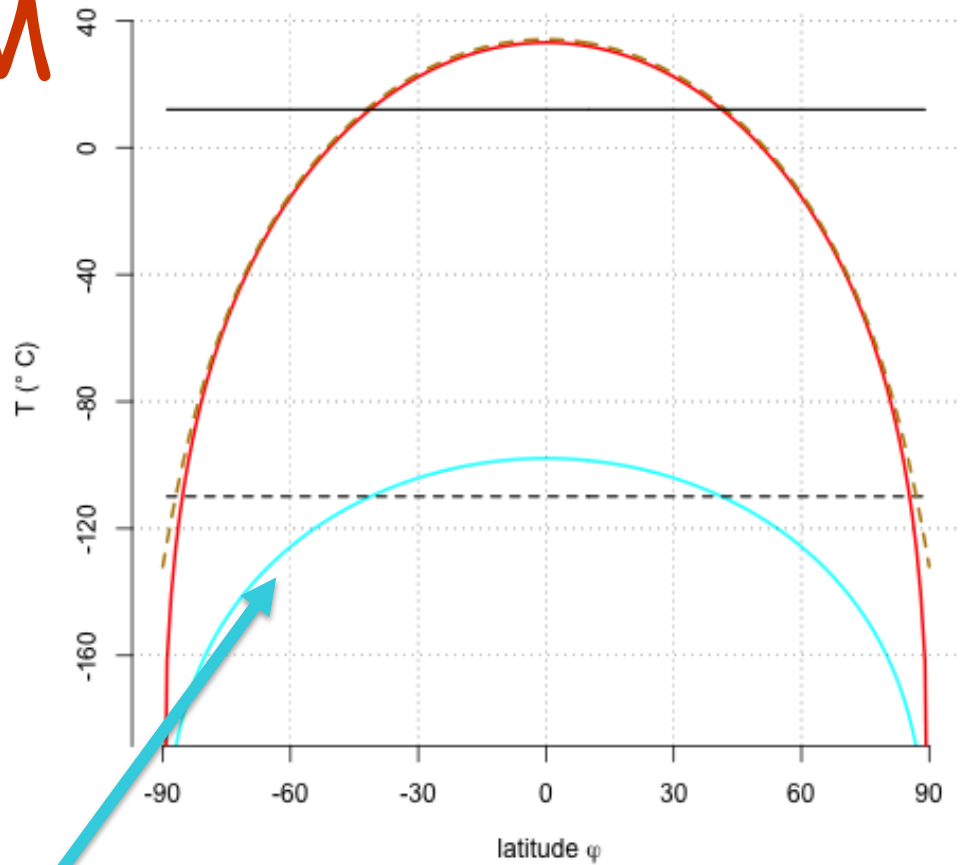
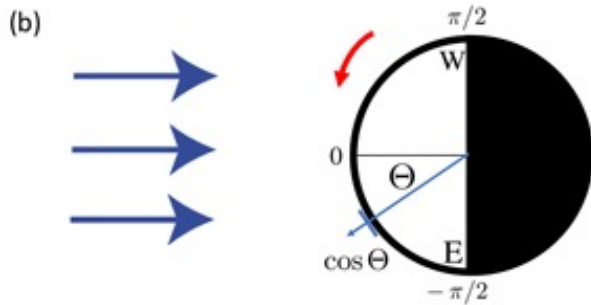
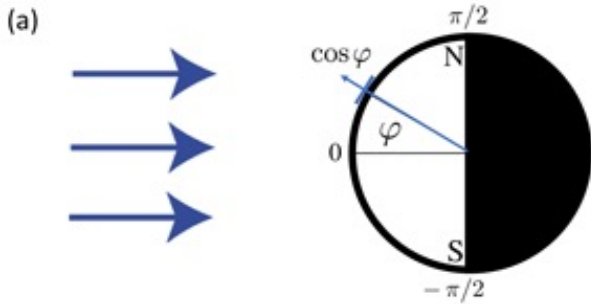


$$HT = -k\nabla T$$

Figure 4. Equilibrium temperature of (15) using different diffusion coefficients. The blue lines use $1.5 \cdot 10^6 \text{ m}^2/\text{s}$ with no tilt (solid line), a tilt of 23.5° (dotted line), and as the dashed line a tilt of 23.5° and ice-albedo feedback using the representation of Sellers (1969). Except for the dashed line, the global mean values are identical to the value calculated in (12). Units are $^\circ\text{C}$.

In the exercise, long-wave radiation as $A + BT$

EBM



$$\bar{T} = \frac{1}{2} \int_{-\pi/2}^{\pi/2} T(\varphi) \cos \varphi d\varphi$$

$$= \frac{1}{2} \frac{\Gamma(5/8)}{\sqrt{2\pi} \Gamma(9/8)} \cdot \sqrt[4]{\frac{(1-\alpha)S}{4\epsilon\sigma}} \underbrace{\int_{-\pi/2}^{\pi/2} (\cos \varphi)^{5/4} d\varphi}_{\sqrt{\pi} \Gamma(9/8) / \Gamma(13/8)}$$

$$= \frac{1}{2} \frac{1}{\sqrt{2}} \frac{\Gamma(5/8)}{\Gamma(13/8)} \cdot \sqrt[4]{\frac{(1-\alpha)S}{4\epsilon\sigma}}$$

$$\frac{\sqrt{2} 8}{4 \cdot 5} = 0.4\sqrt{2}$$

Heat capacity of the climate system

Fast rotation

Practical Jan 11, 2022

Exercise

EBM analysis

- <https://1drv.ms/u/s!AnZSDMNwdkDMgccDeuhjFFrmQHaqvw?e=ZaHqPA>